# Morogoro Domestic Water Supply Plan 

Volume IV
Hydrogeology-Annexes and Data

Final Report
August 1980

## IJHV

DHV Consulting Engineers

Kingdom of the Netherlands Ministry of Foreign Affairs DGIS

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ANNEX DA 1

THE GEO-ELECTRICAL METHOD

DA 1 THE GEO-ELECTRICAL METHOD

## General

The geo-electrical method is described in many handbooks and reports, to which the reader is referred (see references sub-par. 2.4., No. 5, 11, 17, 27, 29 and 50). Here just some aspects will be mentioned.

All matter has the intrinsic property of an electrical resistivity ( $\rho$ ) , which is expressed in ohmm ( $=\int \mathrm{m}$ ). Therefore each earth layer has also this property which is usually called the formation resistivity ( $\rho_{f}$ ). When the pores of the formation contain water, the following relation between the formation resistivity and the resistivity of the pore water $\left(\rho_{\mathrm{w}}\right)$ has been found.

$$
\begin{equation*}
\rho_{\mathrm{f}}=\mathrm{F} \times \rho_{\mathrm{w}} \tag{1}
\end{equation*}
$$

F is called the formation factor. Its value depends among other things on:

- porosity
- mineral content, e.g. clay and/or sand
- grain size, shape and arrangement of the particles.

For clayey layers the formation factor depends also on the water resistivity.
The resistivity values of different formations can vary from less than 5 ohmm up to more than 1000 ohmm. Due to the occurrence of these high contrasts between the formation resistivities, the geo-electrical method is often very suitable to distinguish the different layers.

In general the conductivity of water (EC) is given. The relation between the conductivity (in $\mathrm{mS} / \mathrm{m}$ ) and resistivity (in ohmm) is:

$$
\mathrm{EC}=\frac{10^{3}}{\rho}
$$

When there is no water in the formation the current conductivity is not electrolytic and the formation resistivity is always very high ( $\rho_{f}>500$ ohmm). In fissured bedrock the current conductivion takes place through the water in the fissures.

From formula (1) it is clear that when the salinity of the formation water increases, the water resistivity as well as the formation resistivity decrease. Moreover, if the clay content of a formation increases, the formation factor and with this the formation resistivity decrease. It has been found, just because of these two effects, that a distinction in the formation resistivities can be made between water-bearing layers with fresh water on the one hand and clayey layers or saline waterbearing layers on the other hand.


Geo-electrical sounding in Schlumberger arrangement of a subsurface built up by two layers with formation resistivities $\rho_{1}=100 \Omega \mathrm{~m}$; $\rho_{2}=$ 2 sm .

1, 2 and 3: three steps of a sounding. Full curves indicate current lines; dashed curves indicate equipotential lines.

The 4th figure gives the resulting sounding graph, with the interpretation in the lower part.

Fig. DA 1-1: Principle of geo-electrical sounding technique.

As in hydrogeologic studies the occurrence of fresh water-bearing layers is very important, the geo-electrical method is often applied to trace these layers. It is also clear that it is very difficult or even impossible to distinguish clayey layers from saline aquifers.
By the execution of measurements on the groundlevel the true formation resistivity of the successive layers cannot be measured directly. Carrying out surface measurements (a sounding) only the so-called apparent resistivity ( $\rho_{\mathrm{a}}$ ) can be determined. Such a sounding has to be interpreted to true formation resistivities.

## Field technique

A geo-electrical sounding is carried out by means of an arrangement comprising four electrodes. For the Schlumberger configuration used in the present survey, the electrodes are placed in a straight line, symmetrical in respect to a central point (fig. DA 1-1). Via the outer electrodes $A$ and $B$ an electrical current is passed through the soil. This generates a potential distribution inside the earth and therefore also a potential difference between the two inner electrodes $M$ and $N$. This potential difference (V) and the strength (I) of the current applied is measured. The apparent formation resistivity value is calculated then by using the formula:

$$
\rho_{a}=\frac{\pi}{4} \frac{L^{2}-a^{2}}{a} \frac{V}{I}
$$

where $a$ and $L$ are the spacings between $M$ and $N$, respectively $A$ and $B$. This calculated resistivity value is plotted against the value of $\mathrm{L} / 2$ belonging to it.
A sounding consists of a series of observations of current strengths and potential differences with varying electrode distances.
The distance between the current electrodes ( $L$ ) is at first usually 3 metres. The potential electrodes are then placed at a distance (a) of 1 metre from each other. The distance between the current electrodes is subsequently increased in steps. As the arrangement expands, the current penetrates more and more deeply into the earth. As a result, more and more strata will influence the measurements. The distance between the potential electrodes remains unchanged, unless $V$ becomes too small to be accurately read. Then the spacing between the potential electrodes is increased, after which the expanding of the $A B$-spacing is continued. The maximum distance between the current electrodes is determined by the depth to which information is desired.
In this way the relation between the horizontal distance $L / 2$ and the apparent formation resistivity $\rho_{a}$ is determined at several discrete points.

## Interpretation

After a sounding is carried out a smooth line is drawn which fits the observation point as well as possible. It is often possible then to make immediately a qualitative interpretation of the soundings. This means that often a qualitative prediction can be given concerning the layer structure, i.e. the occurrence of clayey and/or saline water-bearing layers (low resistivities), prospects for sandy or gravelly aquifers with fresh water (medium to high resistivities), the depth to the bedrock (very high resistivity) etc.
The quantitative interpretation of the soundings is a much more complicated problem. The interpretation is carried out in two stages.

- In the first stage the surface measurements have to be converted into a vertical sequence of layers with each a different formation resistivity, so:

$$
\rho_{\mathrm{a}}(\mathrm{~L} / 2) \rightarrow \rho(z)
$$

- There upon these interpreted layers have to be translated in hydrogeological terms, so:
$\rho(z) \rightarrow$ hydrogeological sequence of layers
The first stage of the interpretation is done by a curve-fitting method. For this purpose three-layer master curves 1) are used and a programmable pocket calculator (HP 97).
To execute the second stage, the relation between the formation resistivity values and the hydrogeological strata, which holds for the area under investigation has to be known.
The use of the geo-electrical method is limited, which is mainly caused by the following factors (see also sub-par. 3.2.3.):
- interpreting a sounding, it is assumed that the layers are horizontal and homogeneous over a rather large area;
- non uniqueness of interpretation; this means that one sounding may represent different sequences of resistivity layers (see the examples presented in sub-par. 3.2.3.);
- the actual sequence of layers has to be (strongly) schematized, due to the restricted resolution of the geo-electrical method;
- different hydrogeological strata may have the same value of formation resistivity (see table D 3.2.-2) .

Due to these limiting factors, a sounding has not to be regarded as an isolated one. Interpreting a sounding all the available data of neighbouring boreholes and all the information of the nearby soundings have to be taken into account. Moreover, it is in general very helpful to have in advance a conception of the hydrogeological structure to start with. This conception has to be based on the available hydrogeological data, such as borehole data.

1) Rijkswaterstaat, the Netherlands (1969). Standard graphs for resistivity prospecting. European Association of Exploration Geophysicists, The Hague.

## Approach of a geo-electrical investigation

In areas where little is known about the hydrogeological structure no final interpretation of the soundings can be given directly.
In the first instance the soundings can be interpreted then only tentatively into simple resistivity models. Because the interpretation is non-unique, the actual sequence of layers may be much more complicated than the simple models. In an interpreted model several actual layers are taken together and replaced by one "substitution" layer which has an "average" resistivity value.
After an evaluation of the tentative interpretations of all the soundings some locations are chosen where exploratory boreholes have to be drilled in which resistivity well logging measurements have to be executed. Based on these required informations the relation between the formation resistivity and the hydrogeological strata can be established. Moreover, on the ground of the boreholes and the tentative interpretations of the soundings a conception about the hydrogeological structure can be made. Backed with all this knowledge the final interpretation of the soundings can be made and finally a coherent model of the hydrogeological structure of the area and the prospects for the occurrence of fresh ground water can be presented.

## ANNEX DA 2

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## DA 2.1. SUMMARY

Within the scope of the hydrogeological survey detailed knowledge was required about the occurrence of shallow ground water and its quality. This information could be obtained from hand-drilled boreholes, the execution of which is time consuming.
Geo-electrical measurements have been carried out and compared with hand-drilled borehole data, to investigate the possibilities of using this method as an expedient in the investigations of shallow ground water and to optimize the selection of suitable shallow well sites.

The use of the geo-electrical method in the investigations for shallow ground water appeared to be possible, although with a low accuracy. By means of a qualitative interpretation of the sounding curves only a general impression can be obtained about the occurrence and water quality (fresh versus saline) of shallow ground water. Optimization of the selection of shallow well sites does not seem feasible with this method because of its low accuracy and the time required for execution and interpretation; this apart from the fact that this method requires a skilled geophysicist/operator, labourers, equipment and transport.

## DA 2.2 INTRODUCTION

Within the scope of the hydrogeological survey of the MDWSP, detailed knowledge was required about the occurrence, the quantity and the quality of shallow ground water. The required data have been obtained from existing boreholes, shallow wells, pools and hand-dug holes.
It has been found, however, that often additional information is necessary. This information could be collected among other things from handdrilled boreholes, the execution of which is time consuming.
In July 1978, the Morogoro Wells Construction Project (MWCP) started to operate in the Morogoro Region. Therefore, a skilled siting crew, from the Shinyanga Shallow Wells Project (SSWP), executed a survey to locate about fifty suitable shallow well sites in the period from May up to September 1978.
Owing to these circumstances it has been decided to study the following questions:

- to what extend is it possible, by means of the geo-electrical method, to predict the occurrence of shallow water-bearing layers, their thickness and the quality of the water;
- is it possible by means of the geo-electrical method to optimize (that is to speed up and to economize) the siting of shallow well locations.

This study has been integrated in the running programme of the siting crew. In this way the results of the geo-electrical measurements could be compared directly with the data of the hand-drilled boreholes such as sample descriptions, resistivity well-logging measurements and the quality of the ground water expressed in EC-values.

The survey area was situated in the northern part of the Morogoro Region along the northern border of the Wami valley. The investigations have been carried out near villages along the road between Dumila and Turiani (see e.g: map D1).
For a hydrogeological description of the survey area the reader is referred to par. 4.1.

## DA 2.3 APPROACH TO THE STUDY

The investigations have been performed in two stages.
In the first one, which took place near the villages between Dumila and Mvomero, the geo-electrical group followed the SSWP siting crew. A geo-electrical sounding was carried out near most of the hand-drilled holes. In this way experience was gained concerning the types of stratification occurring in the investigated area and the corresponding sounding curves.
In the second stage the geo-electrical group preceded the hand-drillers.
In this stage the villages between Mvomero and Turiani were visited together with the SSWP hydrogeological assistant. Pursuing the SSWP procedure (see DA 2.4.), in each village areas were located which showed good prospects for shallow well sites. In every promising area one or more geo-electrical soundings were executed. On the basis of these soundings the decision has been made whether the location should be considered for further investigations by means of hand-drillings or not. All these locations have been shown to the surveyor of the siting crew after which the siting has been completed with hand-drillings, EC measurements, well loggings and if necessary a yield test. Then the comparison between the geo-electrical soundings and the boreholes could be carried out and an evaluation could be made.

DA 2.4 THE EXECUTED STUDY

## DA 2.4.1 Working method of the SSWP siting crew

Because one of the objectives of the study was to investigate the possibilities of optimizing the selection of shallow well sites and because the execution of the study has been integrated in the running programme of the SSWP siting crew, it is necessary to discuss the working method of this crew. This working method will be summarized below. For a comprehensive description the reader is referred to the booklet Shallow Wells 1).

The survey for selecting shallow well sites consists of two phases. In the first one a hydrogeological assistant selects the areas near villages for further investigations. The following procedure is applied:

- gathering of information about the existing water supply (pools, springs, etc.) of the villages with the aid of ward heads and village-leaders;

1) DHV (1978): Shallow Wells, DHV Consulting Engineers, Amersfoort, the Netherlands

- establishing the number of the necessary well sites for the villages (one well for about 300 inhabitants);
- selections of promising areas on the ground of the hydrogeological prospects, based on the geology and topography.

The selected areas have to fulfil the following conditions:

- the distance to the village must be less than 1 to 2 km ;
- they must be accesssible for the construction group;
- they must not be situated within 100 m of a source of contamination;
- they must be safeguarded against flooding.

These selected areas are successively numbered per 1:50.000 mapsheet
(e.g. 165/4-7).

In the second phase a detailed investigation is carried out by the siting crew by making several hand-drilled boreholes. The following items have to be mentioned:

- the hand-drillings are carried out with light-weight equipment of small diameter; the maximum drilling depth is 10 metres;
- yield tests; if a water-bearing layer of reasonable thickness is found and if the EC of the water is less than $200 \mathrm{mS} / \mathrm{m}$ a yield test is performed by means of a simple handpump during one hour; drawdown and yield are measured every 10 minutes;
- water quality checks, mainly concerning EC values are carried out during drilling and yield test.
All hand-drillings are successively numbered per area (e.g. 165/4-7-3).


## DA 2.4.2 Equipment

The geo-electrical soundings have been carried out with simple portable equipment. A Bison, type 2350A, earth resistivity meter has been used in combination with a special cable, normally used for "profiling" measurements. With this equipment soundings could be executed rapidly, with a maximum length ( $\mathrm{L} / 2$ ) of 20 meter.
Resistivity well-logging measurements have been performed in the boreholes with a Geohm resistivity meter in combination with a self-made probe with an electrode distance of 0.2 m . These measurements have been carried out with depth intervals of 0.5 m .

## DA 2.4.3 Fieldwork and interpretation

The actual fieldwork took two weeks, in which fifty two geo-electrical soundings have been carried out in thirty three promising areas. The soundings were numbered by adding successive letters to the site number, e.g. 165/4-7-b. Well logging measurements got the same number as the hand-drilled boreholes, e.g. 165/4-7-3.

The geo-electrical soundings, plotted on double logarithmic paper have been interpreted with the curve fitting method using three layer master graphs (see also annex DA 1). In table D 3.2 .2 the relation between the formation resistivity values and the hydrogeological strata is presented.

The curves and their interpretations are given in figure DA 2.1. In the same figures the lithological profile of the corresponding borehole and the data of the resistivity well logging are given. The interpretation took about one week.

The SSWP siting crew carried out fifty three hand-drillings in the investigated area, which has taken six weeks. Thirty seven of these have been executed at the same location as a geo-electrical sounding.

A review of the hand-drillings, geo-electrical soundings and well-loggings per area and per village is presented in table DA 2.2. The results of the different interpretation methods as discussed in DA 2.5.1 and DA 2.5.2 are also given in this table.

DA 2.5 RESULTS

## DA 2.5.1 Quantitative interpretation

The geo-electrical soundings have been interpreted quantitatively into a hydrogeological model. These interpreted models have been compared with the lithological profile gathered from the borehole descriptions. A distinction has been made between a good, bad and "not clear" correspondence of the interpreted model with the hand-drilled lithological profile in view of depth, thickness and resistivities of the layers. These results are given in table DA 2.2 . Ten soundings turned out to have an interpreted model that corresponds well with the drilled profile (e.g. 166/1-2-a). On the contrary the interpreted models of thirteen soundings had a bad correspondence with the borehole data (e.g. 166/1-3-a), while fourteen interpreted geo-electrical soundings have been indicated as "not clear" as compared with the lithological profile (e.g. 166/1-4-a). In some soundings only the top of a layer could be found while the bottom could not be indicated due to the limited length of the soundings. In many cases thin (water-bearing) layers could not be distinguished in the sounding curve due to the theoretical restrictions of the geo-electrical method as mentioned in sub-par. 3.2.3. These thin layers, however, are often of great importance for the exploitation of shallow ground water.
In general it can be said that the correspondence between the interpreted models an boreholes is not very good and the accuracy is low. Therefore it is concluded that a quantitative interpretation of the soundings is not suitable in shallow hydrogeological investigations, which do not exceed a depth of 10 to 15 m .

## DA 2.5.2 Qualitative interpretation

The sounding graphs have also been interpreted in a qualitative way. On the ground of the measured (apparent) resistivities and the general shape of the sounding curves, predictions have been made concerning the possibilities for waterbearing layers with fresh ground water.

Fig. DA 2.1 Legend to geo-electrical soundings and corresponding lithologic bore hole profiles and well logs






































So predictions, divided in good, bad and "not clear" prospects, have been made whether a suitable shallow well could be constructed or not. Later on, these predictions have been compared with the results of the siting crew, see table DA 2.2 . If the prediction was in agreement with the drilling result it is indicated as positive, a disagreement is indicated as negative. The predictions of the qualitative interpretation, the results of the siting crew and their mutual correspondence are summarized in the table below.

Table DA 2.1 - Summary of the results

| Qualitative predictions | Results hand-drillings |  | Correspondence |  |  |
| :--- | :---: | :---: | :---: | :---: | :---: |
|  | Approved | Disapproved | Good | Bad |  |
| Good | 20 | 17 | 3 | 17 | 3 |
| Bad | 8 | 1 | 7 | 7 | 1 |
| "Not clear" | 9 | 5 | 4 | - | - |
| Total | 37 | 23 | 14 | 24 | 4 |

In 9 cases no predictions could be drawn from the soundings. From the remaining 28 cases only 4 predictions are in disagreement with the results of the siting crew. When only boreholes should have been drilled at the locations which were indicated as good or "not clear", just one shallow well site should have been missed. This demonstrates clearly that the geo-electrical method can be used in the surveys for suitable shallow well sites.
A qualitative interpretation takes only little time and can be done immediately after the execution and plotting of the sounding, although it requires some experience of the geophysical operator.

In spite of the above, the geo-electrical method appears not to be efficient in the investigations for shallow well sites. The time which will be saved by the siting crew as a result of the geo-electrical survey, is less than the time necessary for the geo-electrical crew. Moreover the geo-electrical investigations require a skilled geophysical assistant, labourers, equipment and transport. The siting crew cannot be replaced by a geo-electrical crew. First of all one quarter of the soundings did not give a decisive answer about the possibilities for shallow wells; in the second place a yield test is necessary and for construction reasons a sample description is required.
For general shallow hydrogeological investigations these qualitative interpretations are of little importance and therefore not useful.

Table DA 2.2. - Review of the soundings, hand-drillings and the results

|  | Site number | Village | Sounding number | Borehole number | Well <br> logging | Approved for construction | Correspondence between quantitative interpretation and hand-drillings | Correspondence between qualitative predic= thons and hand-drillings |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| 165/1-2 |  | Maqubike | a | 1 | yes | yes | $?$ | - |
|  | 8 | Dumila | a | 2 | yes | yes | - | + |
| $\begin{aligned} & 12 \\ & 14 \end{aligned}$ |  | Mabana | a | 3 | yes | yes | + | + |
|  |  | Mabana | a |  |  | N |  |  |
|  |  | Mabana | $b$ | 1 | no | no | - | + |
| 15 |  | Msufini | a | 1 | yes | yes | - | + |
|  |  | Maufini | b | 2 | yes | yes | + | $?$ |
| 16 |  | Msufini | a | - | - | N |  |  |
|  |  | Msufini | b | 1 | yes | yes | $?$ | + |
| 17 |  | Hembeti | a | 1 | yes | yes | - | + |
|  |  | Hembeti | b |  | - | N |  |  |
|  |  | Hembeti | c | - | - | N |  |  |
|  |  | Hembeti | d | - | - | N |  |  |
| 166/1-1 |  | Mkindo | . | 1 | yes | yes | ? |  |
|  | 2 | Mkindo | a | 1 | yes | yes | + | ? |
|  | 3 | Mkindo | a | 1 | yes | yes | - | + |
| 4 |  | Mkindo | a | 1 | yes | no | ? | - |
|  |  | Mkindo | b | 2 | no | no | ? | - |
| 5 |  | Kigugu | a | - | - | N |  |  |
|  |  | Kigugu | b | 1 | no | no | + | ? |
| 6 |  | Kigugu | a | 1 | no | no | ? | - |
|  |  | Kigugu | $b$ | 1 | no | no | - | + |
|  |  | Kigugu | c | 3 | 00 | no | + | - |
|  | 7 | Kigugu | ${ }^{3}$ | 1 | no | no | - | - |
| 8 |  | Kigugu | a | - | - | N |  |  |
|  |  | Kigugu | b | 1 | no | no | - | $?$ |
| 9 |  | Mbogo | a | 1 | no | no | - | $?$ |
|  |  | Mbogo | b | - | - | N |  |  |
|  | 10 | Mbogo | 4 | 1 | no | no | , | - |
| 11 |  | Mbogo | * | 1 | no | no | ? | - |
|  |  | Mbogo | $b$ | - | - | N |  | - |
| 12 |  | Mbogo | a | 1 | - | N | * | + |
|  |  | Mbogo | b | - | - | N |  |  |
|  | 13 | Mbogo | a | 1 | yes | yes | - | + |
| 14 |  | Mbogo | a | 1 | no | yes | $?$ | + |
|  |  | Mbogo | b | - | - | N |  |  |
|  | 15 | Komtonga | a | 1 | yes | yes | $?$ | ? |
|  | 16 | Komtonga | a | 1 | no | no | ? | ? |
| 17 |  | Komtonga | \# | - | - | N |  |  |
|  |  | Komtonga | b | 1 | yes | yes | - | + |
| 166/3-1 |  | Hembeti | a | 1 | yes | yes | $?$ | + |
|  | 2 | Hembeti | a | 1 | no | no | + | - |
|  | 3 | Hembeti | a | 1 | yes | yes | $?$ | + |
|  | 4 | Hembeti | a | 1 | yes | yes | ? | + |
|  | 5 | Dihombo | a | 1 | yes | yes | + | ? |
|  | 6 | Dihombo | a | 1 | yes | yes | + | + |
|  | 7 | Dihombo | a | 1 | yes | yes | + | + |
| 9 |  | Mkindo | a | 1 | no | no | 7 | + |
|  |  | Mkindo | a | - | - | N |  |  |
|  |  | Mkindo | b | - | - | N |  |  |
|  |  | Mkindo | c | 1 | no | yes | . - | + |
|  | N = | not drill |  |  |  |  |  |  |
|  | + $=$ | correspon | dence is 8 | ood |  |  |  |  |
|  | - = | correspon | dence is |  |  |  |  |  |
|  | ? $\times$ | correspon | dence is " | not clear" |  |  |  |  |

## ANNEX DA 3

SEISMIC INVESTIGATIONS TO SHALLOW DEPTH

## DA 3

SEISMIC INVESTIGATIONS TO SHALLOW DEPTH (special study)
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## DA 3.1 INTRODUCTION

### 3.1.1 General

Within the scope of the hydrogeological investigations the seismic refraction method has been applied. Its main purpose was to study the usefulness of seismic refraction investigations with portable equipment in hydrogeological surveys to shallow depth ( $<10 \mathrm{~m}$ ). The objectives were mainly to determine the depth to the water table and/or to the bedrock. Moreover, it was examined whether this method was suitable in the siting for shallow well sites.

### 3.1.2 Principles of the seismic refraction method

For a detailed description of the seismic refraction method, i.e. the theory, the fieldwork, the interpretation method and the applications, the reader is referred to the literature (see also references sub-par. 2.4., No. 11, 17 and 50). To understand the seismic method, the following should be kept in mind:

- Refraction seismics is based on contrasts in propagation velocity of elastic waves through the earth, which depends upon its elastic properties.
- Seismic (= elastic) waves are bent (refracted) when they cross the boundary between two layers with different wave-velocities. The measure of bending is described by Snell's Law, as in optics.
- The function of a seismograph is to measure the time interval between the initiation of a seismic wave at the source and its arrival at a geophone. The seismograph does not indicate which path the wave followed.
- By measuring this time interval for different spacings between source and geophone a time-distance graph can be constructed.
- Interpreting these graphs, values of two parameters are deduced: the depth to the interface between the different layers and the wave velocities in these layers.
The deduced values for the wave velocities have to be translated into lithologic terms. The expected relation between the wave velocity and lithology, based on literature, is presented in table D 2.2-2.

The seismic refraction method has some limitations. Wrong depth values are obtained e.g. in the cases that:

- a layer is relatively thin (about less than $\frac{1}{4}$ of the depth to its top);
- the wave velocity does not increase with depth for each layer. In these cases one or more layers cannot be distinguished with this method (hidden layer problem), and as a consequence wrong depth values will be interpreted.


## DA 3.2 FIELDWORK

The measurements have been carried out with a portable Bison 1507 B Signal Enhancement Seismograph, which registered the time interval between the initiation of a seismic wave at the source and its arrival at the geophone. As impact source a 5 kg sledge hammer was used. The location of a geophone was fixed during the execution of a spread; the "shotpoint" was shifted successively.

Single spreads and profiles were made along a straight line with equally spaced "shooting" intervals of 2.5 m . The maximum length of a spread amounts to 50 m . If possible, all spreads have been shot "up" and "down". The execution of each spread took about one hour. The measurements have been executed in the period of June to August 1978.

The time distance graphs and the interpreted velocities and depths are given in fig. DA 3.1 up to 12 (at the end of this annex).

Because of the experimental character of the survey, the fieldwork was mainly carried out at locations where either the structure of the underground was well known from hand-drilled boreholes, or the depth to the water table was well known from nearby shallow wells. Most of these boreholes were executed by a MWCP-siting crew to locate suitable shallow well sites (see also DA 2).
All profiles have been numbered succesively per map sheet with the prefix S (e.g. 165/4-S1). Seperate spreads of a profile were succesively numbered per profile with the addition of a letter (e.g. 165/4-S1-a). Two profiles were shot in terrace deposits of the Ngerengere River near Kihonda in order to assess the possibility of measuring the depth to the water table contínuously along a profile.
In these terrace deposits single spreads have been made near a wellknown borehole.
Most of the measurements have been made near the village of Manza in valleys, filled up with alluvial material, in the foothills of the Uluguru Mountains.
Profiles as well as single spreads have been carried out next to boreholes with clear succesion of clayey and sandy layers and the basement most probably at shallow depth.
Near Mirama single spreads have been made near an existing shallow well to establish the possibility of determining the water table depth. In the Berega catchment area an attempt has been made to assess the depth to the ground-water table and to the bedrock in dry riverbeds. No records could be made at these locations because the impact energy of the sledge hammer appeared to be too small in the loose sands of these riverbeds.

## DA 3.3 INTERPRETATION

The readings belonging to one spread, measured with the seismograph, have been plotted on linear graph paper and a time-distance graph has thus been constructed. From this graph velocity values could be derived. After that the depths to interfaces could be calculated with the help of time-intercept formulas.
Later on the records have been interpreted once again according to the "Plus-Minus" method ${ }^{1}$ ). In this method the depth to the interfaces are derived from the "plus" values, i.e. the sum of the travel times from two symmetrical shot points. From the "minus" values, i.e. the difference between the same pairs of travel times the velocities of the refractor along the profile are calculated.

## DA 3.4 RESULTS

The results of the seismic investigations will be discussed successively per location. A brief description will be given of the results with the different interpretation methods. Where the spreads were situated near hand-drilled boreholes or shallow wells the correlation between the results and the lithologic data or with the depth to the water table is dealt with.

DA 3.4.1 Kihonda: 183/3-S1 - fig. DA 3-1
This profile has been interpreted with the "plus-minus" method of which the "minus" graph indicates a rather constant velocity of the second layer ( $1400 \mathrm{~m} / \mathrm{sec}$ ), while the velocity of the upper layer varies widely ( $360-500 \mathrm{~m} / \mathrm{sec}$ ).
Using different velocities for the upper layer the calculated depths of the interfase vary between 1.5 and 5 m , which is rather unlikely for the ground-water table.
Possible causes of this might be:

- differences in the velocity $v_{1}$ of the first layer;
- the occurrence of an intercalated layer will an intermediate wave velocity (hidden layer).
The first possibility seems the most probable however. Especially in cases in which the velocity varies not only verticaly but also laterally (e.g. due to differences in compaction), considerable differences in the interpreted depths can be found.

DA 3.4.2 Kihonda: 183/3-S2-fig. DA 3-2
From the "minus" graph an average velocity of the second layer has been derived of $1670 \mathrm{~m} / \mathrm{sec}$, but considerable differences occur per spread e.g. S2-a: $\mathrm{v}_{2}=1400 \mathrm{~m} / \mathrm{sec}$ and $\mathrm{S} 2-\mathrm{b}: \mathrm{v}_{2}=1900 \mathrm{~m} / \mathrm{sec}$.

1) Hagedoorn, J.G. (1959). The Plus-Minus method of interpreting seismic refraction sections. Geophysical Prospecting, vol. 7, $n^{\circ} 2$.

The "plus" graph shows depths to the first. interface varying between 3 and $5,5 \mathrm{~m}$, which is probably due to a sloping layer. This does not. seem unlikely in these fluvial deposits of the Ngerengere River.

DA 3.4.3 Kihonda: $183 / 3-$ S3 and S4 - fig. DA $3-3$ and 4
These measurements have to be mistrusted because the travel times between the source and geophone are not equal in the forward and reverse direction.
Nevertheless, $S 2$ demonstrates a three layer structure with velocities of $390,600-1000$ and $1500-2400 \mathrm{~m} / \mathrm{sec}$ respectively. This measurement clearly demonstrates that the (dry) upper layer has a higher velocity at greater depth. This velocity, however, is still too low to be correlated with a saturated zone. The calculated depths do not correspond with the handdrilled lithological profile, nor with the depth of the watertable. Confined groundwater may be the cause of this difference.

DA 3.4.4 Manza: 201/1-S1 - fig. DA 3-5
The interpreted depth $d_{1}$ calculated with the time-intercept formula, varies between 3.4 and 2.2 m , which demonstrates that the interpreted depth is not very accurate. The apparent velocity $v_{1}$ of the first layer is rather constant ( $370 \mathrm{~m} / \mathrm{sec}$ ) and indicates a dry upper layer. The interpreted value of $v_{2}$ of the second layer varies between 1175 and 1860, which is a too wide range for a proper prediction about the concerning lithology.
The interpretation of the "minus" graph indicates a velocity of the second layer between 1300 and $1700 \mathrm{~m} / \mathrm{sec}$. A slight "undulation" in the vertical part of the graph suggests in most cases a lateral heterogeneity. The velocity differences might be the result of lithological transitions, e.g. from clay to sand.

DA 3.4.5 Manza: 201/1-S2 - fig. DA 3-6
These time-distance graphs seem to indicate that the vertical velocity distribution consists of many velocity values. In fact these graphs demonstrate an example of "diffraction" which is the result of a considerable lateral heterogeneity (see "burried step model" in Zohdy et al., 1974, references sub-par. 2.4. n ${ }^{\circ} 50$ ).
The "minus" graph shows a s-shaped part which is a clear indication of a lateral inhomogeneity.
Because such a phenomenon is difficult to interpret it will not be discussed further.

DA 3.4.6 Manza: 201/1-S3 - fig. DA 3-7
This single spread has been carried out next to a hand-drilled borehole. The lithological profile shows a relatively thick sand layer between more clayey deposits.

The time-distance graph demonstrates a three layer structure. Two interfaces can be calculated at $3,1 \mathrm{~m}$ and at $8-11 \mathrm{~m}$ respectively of which the first one does not correspond with the drilled profile. The "minus" graph indicates a velocity $v_{3}$ of the third layer of $2500 \mathrm{~m} / \mathrm{sec}$ which may be correlated with the weathered zone. The "plus" graph demonstrates clearly that the interface is slightly sloping.

## DA 3.4.7 Manza: 201/1-S4 and S5 - fig. DA 3-8 and 9

These spreads have also been executed near a borehore. The time-distance graph of S 4 demonstrates a three layer structure. The velocity of the middle layer can not be derived from the "minus" graph but the time distance graph indicates an apparent velocity of about $1400 \mathrm{~m} / \mathrm{sec}$. The "minus" graph indicates a velocity of $2500 \mathrm{~m} / \mathrm{sec}$ for the third layer, probably weathered basement. This is in good correlation with the borehole data, moreover its calculated depth corresponds well with the second interface of 5 m .
The first calculated interface on the other hand does not correspond at all with the hand-drilled profile.
The graph belonging to spread 55 is strongly distorted and no further interpretation has been executed.

DA 3.4.8 Mirama: $165 / 4-\mathrm{S} 1, \mathrm{~S} 2$ and S 3 - fig. DA $3-10$, 11 and 12
These spreads have been made near an existing shallow well in order to assess the possibility of determining the depth to the water table. The time-distance graphs clearly demonstrate a wide variety in velocities in the upper layer. The calculated depths and velocities do not correspond at all with a water table of $1,80 \mathrm{~m}$ below groundlevel. The measured velocities indicate a saturated zone at greater depths ( 10 m ). It is therefore concluded that the watertable in the shallow well is confined. As a consequence a (thin) clay layer has to be supposed on top of the aquifer, which however cannot be recognized in the time-distance graph due to its low velocity value (hidden layer).

## DA 3.5 CONCLUSIONS

An interpretation of the derived velocity values gives reasonable results. Layers with an average velocity of $1500 \mathrm{~m} / \mathrm{sec}$ can in most cases be correlated with saturated sediments, although moist clay layers may have the same values. Layers with velocities of about $2500 \mathrm{~m} / \mathrm{sec}$ are probably semi-consolidated deposits or weathered basement; the second possibility seems the most likely one. This could be confirmed in one case (DA 3.4.7).

On the contrary, the comparison of the calculated depth to the water levels and those found in hand-drilled boreholes or wells was in general very disappointing. The calculated depths to interfaces are not accurate, due to lateral as well as vertical differences in velocity values which occur frequently in the top layers.

The determination of the depth to the bedrock (weathered or unweathered) seems possible with this method although this could be confirmed only once.

The application of the seismic refraction method in hydrogeological investigations for shallow ground water ( $<10 \mathrm{~m}$ ) seems not to be very feasible in the project area, because of its low accuracy in determining the depths to the several interfaces, in particular the depth to the water tables, i.e. water-bearing layers. For the same reason this method is also not suitable in the surveys for shallow well sites.
Apart from this, the execution of seismic refraction measurements and especially their interpretation according to the "plus-minus" method is rather time consuming in comparison with the execution of handdrillings.





Figure DA 3-5 Manza :


Figure DA 3-6 Manza :
201/1-S2
(



Figure DA 3-7 Manza :
201/1-S3



Figure DA 3-8 Manza :
201/1-S4


Figure DA 3-9 Manza : 201/1-S5


Figure DA 3-10 Mirama: 165/4-S1 Figure DA 3-11 Mirama : 165/4-S2 Figure DA 3-12 Mirama : 165/4-S3


ANNEX DA 4
ground water in the ngerengere area SPECIAL STUDY

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DA 4 GROUND WATER IN THE NGERENGERE AREA
DA 4.1 Identification of problems
In the First Report of the MDWSP the Ngerengere area was identified as an area where a large number of villages experience major difficulties in obtaining sufficient drinking water.
The area does not have sufficient surface water to supply all villages with water throughout the year and ground water sources - tapped mainly by means of handdug holes and some shallow wells - give low yields during the dry season.
The chemical quality of part of the ground water in the area was reported to be very poor.

During the hydrogeological survey of the Ngerengere area, carried out between August and November 1978, the water supply problems of this area were quantified and qualified in terms of ground water availability and salinity (table DA 4.6.1). The results of the hydrogeological survey were too general however, to enable the assessment of definite recommendations on the development of potable ground water sources.

Hence the aim of this special study was:

- to study in detail the availability of ground water in the Ngerengere area and near its problem villages in particular.

DA 4.2 Approach to the Study
Evaluation of borehole data, drilled in the Ngerengere area (see table DA $4.4-3$ ) showed that the ground water at greater depth is highly mineralized and unsuitable for domestic purposes.
Hence the study was focussed on the occurence of shallow ground water. Because the geological and consequently also the hydrogeological situation throughout the area varies considerably (see D 4.2), each problem village had to be studied separately.

The field investigations can be split up into three parts:

- study of existing shallow wells and hand-dug holes during the last part of the dry season, when ground water levels decline to a minimum and salinities are increasing accordingly;
- execution of geo-electrical soundings;
- drilling of exploration holes with hand-drilling equipment.

Further activities included the examination and evaluation of water sample analyses results (see table D 10) and the study of aerial photographs.

DA 4.3 Description of the area
DA 4.3.1 General (fig. DA 4.3-1)
The Ngerengere area is the catchment area of the Ngerengere River which takes its rise in the NW Uluguru Mountains and flows around the northern Ulugurus in an eastern direction.
From its sources till the point where it leaves the project area, near Seregete, the Ngerengere River is about 100 km long.
The area is approximately $2210 \mathrm{~km}^{2}$ and contains 42 villages with a total population of more than 50,000 . Morogoro town, with 74,000 inhabitants, which is also situated in the Ngerengere catchment area, is not included in this study.

DA 4.3.2 Topography and geology (fig. DA 4.3-2 and map D 2)
The Ngerengere River takes its rise at elevations between 1200 and 1600 m in the Ulugurus and descends steeply into the 14 km wide valley SW of Morogoro, situated at elevations between 520 and 500 m above sealevel.
From here, down to the junction with the Ruvu River, the slope of the main course is $0,25 \%$, with two knick points caused by lithological differences of the rocks underlying the valley.


Fig. DA 4.3-2 - Longitudinal profile of Ngerengere valley

West of the Uluguru Mountains the Ngerengere River follows $\mathrm{SW}-\mathrm{NE}$ oriented faults that separate the Uluguru granulites from banded migmatites. North of the mountains, the river follows the SSW-NNE oriented Kingolwira Fault, which separates the migmatites in the West from a biotite gneiss complex in the East. In this part of the catchment area, the Precambrian rocks are covered with red soil of varying thickness. Between Mzumbe and Mavulu the Ngerengere River has deposited alluvial material in a $500-1000 \mathrm{~m}$ wide and $10-25$ deep valley.
East of the Kingolwira Fault, the catchment is underlain by biotite gneiss, intruded by a migmatised gabbro body just east of Ngerengere Darajani.

At Ngerengere the river crosses the Kiwege Fault which separates the biotite gneiss in the West from Jurassic sandstones and limestones in the East.

DA 4.3.3 Meteorology and river flow
In the catchment area, rainfall data have been collected from 31 raingauge stations.
The mean annual rainfall in $95 \%$ of the catchment area is about 900 mm ; in the Ulugurus the annual rainfall exceeds 1500 mm .

The annual potential evaporation (Penman) calculated from evaporation data collected at Morogoro, is 1800 mm . Actual annual evaporation may be in of order of 0.4 times the annual potential evaporation, thus 720 mm .

The mean annual discharge at Utari Bridge, near the mouth of the Ngerengere River is $144,000.000 \mathrm{~m}^{3}$. The catchment area here is $2840 \mathrm{~km}^{2}$ hence the annual discharge equals an annual rainfall of 51 mm .

Water from the Ngerengere River is being used by some sisal estates and factories.
The present potential consumptive use is estimated at $0.5 \mathrm{~m}^{3} / \mathrm{sec}$.
A very rough annual water balance over the catchment area may thus be calculated.

Rainfall $=$ discharge + evaporation $\pm$ change in ground water storage + underground outflow of ground water runoff
$900 \mathrm{~mm}=56.5 \mathrm{~mm}+720 \mathrm{~mm}+0 \mathrm{~mm}$ (average annual change in groundwater storage is zero) + underground outflow.

The approximate amount of ground water that probably leaves the catchment area underground thus equals an annual rainfall of 123.5 mm . As explained in 4.2., the only rocks with a high (secondary) permeability in the Ngerengere area, are the Jurassic limestones. The area underlain by the limestones is a karst area and considerable amounts of ground water may flow through subsurface solution channels, and leave the catchment area of the Ngerengere. The figure of 123.5 mm is, however, certainly much too high due to inaccuracies in the mean figures used in the tentative water balance. The occurrence of water losses in the karst area is highly probable but they cannot be quantified.

DA 4.4 Medium-depth and deep ground water
Since 1931, 47 registered boreholes have been drilled in the Ngerengere area.
Some borehole data are listed in table DA 4.4-3; additional information is presented in DD 1,2 and 3 (data-part of this report). At present at least 43 out of these 47 boreholes are abandoned because of high salinity of the ground water, or a very low yield at the time of drilling or after some time of operation.
Nineteen boreholes were drilled in Jurassic sandstones and limestones in the eastern part of the catchment area. The electrical conductivity of the ground water varied from $42 \mathrm{mS} / \mathrm{m}$ to $1050 \mathrm{mS} / \mathrm{m}$ and was $335 \mathrm{mS} / \mathrm{m}$ on an average.

Table DA 4.4-1 - Summary of Ngerengere boreholes drilled in Jurassic

| $N^{\circ}$ of boreholes | EC $\mathrm{mS} / \mathrm{m}$ | Classification |
| :---: | :---: | :---: |
| 1 | $<r 75$ | good |
| 0 | $75-125$ | fair |
| 4 | $125-200$ | poor |
| 9 | $>200$ | beyond acceptable limits |

Chemical data of 5 boreholes are lacking.
Tested yields varied from 0.3 to $2.41 / s$ and were $1 / \mathrm{s}$ on an average.
The remaining 28 boreholes were drilled in Precambrian gneiss. The electrical conductivity varies from $17 \mathrm{mS} / \mathrm{m}$ to $950 \mathrm{mS} / \mathrm{m}$ with an average of $420 \mathrm{mS} / \mathrm{m}$. Borehole $9 / 49,21 / 49$ and $90 / 75$, drilled between Tungi and Kihonda, had extremely high salinities (EC $1300-1770 \mathrm{mS} / \mathrm{m}$ ). Although they were drilled into gneiss, the saline water probably originated from the overlying alluvial deposits.

Table DA 4.4-2 - Summary of Ngerengere boreholes drilled in Precambrian gneiss

| $N^{\circ}$ of boreholes | EC mS/m | Classification |
| :---: | :---: | :---: |
| 3 | $<75$ | good |
| 0 | $75-125$ | fair |
| 3 | $125-200$ | poor |
| 15 | $>200$ | beyond acceptable limits |

Although the EC of 6 boreholes was lower than $200 \mathrm{mS} / \mathrm{m}$, the yield of only one of those boreholes was sufficient for exploitation.

The conclusion of this review is, that it is very risky to try to drill boreholes in the Ngerengere area because yields would probably be very low and the quality of the water would be very poor.

Table DA 4.4-3 - Summary of boreholes, drilled in the Ngerengere area

| Number Location | EC | Tested <br> yield <br> l/s | Total <br> depth <br> (m) | Aquifer(s) | Remarks |
| :--- | ---: | :--- | ---: | :--- | :--- |
|  | $\mathrm{mS} / \mathrm{m}$ |  |  |  |  |
| 1/31 Ngerengere Railway st. 130 | 0.4 | 73.14 | Jura | abandoned |  |
| 2/34 Fatemi S.E. | 130 | 2.4 | 146.31 | Jura | abandoned |
| 4/34 Kizuki S.E. | 42 | 1.3 | 145.70 | Jura | abandoned |
| 6/34 Kiwege S.E. | 250 | 2.5 | 119.48 | Gneiss | abandoned |
| $7 / 34$ Kiwege S.E. | 190 | 5.0 | 50.90 | Gneiss | abandoned |
| 8/34 Mgugo S.E. | 950 | 9.3 | 118.26 | Gneiss | abandoned |
| 14/34 Pangawe S.E. | 460 | 3.8 | 68.57 | Gneiss | abandoned |
| $1 / 35$ Morogoro town | 950 | 0.4 | 48.77 | Gneiss | abandoned |
| 3/35 Kingolwira Farm |  | 0.01 | 150.67 | Gneiss | abandoned |
| 6/35 Ngerengere S.E. | 240 | 4.6 | 63.09 | Gneiss | abandoned |
| 2/36 Fatemi S.E. | 400 | 0.4 | 145.70 | Jura | abandoned |
| $7 / 38$ Tungi S.E. | 170 | 0.4 | 86.26 | Gneiss | abandoned |
| 5/39 Kizuka S.E. | 500 |  | 80.15 | Jura | abandoned |
| 8/47 Kingolwira Prison | 490 | 0.2 | 142.65 | Gneiss | abandoned |


| Number Location | EC $\mathrm{mS} / \mathrm{m}$ | Tested <br> yield <br> 1/s | Total depth (m) | Aquifer(s) | Remarks |
| :---: | :---: | :---: | :---: | :---: | :---: |
| 17/47 Kingolwira Prison | Saline | 0.3 | 76.20 | Gneiss | abandoned |
| 9/49 Tungi S.E. | 15.15 | 0.1 | 121.92 | All/ Gneiss | abandoned |
| 21/49 Tungi S.E. | 17.70 |  | 33.53 | All/ Gneiss | abandoned |
| 3/53 Ngerengere | 300 |  | 117.09 | Jura | abandoned |
| 23/53 Kidugallo | 180 | 1.3 | 121.92 | Jura | abandoned |
| 52/55 Kidugallo Prison |  | dry | 39.62 | Jura | abandoned |
| 6/56 Kidugallo | 400 | 2.2 | 111.66 | Jura | abandoned |
| 21/56 Kihonda S.E. | 17 | 0.03 | 74.07 | Gneiss | abandoned |
| $18^{\text {a } / 59 ~ P a n g a w e ~}$ |  | 0.3 | 85.04 | Gneiss | abandoned |
| $18^{\text {b } / 59 ~ P a n g a w e ~}$ |  | 0.6 | 49.07 | Gneiss | abandoned |
| 23/59 Pangawe | 94 | 8.8 | 56.08 | Gneiss | abandoned |
| 7/62 Sanga sanga | 10.50 | 10.50 | 121.92 | Jura | abandoned |
| 8/62 Mhulazi school | 227 | 227 | 107.90 | Jura | abandoned |
| 9/62 Gomero | 500 | 0.4 | 111.56 | Jura | abandoned |
| 5/63 Mgugu | 650 | 2.2 | 37.19 | Gneiss | abandoned |
| 31/64 Kingolwira |  | 2.3 | 76.20 | Gneiss | abandoned |
| 41/67 Mikese | 300 | 0.6 | 60.69 | Gneiss | abandoned |
| 49/67 Pangawe | Saline |  | 22.25 | Gneiss | abandoned |
| 17/68 Ngerengere |  | Noyield | 67.36 | Jura | abandoned |
| 46/71 Kidugalo | 150 | 1.2 | 201.78 | Jura | in |
|  |  |  |  |  | production |
| 38/72 Magadu | 650 | 1.3 | 135.33 | Gneiss | abandoned |
| 79/72 Morogoro | 800 | 0.1 | 16.80 | Gneiss | abandoned |
| 265/74 Kizuka | Saline |  | 122 | Jura | abandoned |
| 5/75 Kizuka | 680 |  | 136 | Gneiss | abandoned |
| 23/75 Sanga sanga | 360 |  | 52 | Jura | no data, restricted area |


| Number Location | EC $\mathrm{mS} / \mathrm{m}$ | Tested <br> yield <br> $1 / \mathrm{s}$ | Total depth (m) | Aquifer (s) | Remarks |
| :---: | :---: | :---: | :---: | :---: | :---: |
| 70/75 Sanga sanga | Good |  | 85.34 | Jura | no data, restricted area |
| 79/75 Sanga sanga | Good |  | 108 | Jura | no data, restricted area |
| 80/75 Sanga sanga | Good |  | 136 | Jura | no data, <br> restricted <br> area |
| 90/75 Kihonda | 1300 |  | 30.48 | All/ Gneiss | abandoned |
| 77/77 Sanga sanga ujama |  | 1.3 | 36.58 | Gneiss | abandoned |
| 149/77 Fulwe | 23 | 0.5 | 25.60 | Alluvium | abandoned |
| 104/78 Mikese |  | No yield |  | Gneiss | abandoned |
| 117/78 Mikese | 162 |  |  | Gneiss | not yet in production |

DA 4.5 Selection of villages to be studied in detail
DA 4.5.1 General
In order to select priority villages, existing water supply systems, proposed water supply systems as well as systematically collected hydrogeological data have been taken into consideration and are referred to briefly.
Additional information on these subjects can be found in B 3.1, E 3.4.4 and D 3.1 respectively.

DA 4.5.2 Villages to be excluded from the detailed study
The eleven villages, situated in the upper part of the Ngerengere catchment, south-west of Morogoro Town, have been excluded from this detailed study because of the reasons listed below:

- Mlali and Kipera are connected to a piped supply, pumped from a riverbed infiltration well.
- Mzumbe, Changarawe and Kauzeni are connected to a small piped gravity system.
* Sufficient shallow well sites have been found by the MWCP survey department, in Manza and Konga.
- With very little effort and low costs, the other four villages, Tangeni, Vikenge, Sanga Sanga and Mindu can be supplied with gravity piped water (E 3.4.4).
- Construction of the Mindu Dam will start in 1980 , therefore, construction of shallow wells in the Ngerengere Valley within the boundaries of the future reservoir would only be a very temporary solution.

A second group of villages, excluded from this special study are those, directly east of Morogoro Town, in the foothills of the Northern Ulugurus, because Misongeni, Legezamwendo, Kitungwa, Pangawe and Kisinga are all connected to a piped gravity system.
Bigwa also has a small gravity supply.
Both Ngerengere and Kidugalo have pumped, piped supplies from a riverbed infiltration well and a borehole respectively. The riverbed well at Ngerengere, however, dries up in the course of the dry season.

DA 4.5.3 First priority villages
First priority is given to ground water investigations in the villages: Diguzi Mkambarani
Fulwe Mkonowarama
Kihonda Muhugamkola
Kinonko Ngerengere
Kiwege Ngerengere Darajani
Lubungu Seregete A
Maseyu Seregete B
Mikese Tungi
Mlilingwa
These villages experience major difficulties in obtaining sufficient drinking water throughout the year and the water of many holes and wells contains high concentrations of salt.

DA 4.5.4 Second priority villages
The villages Visarake, Magela, Lubuma, Matuli, Lulongwe and Kwamba are situated along the Ngerengere River and some of its main tributaries. River water is used during the wet season and the main part of the dry season.
Hand-dug holes in the riverbeds give sufficient water during the short periods when there is no surface flow.
Construction of shallow wells along or in the riverbeds will probably not offer serious problems.

## Table DA 4.6-1 - Hydrogeological field data Ngerengere area



## DA 4.6 Hydrogeological field survey

Between July and October 1978, hydrogeological fieldwork was carried out in the Ngerengere area.
The existing ground water supplies of more than 20 villages in this area were examined, the hydrogeology of the area as a whole was studied in the field and locations were selected where the occurrence of ground water could be studied by means of geo-electrical investigations and handdrilling.
Data on existing ground water supplies collected during the field work are presented in table DA 4.6-1.
In 16 out of the 20 villages, ground water is being exploited by means of hand-dug holes or shallow wells.
However most of the holes and wells dry up during the dry season and some of them have saline ground water and are not being used therefore.

## DA 4.7 Geo-electrical investigations

In the Ngerengere area, 37 geo-electrical soundings have been executed; their locations are given on fig. DA 4.3-1.
The numbers of these soundings are:
$183 / 1-17,-18,-19,-20,-21,-22,-23,-24,-25$
183/2-1
$183 / 3-1,-2,-3,-4,-5,-6,-7,-8,-9,-10,-11,-12,-13,-14,-15$, $-16,-17,-18,-19$.

The sounding curves and their interpretation are presented in DD 7.
With the exception of soundings $183 / 3-2,-9$ and -16 , all soundings show weathered basement at shallow depth.
From these soundings, information on the occurrence of ground water in the weathered zone, and its salinity, cannot be detected because the degree of weathering of the basement varies considerably form place to place and the resistivities vary accordingly.
Only the depth to the fresh bedrock can be found from these soundings. Soundings $183 / 3-2,-9$ and -16 were located in the alluvium of the Ngerengere River.
Up to a depth of 25 m , very low resistivity values ( $2-30 \mathrm{hmm}$ ) occur in these soundings, indicating the presence of highly saline ground water.

## DA 4.8 Execution of the hand-drillings

Between mid-January and the end of April, 1978, 101 exploration holes were drilled at various locations, across the valleys of the Ngerengere River and some of its tributaries.

Due to lack of time, not all priority villages, listed in the previous paragraph, could be studied. The collected data have been sufficient however, to gain insight into the hydrogeology of the Ngerengere area and the prospects for shallow ground water for the villages that were not studied may be deduced from this knowledge. Profiles were drilled near the following villages:

| Village 1- | $\mathrm{N}^{\circ}$ of hand-drillings | Profile |
| :--- | :---: | :---: |
| Kihonda | 14 | A |
| Tungi | 20 | B |
| Mkambarani | 11 | $\mathrm{C}^{1}, \mathrm{C}^{2}$ |
| Mkonowarama | 4 | $\mathrm{D}^{1}, \mathrm{D}^{2}$ |
| Mikese | 5 | $\mathrm{E}^{1}, \mathrm{E}^{2}$ |
| Lubungo | 5 | $\mathrm{~F}^{\mathbf{1}}, \mathrm{F}^{2}$ |
| Ngerengere Darajani | 6 | G |
| Kinonko | 4 | H |
| Muhungamkola | 6 | I |
| Vidaraka | 8 | J |
| Masimbu | 12 | (K) |
| Mkundi | 6 | (L) |

The borehole data are presented in DD 9.
The drilled profiles are described in the next paragraph.

## DA 4.9 Description of hand-drilled profiles

DA 4.9.1 Profile A at Kihonda (fig. DA 4.9-1)
At Kihonda, profile $A_{1}-A_{12}$ was drilled, north-south accross the Ngerengere Valley.
Boreholes A-1 to A-4 presented the same sequence.
From top to bottom the following layering was found:

- a 1.0 to 1.5 m thick top layer of orange/red loamy soil;
- a 1.0 m thick layer consisting of a mixture of angular quartz stones and calcrete and ferrite nodules in a matrix of orange/red loam;
- weathered biotite gneiss; this basement material was proved only in borehole A-3, at a depth of 2.3 m .

The orange/red loam probably is residual soil material.
There is no single theory explaining the formation of calcrete, different mechanisms can be active under different circumstances. In drilling A-4, ground water was struck in the loamy gravel layer. The electrical conductivity of this water was $925 \mathrm{mS} / \mathrm{m}$.

Figure DA 4.9-1 Hand-drilled profile across Ngerengere Valley, at Kihonda
(s)

Legend: see Fig. D 3.3-18


Alluvial deposits of the Ngerengere River were encountered in boreholes A-5 to A-11, over a distance of 750 m . Boreholes A-5 to A-7 more or less showed the same sequence.

- A 1.5 to 3.0 m thick top layer of very tough darkbrown-black swamp clay; this clay layer also occurred in drillings A-8 to A-11.
- A water-bearing sand layer, varying in thickness from 0.6 m to 2.6 m and in texture from fine to very coarse.

The EC of the ground water in this sand channel varies from 1475 to $295 \mathrm{mS} / \mathrm{m}$.

Below the sand, there is a gravelly brown/grey mottled clay, down to about 10 m below surface, followed by a bluish loam.

Drillings A-8 and A-9 show an identical layering.
Below the 1.4 - 2.0 m thick blackish clay, there is a 7 m thick sequence of brown/grey mottled loam and fine sandy clay, alternated with 3 thin water-bearing sand layers. The EC of the ground water in these layers is $400 \mathrm{mS} / \mathrm{m}$.

At a depth of 9 m , the same bluish loam as present in A-14 was encountered.

In borings A-10 and A-11, ground water was struck in an orange/red loamy gravel, identical to the residual material found in A-1 to A-5.
Some infiltration of water from the Ngerengere River into this layer occurs since the EC is between 127 and $190 \mathrm{mS} / \mathrm{m}$. A 5 m deep shallow well, constructed just south of the river, is not being used by the villagers of Kihonda, because of the salty taste of the water (EC $190 \mathrm{mS} / \mathrm{m}$ ).

DA 4.9.2 Profile B, at Tungi (fig. DA 4.9-2)
At Tungi, a profile of 20 boreholes ( $B-1$ to $B-20$ ) was drilled across the Ngerengere Valley.
The profile resembles the Kihonda profile.
Borings B-1 to B-8 on the northern slope of the valley and B-17 and B-18 on the southern flank present identical sequences: a 1.0 to 2.5 m thick orange/red loam on top of an orange coloured loamy gravel. The gravel consists of angular quartz pieces and lumps of calcrete. Evidence of basement material was not found in these borings but from other drillings it is known that the gravel zone with quartz and calcrete invariably rests on basement rock.
Ground water in the gravel layer was only encountered in borings B-6 and $\mathrm{B}=17$; the EC of the water was 525 to $800 \mathrm{mS} / \mathrm{m}$, which is about the same as the ground water in this gravel layer near Kihonda.

Alluvial deposits were met with in borings $B-9$ to $B-20$, over a width of 1000 m across the valley.
The upper layer in these borings is a 1.2 to 3.0 m thick blackish swamp clay, which occasionally is sandy in its lower parts.

Covered by this clay, a sand channel was found in drillings B-9 to B-15. The salinity of the ground water in the sand body is very high; up to $3800 \mathrm{mS} / \mathrm{m}$.

Much fine to coarse-textured, water-bearing sand was present in borings B-19 and B-20; both near the present riverbed.
The EC of the water here was 500 to $800 \mathrm{mS} / \mathrm{m}$, almost the lowest of the profile.

The lower parts of the drillings are made up of brown/grey mottled, sometimes loamy or fine sandy clay with some quartz pieces. In Maji borehole 41/49, drilled down to 34 m , weathered gneiss was struck at a depth of 10.5 m .

DA 4.9.3 Profiles $C^{1}$ and $C^{2}$ at Mkambarani (fig. DA 4.9-3)
As can be seen in profile $C^{1}$, the alluvium deposited by the Lukonde River in the neighbourhood of Mkambarani, is more than 100 m wide and $4-5 \mathrm{~m}$ thick.
The uppermost $2-4 \mathrm{~m}$ of the alluvium, is a sticky, brown coloured clay which confines the sand layers that occur between this clay and the weathered gneiss.
In the profile three different sand lenses were encountered. These
lenses probably are cross-sections through old channels and therefore the sand bodies may have a considerable length.
In two of the three channels, saline ground water is present under confined conditions. The EC varies from 350 to $800 \mathrm{mS} / \mathrm{m}$.
In the sand, encountered in borings $C^{1-4}$ and $C^{1-7}$, ground water with an
EC of $135 \mathrm{mS} / \mathrm{m}$ was found. It is possible that some recharge of this aquifer takes place, the EC of the ground water in boring $c^{1}-6$ (upper parts of weathered basement) was $120 \mathrm{mS} / \mathrm{m}$.
There may be a hydraulic connection between the outcropping wheathered basement south of boring $C^{1}-6$, and the aquifer of $C^{1-4}-C^{1}-7$. In boring $C^{1}-7$ a pumping test will have to be carried out to see whether salt water is attracted by pumping this aquifer.
If this doesn't occur, than there are some possibilities for exploitation of this aquifer by means of shallow wells with hand pump.

Profile $C^{2}$ is drilled across the valley of a small tributary of the Lukonde River.
The alluvial deposits are more than 50 m wide.
Confined ground water occurs in a $0.5-2.5 \mathrm{~m}$ thick sand layer, at depths between 3.5 and 6.5 m below ground level.

Figure DA 4.9-3 Hand-drilled profiles at, Mkambarani


Below the sand, weathered basement was found.
The salinity in most part of the aquifer is relatively low: 56-
$80 \mathrm{mS} / \mathrm{m}$; in boring $\mathrm{C}^{2}-4$ an EC of $200 \mathrm{mS} / \mathrm{m}$ was found.
Location $C^{2}-2$ seems an attractive site for constructing a shallow well with hand pump. The sand is coarse and free of clay particles; the bottom of the sand layer is gravelly here. Again a pumping test will have to be carried out first, to assess the yield and the quality of the ground water after some time of pumping.

DA 4.9.4 Profiles $D^{1}$ and $D^{2}$ at Mkonowarama (fig. DA 4.9-4)
Near the village of Mkonowarama, 5 boreholes were drilled at three different locations along the Ngerengere River.

In boring $\mathrm{D}^{\mathbf{1}}-1$, water-bearing sand was found below 3.5 m of sticky brown clay. The sand layer is only 0.5 m thick and rests on weathered basement. The ground water was unconfined; the water level is the same as the water level of Ngerengere River.
The ground water is extremely saline; the EC was $1900 \mathrm{mS} / \mathrm{m}$.
In boreholes $D^{2}-1,-2$ and -3 , predominantly sandy material occurs. The water levels correspond with the water level of the Ngerengere River. Close to the river, the $E C$ is $650 \mathrm{mS} / \mathrm{m}$; further from the river, EC's up to $1900 \mathrm{mS} / \mathrm{m}$ occur.

DA 4.9.5 Profiles $\mathrm{E}^{1}$ and $\mathrm{E}^{2}$ at Mikese (fig. DA 4.9-5)
Although the Mikese River is only a very small stream, it has deposited a considerable volume of alluvium.
North of the river, only clays were encountered on top of the weathered basement.
Just south of the river, in boring $E^{1-4}$, two water bearing sand layers were found. The EC of the ground water was $120 \mathrm{mS} / \mathrm{m}$. The total thickness of the aquifer, 3.0 m and its coarse texture make it interesting for the construction of a shallow well at the location of $E^{1-4}$.

Profile $\mathrm{E}^{2}$ shows the presence of a 4 - 5 m thick slightly clayey sand body, below 2 meters of clay. Confined ground water with an EC of 88 to $120 \mathrm{mS} / \mathrm{m}$, occurs in this aquifer. Because of the fact that neither in profile $E^{1}$ nor in profile $E^{2}$, saline ground water was present, the prospects for shallow ground water seem fair in this valley.

Figure DA 4.9-4 Hand-drillings in Ngerengere River Valley, at Mkonowarama




DA 4.9.6 Profiles $F^{1}$ and $F^{2}$ at Lubungo (fig. DA 4.9-6)
As can be seen on profile $F^{1}$, alluvium occurs only west of the present course of the Lubungo River.
East of the river, orange/red loamy soils were found on top of the weathered basement.
A small isolated pocket of ground water with an EC of $45 \mathrm{mS} / \mathrm{m}$ was found in boring $F^{1-2}$; it offers no possibilities however.
The alluvium of the Lubungo River is about 80 m wide.
Below 3-4m of sandy clay, there is a sand body that wedges out against the basement in a western direction.
 water occurs (EC $577 \mathrm{mS} / \mathrm{m}$ ).
Because of this, there seem to be no possibilities for exploitation of shallow ground water here.

In profile $\mathrm{F}^{2}$, predominantly clays were encountered.
Ground water occurs in a thin sand layer in boring $\mathrm{F}^{2}-3$ and in the gravel layer on top of the weathered basement in boring $\mathrm{F}^{2}-2$.
Because of the presence of highly saline ground water, there seem to be no possibilities here for shallow wells.

DA 4.9.7 Profile G, at Ngerengere Darajani (fig. DA 4.9-7)
In Ngerengere Darajani, the Ngerengere River has incised itself relatively deeply into the basement.
Alluvial deposits of limited extend have locally been sedimented along the river. In borings G-4 to G-7, such alluvium was encountered. Borings G-1 to G-3 and G-8 were drilled in orange/red loamy soils.
Ground water of a low salinity (EC 50 up to $150 \mathrm{mS} / \mathrm{m}$ ) occurs in an aquifer with considerably varying thickness and texture from borehole to borehole.
No saline ground water was found at this location and therefore there seem to be possibilities for exploitation of shallow ground water. Wells should preferably be constructed as close to the river as possible so that pumping of a well will induce a ground water flow from the riverbed into the sandy aquifer.

DA 4.9.8 Profile $H$, at Kinonko (fig. DA 4.9-8)
At Kinonko, ground water was found to occur in a 60 m wide 2 m thick, slightly clayey coarse sandlayer, confined by 2-3 mof sandy clay. The low salinity of the ground water, and the coarse texture of the sand offer some prospects for the exploitation of this aquifer. Pumping tests will have to be carried out to find out if the EC does not increase after some time of pumping, and also to determine the possible yield of a shallow well in this aquifer.


Figure DA 4.9-7 Hand-drilled profile across Ngerengere Valley, at Ngerengere Darajani

(s)

Figure DA 4.9-8 Hand-drilled profile across Matule River, at Kinonko


DA 4.9.9 Profile $I$, at Muhungamkola (fig. DA 4.9-9)
The alluvium deposited by the Muhungamkola River at the village Muhungamkola, is about 60 m wide and $6-8 \mathrm{~m}$ thick.
Below sandy clay which is $1-2 \mathrm{~m}$ thick, water bearing sand is present. The ground water is confined and its EC varies in the different borings from 75 up to $220 \mathrm{mS} / \mathrm{m}$.
A second aquifer in boring $I-6$, contains saline ground water.
Location I-4 seems to offer some prospects for shallow well construction.
However, pumping tests are necessary to be able to decide if this tentative conclusion is justified.

DA 4.9.10 Profile J, at Visaraka (fig. DA 4.9-10)
This profile, drilled across the valley of the Ngerengere River, downstream from the village Ngerengere, differs from the profiles described thus far. In boring $\mathrm{J}-1$ to $\mathrm{J}-4$, an up to 9 m thick sand layer occurs below 1-2 m of sandy clay and clayey sand. The sand is coarse to very coarse and only occasionally slightly clayey.
The sand is water-bearing: the ground water seems to be unconfined, and the EC is low: $65-75 \mathrm{mS} / \mathrm{m}$.
In borings J-5 to J-8 predominantly clay was found alternated with thin sand or gravel layers. Here saline ground water occurs (EC up to $350 \mathrm{~ms} / \mathrm{m}$ ).
The best prospects for shallow ground water are therefore present, near the river.
The aquifer is certainly thick and permeable enough, to install shallow wells with motor pumps but it is possible that after some time of pumping saline ground water will flow towards the well.

DA 4.9.11 Profile $K$, at Masimbu (fig. DA 4.9-11)
Masimbu is not a village, but a camp, Schools for people from of a Liberation Movement are being built here. A request was made to the MDWSP to assess the ground water potential of the area around Masimbu. To this end, 12 holes were drilled along the Ngerengere River, at intervals of approximately 50 m .

Fig. DA 4.9-11 shows that the individual boreholes are lithologically very different from each other. The thickness of aquifers and the salinity of the ground water varies also considerably. Generally, most boreholes have very clayey sequences, alternating with sands containing saline ground water (EC's up to $1200 \mathrm{mS} / \mathrm{m}$ ). Locations that probably are suitable for construction of shallow wells with handpump are: K-7 and K-11.

Figure DA 4.9-9 Hand-drilled profile across Muhumgamhola Valley, at Muhumgamhola
I-1


Figure DA 4.9-10 Hand-drilled profile across Ngerengere Valley, at Visaraka


Figure DA 4.9-11 Hand-drilled holes at Masimba, along the Ngerengere River


Pumped supplies from shallow ground water in this area is not possible because of the clayey character of the aquifers and the presence of saline ground water.

DA 4.10 Conclusions
A. Examination of data from borehores drilled in the Ngerengere area proves, that the prospects for medium-depth and deep ground water in most parts of the Ngerengere area are poor to extremely poor. The yield of boreholes generally is low ( $<3 \mathrm{l} / \mathrm{s}$ ) to very low ( $<11 / s$ ) and almost everywhere the ground water is saline. EC's up to $3800 \mathrm{mS} / \mathrm{m}$ have been measured. Fair prospects for exploitation of medium-depth to deep ground water are only present in the narrow (2 - 3 km wide) $\mathrm{N}-\mathrm{S}$ orientated zone, underlain by Jurasic limestones (see D 4.2.7.).
B. In the Ngerengere valley upstream from Sanga Sanga settlement (fig. DA 4.3-1) 10 m up to 30 m thick and up to 900 m wide alluvial deposits occur. Since downstream from Sanga Sanga the Ngerengere River cuts relatively deeply into the basement and no alluvium is present there, the valley upstream from Sanga Sanga must have undergone a relative subsidence. Clays and sandy clays make up more than $90 \%$ of the upper 12 m of the sediments. Sands are usually water-bearing, the ground water being confined, but the salinity of the ground water is extremely high so that there are no prospects for shallow ground water in this part of the area. Only close to the river, shallow wells may locally be possible.
C. Between Sanga Sanga settlement and Ngerengere, the Ngerengere River is incised into the Precambrian gneiss and almost no alluvium has been deposited here.
As can be seen on Map D 4, this part of the catchment area is bounded by NE-SW running major faults, and is relatively uplifted. Alluvium occurs only locally in the valley of the Lukonde, Matule and Mikese Rivers, and some other streams. The alluvium is up to 100 m wide and 5 - 10 m thick. Approximately $80 \%$ of the deposits are clays; sands and clayey sands make up $20 \%$ of the alluvium. The ground water is confined and the EC's are mostly below $200 \mathrm{mS} / \mathrm{m}$ although saline ground water was found locally also.

Exploitation of shallow ground water by means of shallow wells with handpump might be possible in the villages: Mkambarani, Mikese, Muhungamkola and Kinonko.
No data are available yet on the water level fluctuations in these aquifers; some aquifers probably will run dry during the dry season.

Also the presence of saline ground water very close to the areas where fresh ground water was found, forms a potential danger. Pumping of ground water might induce the saline ground water to flow towards the wells.
D. Downstream of Ngerengere, the river has deposited alluvial sediments which are more than 12 m thick and 500 m up to more than 1000 m wide. This area is not underlain by Precambrian gneisses, but mainly by Jurassic sandstones and this probably is the reason for the abundance of sands in the alluvial deposits. At Visaraka, medium to very coarse sands make up more than $75 \%$ of the alluvial sediments. The sands are water-bearing and the EC is Low ( $<75 \mathrm{mS} / \mathrm{m}$ ). Saline ground water was found to occur in clays and loams on the slopes of the valley. Shallow wells therefore will have to be located close to the river. Because of the high permeability of the sands, and the considerable thickness and extend of the aquifer, it is possible to withdraw a few $1 / s$ from a well. Larger yields however will increase the chances of salinisation of the aquifer. Careful planning and testing of aquifers is recommended before pumped supplies from shallow wells are taken into consideration.
E. By extrapolation of the collected and evaluated hydrogeological data, predictions can also be made concerning the prospects for shallow ground water in other problem villages.
The summary below describes the prospects for shallow ground water of all problem villages in the Ngerengere area.

Diguzi
No field data were collected on the occurrence of shallow ground water near Diguzi.
The valley, approximately 1 km south of this village was visited and alluvium was found to occur hexe over a width of 100 m across the valley. Possibilities for shallow ground water are present here.
Since the EC's of the ground water in Kiwege, upstream from Diguzi, and Visaraka, downstream from Diguzi are low, low salinity of the ground water near Diguzi is expected also.

Fulwe
The topographical location of Fulwe is very unfavourable; the village is situated almost on the watershed between the Ngerengere and Lukonde Rivers.
Ground water is present in small valleys in and near the village. However during the dry season the wells and holes dry out and the ground water is saline.

Approximately 1.5 km south of the village there is a small alluvial valley, where ground water of a low salinity ( $25 \mathrm{mS} / \mathrm{m}$ ) occurs. The two shallow wells in this valley are hardly being used however; because of the distance. If shallow wells with handpump are to be constructed in this valley, a road will have to be constructed first.
A borehole was drilled in Fulwe (149/77), but its yield is very low, $0.4 \mathrm{l} / \mathrm{s}$ with a drawdown of 9.3 m .

## Kinonko

At Kinonko shallow wells can be constructed in the valley of the Matule River. Average depth of the wells will be 7.0 m ; the aquifer thickness is approximately 1.6 m . The EC of the ground water generally is below $200 \mathrm{mS} / \mathrm{m}$ but saline ground water occurs also and salting up of the aquifers during the dry season may be expected.

Kiwege and Mlilengwa
These villages are situated on water sheds.
Small quantities of shallow ground water occur in a few very small alluvial valleys around the villages. The yield of hand-dug holes in these valleys is extremely low; the EC of the ground water is only $9 \mathrm{mS} / \mathrm{m}$, which points to very "young" water. There may be small possibilities for concrete ring wells.
Yields will still be very low and during the dry season there may be a shortage of ground water.

Lubungo
Near Lubungo there are two small valleys where alluvium was deposited. The ground water in these alluvial sediments is saline. There are no possibilities for shallow wells.

Maseyu
This village is situated on a water shed.
Ground water of low salinity occurs in very small valleys, filled with sandy alluvium. The ground water occurs in pockets and many of these pockets dry up during the dry season.

A detail study of the aquifers around this village will have to be carried out, to find out whether there are possibilities for shallow wells here.

## Mikese

Shallow wells can be constructed in the valley of the Mikese River. The EC of the ground water is $100 \mathrm{mS} / \mathrm{m}$ on an average.
The seasonal water level fluctuations in this valley are considerable. Water levels vary from 1.5 to 7.0 m below ground level.

## Mkambarani

Small amounts of ground water can be withdrawn from aquifers in valleys south and east of the village.
Because of the presence of saline ground water in the same valleys, the wells may become more saline during the dry season.

Mkonowarama, Kihonda, Tungi
All ground water occurring in the alluvium of the Ngerengere Valley is extremely saline.
There are no possibilities for shallow wells therefore.

## Muhungamkola

Shallow wells with handpump can be constructed in a small valley, east of the village.
There is danger that the aquifer becomes more saline during exploitation, however.

## Ngerengere

Shallow wells can be constructed along the Ngerengere River in sandy alluvium. The EC is low; $75 \mathrm{mS} / \mathrm{m}$.

Ngerengere Darajani
Narrow isolated patches of alluvium occur along the Ngerengere River.
There are some possibilities of exploitation of shallow ground water from the alluvium, especially close to the river, where hydraulical connections exist between the riverbed and the alluvium along the river. Because of the presence of saline ground water, there is danger of salinisation of the aquifers.

## Seregete A and B

There are no possibilities for extraction of shallow ground water within a distance of 1.5 km from these villages.
The nearest source of ground water is the riverbed and the alluvium of the Ngerengere River, at a distance of respectively 5 and 3 km from these villages.

Kwamba Lubumu, Lulogwe, Magela, Matuli, Visakara
Shallow wells can be constructed in the extensive alluvium of the Ngerengere River and its main tributaries.
There is an abundance of sandy material in the alluvium.
Yields of shallow wells will be high and the EC of the ground water is between 75 and 125 mS 7 m .
Because of the presence of saline ground water on the higher slopes of the valley, in weathered basement material, wells will have to be constructed quite close to the river.

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SPECIAL STUDY
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DA 5.
DA 5.1.
INTRODUCTION
In the First MDWSP Report, the northern part of the Kilosa district was identified as an area experiencing most serious domestic water supply problems. Hence priority was given to the setting up of a study to assess the availability of ground water in this area, corresponding with the Berega Catchment and shown in fig. DA 5.1-1.
In a separate study (Volume III), the surface hydrology section of the MDWSP established the quantity of surface water available for domestic use.

Activities started in June 1978, with a hydrogeological reconnaissance survey of the area.
During this survey, data collected within the scope of the identification of conditions and problems of rural water supply in the northern part of the Kilosa district, dealt with in the First MDWSP Report, were quantified and qualified.
Hydrogeological fieldwork was carried out in 40 villages in the area. Hydrogeological field data were systematically collected between June and July 1978, and are tabulated in DD 4.

An evaluation of data on deep boreholes drilled in this part of the project area is presented in D 3.1.1.
On the basis of the hydrogeological survey, the evaluation of boreholes data and the geological maps of the area, a general description of the hydrogeology of this area could be outlined in D 4.2.4. and D 4.2.13.

Three different waterbearing formations could be distinguished.
Shallow ground water occurs in:

- alluvial sands and gravels that fill up the riverbeds of the Berega river and its main tributaries
- . sands and gravels of higher level fluvial terraces

Ground water at greater depth is occasionally present in:

- weathered zones, faults and shear zones of the gneissic basement rocks

Table DA 5.4.6-1 shows that ground water in the basement rocks is invariably highly mineralized and unsuitable for domestic use. In addition, yields of boreholes are very low.
For this reason it was decided to concentrate this detailed study on the occurrence of shallow ground water.
Since all problem villages are situated within the catchment area of the Berega River, it has been found logical from a hydrogeological point of view to confine the investigations to this basin.

Thus, the aim of this detailed study was:

- to assess the quantity of potable ground water at shallow depth ( $0-12 \mathrm{~m}$ ) available for domestic water supply throughout the year
- to find hydrogeological reasons for the occurrence of saline ground water in the basin
- to locate suitable shallow well locations

The Berega Catchment Area; Location of villages


DA 5.2. APPROACH TO THE STUDY
DA 5.2.1. General
A reconnaissance survey, carried out at the beginning of the dry season by the MDWSP-hydrogeological section showed that the ground water quality and potential varied widely over the area.

As a result of this survey two sub-areas could be distinguished within the Berega River Basin (fig. DA 5.2.1-1).

Sub-area $I$, the western part of the catchment, has been characterized as one with poor prospects for the construction of shallow or river wells, due to a considerable salinity of the ground water and insufficient ground water potential of the riverbed aquifers.
Ground water quality and potential of sub-area II, the eastern half of the catchment, seemed much better although considerable differences in salinity were measured at relatively short distances.

Because prospects for the construction of shallow or river wells seemed fair in sub-area II, a detailed study was set up to establish the ground water potential of the riverbed aquifer during the dry season and to find an explanation of locally high salt contents.

The main purpose was:

- to determine the thickness of the riverbed aquifer
- to find out if a vertical and/or horizontal layering of ground water with different salinity content exists
- to estimate the hydraulic permeability
- to measure the water table and salinity changes regularly during the dry season
- to locate good river well sites

At those locations where older terrace deposits occupy part of the valley, the cross-sections have also been extended to cover these. The aim of this was:

- to find out whether a hydraulic connection exists between the riverbed aquifer and other waterbearing zones
- to examine the chemical quality of ground water in these deposits
- to locate good shallow well sites

Fig. DA 5.2.1-1

## Sub-areas based on fieldwork



DA 5.2.2. Fieldwork
To determine the ground water potential of the riverbed aquifer the following fieldwork was executed:

- as close as possible to a problem village, hand borings were drilled along cross-sections in the riverbeds
- samples were taken for examination from every drilled metre and the salinity of the ground water, expressed in EC-values, was measured
- $\quad 1$ inch diameter PVC piezometer tubes were installed in every borehole

Geo-electrical soundings, covering the hand drilled profiles, were carried out in the riverbeds themselves and on the older terrace deposits, if present. The soundings were made in order to get more information about thickness of the weathered zone and to establish the depth to the unweathered bedrock.

Boreholes, drilled in the vicinity of geo-electrical soundings were well-logged, in order to gather resistivity data that could be of use during the interpretation of the geo-electrical sounding curves.

All profiles, drilled boreholes, locations of geo-electrical soundings and installed piezometer tubes were levelled.

Water samples were taken to get an impression of the chemical quality of the ground water. Extracting water samples from the observation wells was not possible, therefore water samples were taken from hand-dug holes near the cross-sections.

DA 5.2.3. Equipment
The boreholes were drilled with light equipment, manufactured in the Netherlands (Eykelkamp).
This lightweight set comprises the following items:

- $\quad 2$ handles +10 extension rods of one metre each
- 2 Edelmann combination bits of 7 cm diameter
- $\quad 2$ Edelmann combination bits of 10 cm diamter
- $\quad 2$ riverside bits of 7 cm diameter
- $\quad 2$ riverside bits of 10 cm diameter
- $\quad 1$ screw auger of 4 cm diameter
- $\quad 1$ bailer of 63 mm diameter
- $\quad 21 \mathrm{~m}$ slotted galvanized casing pipes of 3 inch diameter
- 81 m unslotted PVC casing pipes of 3 inch diameter
- $\quad 10$ extra extension rods of one meter each
- 1 complete set for pumping (hand pump)

Geo-electrical soundings were carried out with portable equipment comprising the following items:

- Bison 2300A earth resistivity meter
- 2 reels with 250 m cable each
- $\quad 440 \mathrm{~cm}$ iron pens (current and potential electrodes)
- 8 connection cables

Conductivity measurements in the observation wells were carried out with:

- Bison 2300A earth resistivity meter
- resistivity probe of 2 cm diameter
- reel

Well-loggings were carried out with simple, partly self-made, portable equipment which comprises the following items:

- geohm resistivity meter
- probe with an electrode distance of 0.2 m and 10 m cable

Levelling was performed with:

- Wild NAK 2
- $\quad 4$ meter staff

EC-values during drilling were measured with:

- VU-IVA conductivity meter

DA 5.2.4. Elaboration of data
Fifteen sections across the riverbeds of the Berega, Kibedya, Magera, Kinyolisi and Dogomi Rivers were hand drilled and geo-electrically investigated.

These cross-sections have been numbered successively per map sheet e.g. 165/1-1. A total of 65 holes were hand drilled; the total drilling depth was 299.50 metres. The borings have been numbered per cross-section with the addition of a number e.g. 165/1-1-3. Because several holes turned out to be dry, only 57 were used to install piezometer tubes which have been given the number of the hand drilled boreholes in which they are installed.

A total number of 37 geo-electrical soundings were carried out. They are also numbered per cross-section, but with the addition of a letter. Well-loggings are numbered as the boreholes they have been made in.

DA 5.3. DATA COLLECTION AND ELABORATION
DA 5.3.1. Hand drilling
The samples have been described by the MDWSP hydrogeology section. Grain size, colour and mineral content have been determined. The lithological profiles were reconstructed together with the surveyor's borehole descriptions.

These lithological profiles were worked out in cross-sections, together with the levelled measuring points.

Water levels and the EC-values measured during drilling are also given in these sections.

DA 5.3.2. Geo-electrical soundings and well-loggings
Geo-electrical soundings were plotted on double logarithmic paper and interpreted with the curve-fitting method using 3-layer mastergraphs; the sounding curves and their interpretations are given in DD 11. For more theoretical background information about geo-electrical soundings and interpretation methods reference is made to D 3.2 .4 . Well-loggings together with the corresponding lithological profile are given in DD 12.

The results of the interpreted sounding curves are also given in the previously mentioned cross-sections. The depths of geo-electrically uniform layers as well as their average resistivities are indicated.

DA 5.3.3. Water quality
Taking water samples out of the piezometer tubes failed due to lack of a suitable pump. Therefore water samples were taken from hand-dug holes in the vicinity of the cross-sections. All samples were chemically analyzed by the MDWSP water supply section. The results are summarized in table DA 5.7.2-2.

DA 5.3.4. Ground water flow
Ground water flow within the alluvial aquifers can hardly be calculated for the whole river basin. Many small tributaries had not been incorporated in this study and it proved to be difficult to determine the real extention of the riverbed aquifers.
To get an impression of the ground water potential, however, a rough estimation was made of the minimum ground water flow in the riverbed.

A longitudinal profile of the Berega River. and its main tributaries was reconstructed (Fig. DA 5.3.4-1) to determine the hydraulic gradient, assuming that this is equal to the gradient of the riverbed. It appeared that the gradient is rather constant over the river bed and varies between 0.013 and 0.021 (see Fig. DA 5.3.4-1).

Minimum saturated vertical sections of the aquifers have been estimated from the cross-sections. The actual extention of the aquifer was not established exactly in all cases, as drillings often had to be stopped at limited depth due to hard layers or caving in of the boreholes. Minimum sections have been established for situations at the beginning of a dry season (ground water table at ground level), the end of the dry season of 1978 (mid-November), and an extremely dry year (beginning of the rainy season mid-December).
To this end, the decline of the ground water levels, measured during the dry season of 1978, has been extrapolated from the figures presented in DD 13.
Ground water flow through the different sections has been calculated using permeabilities estimated from the borehole samples.
Minimum calculated ground water flows vary between $4110 \mathrm{~m}^{3} /$ day downstream at Ibondo, to $535 \mathrm{~m}^{3} /$ day upstream at Ndogomi.

The area is $1760 \mathrm{~km}^{2}$, hence, assuming an average annual rainfall of 650 mm (see DA 5.4.5.), the total amount of precipitation that falls yearly in the catchment area is about $1144 \times 10^{6} \mathrm{~m}^{3}$. The average ground water discharge of the Berega riverbed near the junction of the Berega with the Chogowale River (Ibindo) is approximately $4500 \mathrm{~m}^{3} /$ day (Table DA $5.7 .1-1$ ) or $1.643 \times 10^{6} \mathrm{~m}^{3} /$ year, which is only $0.15 \%$ of the annual precipitation. The remaining part of the rainfall evaporates or is discharged as surface water.

## DA 5.3.5. Measurement programme

From July up to December 1978, measurements were executed of water levels in piezometers installed by MDWSP. At the same time, the EC's of the ground water in many hand-dug holes were measured.
The data are presented in DD 13.
During the course of the dry season, the ground water table in the alluvial aquifer did not decline more than 0.5 m on an average.
At most locations where the EC was measured at regular intervals during the dry season, there was hardly any change in salinity.
An increase in salinity was only measured at Mabula, Magera, Ndogomi, Nguyomi and Mbili.

DA 5.4. DESCRIPTION OF THE SURVEY AREA
DA 5.4.1. Topography
The Berega River Basin drains the central part of the eastern Ukaguru Mountains. The basin is about 80 kilometres long down to the junction with the Chogowale River, one of the main tributaries of the Mkindu River (see Fig. DA 5.1-1).
The basin drains moderately to semi-arid terrain between elevations of 1500 and 600 metres above sea level.
The drainage area is about $1760 \mathrm{~km}^{2}$ of which roughly $5 \%$ is the main valley floor, $15 \%$ is rugged mountains and the remaining $80 \%$ is upland terrain.

Slopes are moderate, $4-10 \%$, with some inselbergs reaching elevations of 300 to 800 metres above their surroundings.

The valley is a NW-SE orientated depression bounded by faults. To the NE the Nguru Mountains rise steeply, forming a drainage divide at elevations up to 1500 metres. To the $S W$ the divide follows elevations up to 1400 m of the Kaguru Mountains.

DA 5.4.2. Geology (Map D 2)
The catchment area is situated between two mountain ranges; the Nguru Mountains in the NE and the Kaguru Mountains in the SW.
The Nguru Mountains mainly consist of granulites. The Kaguru Mountains are composed of tightly folded gneiss, amphibolite, meta-gabbro and quartzite, locally intruded by granite.

The Berega River Basin itself is underlain by pre-Cambrian biotite gneiss of sedimentary origin, folded into a NW-SE strikingly broad syncline.
In many places the weathered or unweathered bedrock is covered with some metres of red-brown sandy soil, derived mainly from the biotite gneiss.

Remnants of fluvial and colluvial depositional terraces, locally up to several hundred metres wide and a few metres thick are found in the main valleys.
These sediments are well bedded, consisting of dark-brown sandy loams, loamy sands and coarse sands to fine gravels, most probably deposited during early Quaternary times.
The top of these terraces is situated at elevations of a few metres above the present riverbeds.
The present riverbeds of the Berega River itself and its main tributaries, the Kibedya, Magera and Kinyolisi River, are from 5 to 50 metres wide.
They are cut into the old terraces or bedrock and filled up with some metres of coarse to very coarse sand. Geoelectrical measurements carried out during this study indicate the presence of a fairly thick weathered zone beneath these alluvial deposits.

## DA 5.4.3. Hydrology

The catchment boundaries and the main rivers are given in fig. DA 5:1-1. None of the rivers in the Berega catchment area are perennial; they only flow after heavy rains.
The occasional floods do not occur more than a few times a year and the inundations of the terraces are limited to a few days a year.
During these floods a considerable amount of water is discharged, carrying a heavy sand load.
After the rain stops, rivers flow intermittently for some time and from about halfway through the dry season no more surface discharge takes place. During the dry season drainage takes place only through the aquifer formed by the alluvial sands and gravels that filled up the riverbeds of the Berega River and its main tributaries.

DA 5.4.4. Land use and vegetation
Agriculture is limited to small local farms (mashamba) in the vicinity of the villages. The main food crops are maize and sorghum. Cash crops are very rare and consist mainly of cotton and sunflowers and to a lesser extent of beans. No large farms or estates occur in this area.

Most of the cultivated fields are situated on the older terraces.
Sometimes very recent lower terraces bordering the present riverbeds are also used for cultivation.

The vegetation of the remaining upland terrain mainly consists of open forest, scrub and scattered trees. The slopes of the Nguru Mountains and the inselbergs are bare or support light forests.

Stock herding activities are extensive throughout the area. Lifestock units for the different villages in the Berega area are given in Volume II, Table BD 3-1.

DA 5.4.5. Meteorology
Rainfall data were collected from two stations within the Berega River catchment area. The mean annual rainfall in Berega (Station Berega Mission), with reasonably reliable records, amounts to 781 mm . Rainfall data from a station in Gairo are less reliable and show a mean annual rainfall of about 505 mm (records from 1970). Rainfall data have been summarized in table DA 5.4.5.-1.

Evaporation data have been derived from records of the stations Kongwa and Morogoro. Mean annual potential evaporation measured at these stations amounts to 1967 mm and 1760 mm respectively.

Evaporation in the Berega area will probably be somewhat lower because of its higher elevations.
The evaporation may possibly be less than 100 mm in May, June and July and will only exceed 150 mm in October, November and December.

Table DA 5.4.5.-1 Rainfall data

Monthly Rainfall (mm) for Station:
uERECA HISSIOH
Registration Nuriber: 96.3703

| Year | Jan | Feb | Harch | April | \%ay | June | July | Aug | Sept | Oct | Nov | Dec | ```Jan - Dec Total``` | $\begin{gathered} \text { Nov - Oct } \\ \text { Total } \end{gathered}$ |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| 1950 | 103.1 | 150.1 | 143.5 | 159.0 | 59.7 | 4.3 | 8.4 | 0.0 | 0.0 | 18.0 | 16.8 | 99.1 | 762.0 | * |
| 1951 | 51.6 | 198.4 | 68.8 | 55.4 | 70.1 | 4.8 | 8.6 | 0.0 | 0.0 | 29.2 | 49.8 | 125.5 | 662.2 | 603.1 |
| 1952 | 227.1 | 128.3 | 100.1 | 135.4 | 82.6 | 2.8 | 0.0 | 0.0 | 30.0 | 0.0 | 54.9 | 25.4 | 786.4 | 881.1 |
| 1953 | 144.8 | 12.4 | 136.7 | 101.3 | 92.5 | 12.4 | 14.7 | 12.2 | 0.0 | 0.0 | 2.5 | 104.1 | 633.7 | 607.4 |
| 1954 | 158.2 | 191.8 | 39.9 | 95.5 | 53.1 | 1.0 | 0.0 | 0.0 | 0.0 | 21.3 | 5.1 | 77.7 | 643.6 | 667.4 |
| 1955 | 56.1 | 219.5 | 24.9 | 108.7 | 212.3 | 40.4 | 5.1 | 0.0 | 0.0 | 0.0 | 24.6 | 166.1 | 756.7 | 648.8 |
| 1956 | 215.9 | 124.5 | 83.3 | 175.0 | 44.2 | 17.3 | 0.0 | 0.0 | 5.1 | 0.0 | 54.6 | 100.6 | 820.4 | 855.9 |
| 1957 | 178.1 | 45.7 | 141.0 | 264.9 | 154.4 | 11.7 | 0.0 | 0.0 | 3.0 | 25.4 | 27.4 | 245.6 | 1097.3 | 979.5 |
| 1958 | 58.2 | 174.8 | 189.7 | 116.1 | 52.3 | 35.3 | 0.0 | 0.0 | 0.0 | 0.0 | 14.2 | 99.3 | 739.9 | 899.4 |
| 1959 | 133.1 | 218.9 | 191.0 | 94.2 | 32.8 | 0.0 | 0.0 | 30.5 | 0.0 | 0.0 | 0.0 | 77.0 | 777.5 | 814.0 |
| 1960 | 149.6 | 62.7 | 99.8 | 139.2 | 36.4 | 25.9 | 25.7 | 0.0 | 1.3 | 1.5 | 0.0 | 0.0 | 544.1 | 621.1 |
| 1961 | 15.2 | 200.7 | 52.1 | 147.3 | 59.7 | 0.0 | 77.5 | 0.0 | 0.0 | 104.1 | 150.4 | 184.4 | 991.4 | 656.6 |
| 1962 | 315.0 | 97.0 | 83.3 | 81.3 | 30.5 | 0.0 | 25.4 | 0.0 | 0.0 | 0.0 | 3.0 | 148.6 | 784.9 | 968.1 |
| 1963 | . 100.5 | 136.6 | 99.6 | 71.8 | 15.1 | 12.7 | 0.0 | 0.0 | 0.0 | 0.0 | 231.2 | 88.1 | 755.6 | 587.9 |
| 1964 | 198.6 | 172.7 | 206.3 | 85.8 | 55.8 | 55.7 | 28.0 | 0.0 | 0.0 | 20.3 | 0.0 | 218.7 | 1041.9 | 1142.5 |
| 1965 | 238.8 | 194.5 | 78.7 | 35.6 | 35.6 | 0.0 | 0.0 | 0.0 | 0.0 | 20.4 | 63.4 | 104.1 | 771.2 | 822.4 |
| 1966 | 25.5 | 27.9 | 111.8 | 139.7 | 73.6 | 53.3 | 40.6 | 0.0 | 0.0 | 0.0 | 38.1 | (50.0) | (560.5) | 639.9 |
| 1967 | 27.9 | 94.3 | 30.5 | 250.7 | 207.1 | 2.3 | 33.0 | 12.7 | 10.2 | 0.0 | 43.1 | 230.3 | 942.1 | (756.8) |
| 1968 | 150.7 | 139.5 | 286.0 | 135.1 | 51.6 | 25.4 | 0.0 | 0.0 | 0.0 | 0.0 | 58.4 | 20.3 | 967.0 | 1161.7 |
| 1969 | 57.9 | 186.0 | 128.2 | 69.0 | 33.5 | 10.5 | 2.8 | 6.8 | 8.9 | 9.4 | 36.8 | 41.7 | 591.5 | 591.7 |
| 1970 | 284.5 | 179.8 | 121.1 | 53.6 | 62.3 | 0.8 | 0.0 | 0.0 | 0.0 | 2.8 | 7.4 | 84.2 | 796.5 | 783.4 |
| 1971 | 129.5 | 111.3 | 60.6 | 157.0 | 57.4 | 19.3 | 0.0 | 0.0 | 0.0 | 12.9 | 1.4 | 61.2 | 611.8 | 640.8 |
| 1972 | 63.0 | 162.8 | 197.2 | 92.5 | 127.6 | 2.1 | 24.7 | 0.0 | 45.1 | 40.3 | 29.5 | 137.7 | 922.5 | 817.9 |
| 1973 | 267.6 | 218.7 | 30.3 | 217.0 | 24.4 | 12.8 | 0.0 | 5.1 | 0.0 | 1.6 | 18.5 | 115.9 | 911.9 | 944.7 |
| 1974 | 64.9 | 55.2 | 29.7 | 278.7 | 75.7 | 3.1 | 14.0 | 0.0 | 6.4 | 17.1 | 10.0 | 53.9 | 501.2 | 571.7 |
| 1975 | 160.4 | 38.9 | 177.6 | 160.2 | 62.6 | 9.9 | 2.9 | 1.6 | 9.9 | 1.0 | 1.3 | 88.5 | 714.8 | 688.9 |
| 1976 | 66.9 | 164.6 | 86.3 | 107.3 | 71.7 | 18.9 | 15.4 | 0.2 | 0.5 | 0.0 | 2.5 | 54.5 | 588.8 | 621.6 |
| 1977 | 195.6 | 188.8 | 96.1 | 105.6 | 123.1 | 1.1 | 21.1 | 14.9 | 19.4 | 11.0 | 81.3 | 127.0 | 985.0 | 833.7 |
| 1978 | 114.0 | 112.2 | 216.9 | 136.0 | 67.1 | 15.6 | 3.4 | 1.0 | 0.0 | 0.0 | 112.6 | 309.1 | 1086.9 | 874.5 |
| n(1953-77) | 25 | 25 | 25 | 25 | 25 | 25 | 25 | 25 | 25 | 25 | 25 | 25 | 25 | 25 |
| - | 138.3 | 136.8 | 111.2 | 131.3 | 69.9 | 14.9 | 13.2 | 3.4 | 4.1 | 11.1 | 36.2 | 107.2 | 778.1 | 773.0 |

Honthly Rainfall (ma) for Station:
GaIRo
Registration Number: 96.3626

| Year | Jan | Feb | Harch | April | May | June | July | Aug | Sept | Oct | Nov | Dec | Jan - Dec Total | Nov - Oct Total |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| 1970 | 251.2 | 59.1 | 98.7 | 15.4 | 7.4 | 0.0 | 0.0 | 0.0 | 0.0 | 0.0 | 0.0 | 102.2 | 534.0 | * |
| 1971 | 57.3 | 97.1 | 50.7 | 64.9 | 28.3 | 15.0 | 15.3 | 0.5 | 3.0 | 3.3 | 2.5 | 30.8 | 388.9 | 437.6 |
| 1972 | 77.9 | 277.4 | 121.8 | 39.5 | 95.9 | 0.0 | 0.0 | * | * | * | * | 118.4 | * | * |
| 1973 | 121.2 | 184.9 | 90.0 | 51.5 | 53.3 | 0.7 | 0.0 | 0.0 | 0.0 | 0.0 | 0.0 | 108.8 | 610.4 | * |
| 1974 | 57.3 | 15.6 | 31.1 | 114.9 | 45.1 | 1.0 | 4.0 | 1.1 | 0.7 | 21.0 | 28.7 | 45.7 | 366.2 | 430.6 |
| 1975 | 138.0 | 75.0 | 103.1 | 74.4 | 28.8 | 0.0 | 0.0 | 0.0 | 13.2 | 16.9 | 30.1 | 143.7 | 623.2 | 523.8 |
| 1976 | * | , | , | * | , | , | * | * | * | * | . | * | * | , |
| 1977 | 303.0 | * | 29.1 | 26.2 | 11.1 | 0.0 | * | 0.0 | 13.2 | 13.7 | 86.3 | 87.1 | * | * |
| 1578 |  |  |  |  |  |  |  |  |  |  |  |  |  |  |

DA 5.4.6. Ground water
The area's main aquifer is formed by alluvial sands and gravels of recent age that filled up the riverbeds of the Berega River and its main tributaries. Phreatic conditions seem to exist everywhere. Ground water is also present in the older sedimentary terraces and residual soils that occupy parts of the main valley and its smaller tributaries.

Ground water level measurements, executed in piezometer, installed during this special study, indicated a ground water flow from the older terraces and residual soils towards the riverbed aquifer, during the course of the dry season.

The salinity of the shallow ground water varies widely over the river basin; ground water in the Kibedya riverbed has a salinity frequently exceeding the maximum acceptable value of $200 \mathrm{mS} / \mathrm{m}$, while the water in the Kinyolisi riverbed is of a much better quality (EC values up to $65 \mathrm{mS} / \mathrm{m}$ ).

The gneiss of the pre-Cambrian basement complex is virtually impervious and the occurrence of deep ground water appears to be restricted to the weathered zone, fault lines and shear zones.
Table DA 5.4.6-1 shows that yields of deep boreholes are very low and the ground water is strongly mineralized. Therefore chances to find suitable deep well sites with sufficient yield and acceptable quality must be considered very poor.

Table DA 5.4.6-1 - Boreholes drilled in the Berega area and surroundings

| number | location | EC $\mathrm{mS} / \mathrm{m}$ | Q lit/sec. | total depth | remarks |
| ---: | :--- | :--- | :--- | :--- | :--- |
|  |  |  |  |  |  |
| $2 / 53$ | Gairo | 198 | 0.5 | 73.8 | abandoned |
| $41 / 54$ | Chakwale | 249 | 0.9 | abandoned |  |
| $15 / 60$ | Rubeho | - | dry hole | 75.3 | abandoned |
| $16 / 60$ | Meshugi | saline | 0.2 | abandoned |  |
| $34 / 60$ | Rubeho | 240 | 0.5 | abandoned |  |
| $37 / 60$ | Kisitwe | - | dry hole | 68.8 | abandoned |
| $18 / 65$ | Magubike | 490 | 1.3 | 76.0 |  |
| $33 / 65$ | Gaixo | 255 | 0.7 | 90.2 | abandoned |
| $41 / 76$ | Mtumbatu | - | dry hole | 91.4 | abandoned |
| $282 / 77$ | Mtambatu | 175 | 1.1 | 62.5 |  |

DA 5.4.7. River flow and ground water recharge
During the wet season, when rivers and streams are flowing continuously, the riverbed aquifer is saturated with water of relatively low salinity.

After the main rainy season, rivers flow intermittently for some time and about halfway through the dry season no more surface discharge takes place.

During the part of the dry season that the riverbed aquifer is no longer recharged by infiltration of stream runoff, ground water seepage from alluvial and colluvial material on the valley and mountain slopes gradually becomes the only source of recharge. Later in the dry season recharge may stop entirely.

Recharge of the terrace deposits, colluvium and eluvium, as well as the weathered and fissured parts of the basement rocks takes place during the wet season, when parts of the valley floor are flooded and water can percolate down to greater depths.

DA 5.5. PRESENT WATER SUPPLY SITUATION
DA 5.5.1. Introduction
The water supply situation in the Berega area; as described below, is based on data collected for the First MDWCP Report, supplemented with data collected by the hydrogeological as well as the hydrological sections during their respective field surveys in this area.

As stated before, the northern part of the Kilosa District was identin fied as an area experiencing most serious water supply problems during the dry season. The present water supply situation (1978) during this season may be summarized as follows:

- about 34,000 people ( $66 \%$ ) in the area depend on hand-dug holes in the riverbeds
- seven villages (about 7750 people) do not have any domestic water supply within 2 km walking distance
- eight villages (about 12,000 people) use water of very poor quality
- two villages, at the end of a gravity scheme, only receive water for a few days per week
- the pumped, piped water supply of one village suffers from regular breakdowns

The different types of domestic water supply in the Berega area and their specific problems will be discussed separately.

DA 5.5.2. Gravity piped water supply (5 villages; 12,000 people)
Five villages, Gairo, Ukwamani, Majawanga, Kibedya and Chakwale, with about 12,000 inhabitants are connected to a gravity piped water supply that also provides water to five villages outside the basin (Gairo gravity scheme).
The water is diverted from two small perennial streams in the mountains south-east of Kisitwe by means of gravity flow through pipes.
The water is distributed through public taps in the villages.
Only Gairo, Ukwamani and Majawanga are certain of a supply throughout the year. During the dry season, however, the villages Kibedya and Chakwale (about 4700 inhabitants), at the end of the pipeline, only receive water for a few days per week.

DA 5.5.3. $\frac{\text { Pumped, piped water supply from a river well (l village; }}{2650 \text { people) }}$

Only Berega and its two sub-villages Chirihara and Berega Mission are connected to a pumped water supply. The water is pumped from a concrete river well in the Berega riverbed into 3 storage tanks that have to be refilled every day. The water is distributed to public taps.

Apart from occasional breakdowns of the pump and lack of diesel oil, this system works properly for about 8 months a year.
During the last months of the dry season, when the ground water table is much lower, the storage capacity of the river well is not sufficient, consequently it takes several days of intermitted pumping to fill the storage tanks.

DA 5.5.4. Pumped, piped water supply from a spring (1 village; 1350 people)

The whole year round, the village of Maguha receives water from a small spring ( 0.3 lit/sec).
This spring has been developed by means of a small dam and two storage wells. From the wells the water is pumped to seven public taps and one cattle dip.

This system shows, that even with a very small amount of water, not varying much throughout the year many people can be served.

DA 5.5.5. Pumped, piped water supply from a deep borehole (1 village, 2800 people)

Only one village, Magubike (about 2800 inhabitants), is connected to a pumped water supply from a deep borehole. In the past more boreholes were drilled in this area, but most of them have been abandoned (see table DA 5.4.6-1). The main reasons are a very low yield and strongly mineralized ground water (EC values up to $500 \mathrm{mS} / \mathrm{m}$.

The borehole in Magubike is connected to public taps and in spite of the very high EC value ( $490 \mathrm{mS} / \mathrm{m}$ ) these taps are still used. For drinking water purposes, however, the villagers collect water from the river in the wet season and from hand-dug holes during the dry season.

DA 5.5.6. Non-improved water supply (22 villages; 32,550 people)
DA 5.5.6.1. Water supply within walking distances
The villages Ibindo, Mwandi, Kiegeya, Mabula, Nguyami, Madege, Iogwe, Makuyu, Ndogomi, Idibo, Leshata, Magera, Italagwe and Kinyolisi, with a total of about 18,000 inhabitants, all fall into the same category; during the wet season people use water from the Berega River, its tributaries and temporary pools, while during the dry season the only sources of water are hand-dug holes in the riverbeds.

There are no major supply problems since the hand-dug holes give water the whole year around, provided that one digs deep enough. The main problem, however, is not the quantity but the quality of the water in these hand-dug holes. Especially the villages along the Kibedya River suffer from saline water (Nguyami EC up to $450 \mathrm{mS} / \mathrm{m}$ ).

The people of Kiegeya ( 900 inhabitants) have to dig holes in alluvial and residual loamy soils during the dry season.
Due to the low permeability of this waterbearing material, people have to wait a long time before they can fill their buckets.
Moreover, the ground water is of a very poor quality ( $E C>200 \mathrm{mS} / \mathrm{m}$ ).
DA 5.5.6.2. No water supply within walking distance
The villages Mkalama, Meshugi, Ngiloli, Tabu Hotel, Mtumbatu, Mbili, Machatu, Kitaite and Kilama (about 9700 inhabitants) have the most serious water supply problems.

During the wet season people from these villages may collect water from nearby temporary pools, but for more than six months a year the villagers have to collect water at considerable distances.

People from Mkalama and Meshugi go to the public taps in Majawanga (Gairo gravity scheme) at respectively 2 and 5 kilometres distance.

The villagers from Ngiloli and Tabu hotel go to the public taps in Kibedya at distances of 3 and 4 kilometres respectively. If, during the last months of the dry season, the tap in Kibedya does not work these people have to walk to Gairo at 6 and 8 kilometres respectively.
The people of Ibuti go to public taps in Gairo at 2.5 kilometres distance.

The villagers of Mbili, Mtumbatu and Kilama have to walk more than 7 kilometres to reach their water supply at the end of the dry season, when ground water in small streams close by the villages dries up.

DA 5.6. DESCRIPTION OF CROSS-SECTIONS AND INDIVIDUAL HAND DRILLINGS

DA 5.6.1. General
The results of the field investigations will be discussed separately per village. A brief description is given of the lithological situation of the cross-sections and individual hand drillings, drilled in the vicinity of the villages in question. Fig. DA $5.6-1$ shows the location of the cross-sections. A summary of the ground water salinity and prospects for the construction of river or shallow wells is given per village as well.
Ground water quality and potenṭial are discussed separately per village in DA 5.7.1. and DA 5.7.2.

DA 5.6.2. Cross-sections
Fig. DA 5.6-2 gives the legend to the cross-sections presented in Fig. DA 5.6.2-1 to -18.

## Location of the cross-sections



Fig. DA 5.6-2 : Legend to cross sections

1 number borehole


basement outcrop
(E) orientation
drillings stopped due to:
HL : hard layer
CL : stickey clay
ST : stones
CA : caving
GR : gravel

DA 5.6.2.1. Berega, Berega River (165/1-1)
The existing water supply and its problems in the village Berega have been discussed in DA 5.5.3.
The villagers depend on hand-dug holes in the riverbed of the Berega River during the times that the pumped piped system is out of order or insufficient, therefore this village has also been involved in the study.

The profile is situated along the track between the sub-villages Berega Mission and Chirihara.
The whole profile shows clean coarse gravelly sands. Drillings had to be stopped at limited depth, however, due to gravel layers and boreholes caving in.
Geo-electrical measurements failed at this location, probably due to lateral heterogeneity indicated by bedrock outcrops in the vicinity of the spot.
The NE situated terrace consists entirely of weathered basement (borehole $165 / 1-1-5$ ) without a hydraulic connection with the riverbed aquifer. Coarse, gravelly sands underly some finer deposits of the SW situated terrace. A hydraulic connection exists between this terrace and the riverbed and because its $E C$ value is about $100 \mathrm{mS} / \mathrm{m}$ prospects for the construction of a shallow well seem favourable.
EC values gradually increase in NE direction; this is probably due to inflow of mineralized water from weathered material of the SW situated terrace.

DA 5.6.2.2. Mabula, Kibedya River (165/1-2)
This cross-section is situated about 500 m SW of the village.
The riverbed itself consists of coarse sand with occasional gravels, underlain by sandy clay of fluvial origin. The more clayey deposits of borehole 165/1-2-4 on the other hand are underlain by sandy clay, probably weathered basement material.
The NE situated terrace shows a gradual increase of more clayey deposits, probably also weathered basement material.
A hydraulic connection exists between the riverbed and both terraces,
but because of its higher topographical level and its position with regard to the village, the NE terrace is the most suitable for the construction of shallow wells.
The cross-sectional diagram shows rather high EC values ( $190 \mathrm{mS} / \mathrm{m}$ ) but measurements carried out later indicated much lower values of about $140 \mathrm{mS} / \mathrm{m}$, which is permissible for a temporary solution.

Geo-electrical soundings indicate a layer with high resistivities representing bedrock at an average depth of 25 metres.
The top layer on the other hand shows resistivities of about $14-18 \mathrm{~m}$ which indicates (clayey) sands or weathered bedrock with saline ground water. The latter possibility seems the most probably.
(sw)

FIGURE DA 5.6. 2-2 Cross-section 165/1-2 Kibedya Valley at MABULA


DA 5.6.2.3. Mabula, Berega River (165/1-3)
This profile is situated about 500 metres $S E$ of the village and covers only one terrace. The first three boreholes in the riverbed encountered coarse sand underlain by sandy clay which is probably weathered material from the higher older deposits.

This cross-section clearly demonstrates that high EC values of ground water in riverbeds is often due to inflow of highly mineralized water from older terraces and weathered basement material.
Boreholes $165 / 1-3-5,6$ and 9 demonstrate a clear layering in water salinity, not only laterally but vertically also and a gradual increase in salinity (up to $800 \mathrm{mS} / \mathrm{m}$ ) in NW direction.
Therefore prospects for the construction of shallow wells are very poor on this terrace. The quality of the water in the riverbed is reasonable, although EC values tend to be high next to the terrace. EC measurements show values of about $130-150 \mathrm{mS} / \mathrm{m}$ which makes prospects for the construction of river wells fair.

Geo-electrical soundings indicate the average depth to bedrock to be about 17 metres. The middle layer shows a resistivity decrease from 34 to 8 Sm in NW direction; this clearly corresponds with an increasing EC value in that direction.
Drillings show that this layer probably consist of weathered basement with saline ground water.

DA 5.6.2.4. Magera, Magera River (165/1-4)
This cross-section is situated along the track between Mabula and Magera about 500 m south of the village. The profile shows medium coarse to coarse sands interbedded with some clay layers. The northern terrace is overlain by clayey fine sands while the other terrace consists entirely of weathered basement material.

A hydraulic connection exists between both terraces but, because of the water quality, prospects for the construction of shallow wells are good only on the northern terrace.
Ground water quality of the riverbed is in general good (EC $<125 \mathrm{mS} / \mathrm{m}$ ) although some inflow of mineralized water occurs from the weathered material of the southern riverbank.

Geo-electrical soundings show a middle layer with resistivities varying between 32 and 60 sm , which probably indicates a weathered zone with fresh ground water ( $\mathrm{EC}<150 \mathrm{mS} / \mathrm{m}$ ).
Depth to bedrock (resistivity $<100 \Omega \mathrm{~m}$ ) gradually increases in northern direction up to 16 metres below ground level.


DA 5.6.2.5. Magera, Magera River (165/1-5)
The profile is located about 500 metres west of the village near a hand-dug hole. The riverbed itself shows very coarse sand with gravel, underlain by weathered basement (clay). The eastern terrace is overlain by sandy clay and clayey sand, probably weathered basement material. Because of a medium coarse sand layer of reasonable thickness and good quality ( $\mathrm{EC}<125 \mathrm{mS} / \mathrm{m}$ ) of the ground water, prospects for shallow well construction are satisfactory on this terrace.

Geo-electrical soundings failed at this location because of a considerable lateral heterogeneity indicated by bedrock outcrops in the vicinity of this cross-section.

DA 5.6.2.6. Kinyolisi, Kinyolisi River (165/1-6)
The cross-section is situated about 600 metres north of the village near temporary hand-dug holes in the riverbed.
The riverbed itself, only 8 metres wide, consists of very coarse sands underlain by gravels. Drillings had to be stopped at limited depths probably due to these gravel layers.
The northern terrace consists of coarse sands of fluvial origin overlain by some finer loamy material. Due to the general good quality of the ground water, prospects for the construction of shallow wells are good on this terrace.
This cross-section clearly demonstrates that the Kinyolisi River is one of the tributaries of the Berega River system with a considerable better ground water quality; EC values vary between 40 and $80 \mathrm{mS} / \mathrm{m}$.


FIGURE DA 5.6. 2-6 Cross-section 165/1-6 Kinyolisi Valley at KINYOLISI


DA 5.6.2.7. Ndogomi, Ndogomi River (165/1~7)
The profile is located about 1300 metres NE of the village centre. Hand-dug holes in the riverbed are the only source of water supply for the villagers and cattle during the dry season.
The riverbed itself is filled up with clayey medium coarse sand with rather low permeabilities. The north eastern terrace consists entirely of weathered basement material while the more flat south western terrace is of fluvial origin.
Although a hydraulic connection exists between both terraces, the south western one is most suitable for the construction of shallow wells. Salinity content seems reasonable, but salinity checks carried out later showed higher values (up to $170 \mathrm{mS} / \mathrm{m}$ ), which make shallow or river wells only temporarily permissible.

Geo-electrical soundings in the narrow bed failed. Sounding 165/1-7-a, however, indicates the average depth to bedrock to be about 7 metres which correlates with the drilling results.

DA 5.6.2.8. Madege, Ndogomi River (165/1-8)
This cross-section is situated about 700 metres NE of the village. The riverbed itself consists of medium coarse to coarse sands underlain by sandy clay of fluvial origin. The sloping south-western terrace consists of weathered basement material without a hydraulic connection with the present riverbed.
The other side becomes more clayey in north-eastern direction and is obviously of fluvial origin, so that prospects of the construction of shallow wells are good (location of borehole 165/1-8-3) on this terrace, although it must be realized that salinity contents tend to be high (> $150 \mathrm{mS} / \mathrm{m}$ ).
The analysis of a water sample taken from a hand-dug hole in the riverbed shows that the manganese content exceeds the maximally acceptable limits of the WHO-standards.

According to geo-electrical sounding results, layers with high resistivity (weathered bedrock) are present from a depth of 14 metres approximately. The middle layer shows resistivities of about 20 Sm which indicates a high clay mineral content or weathered basement material with saline ground water.

FIGURE DA 5.6. 2-7 Cross-section 165/1-7 Ndogomi Valley at NDOGOMI

(*)

FIGURE DA 5.6. 2-8 Cross-section 165/1-8 Ndogomi Valley at MADEGE


## DA 5.6.2.9. Nguyami, Kibedya River (165/1-9)

The cross-section is located about 700 metres south of the village centre along the track to Machatu. The riverbed consists of medium to coarse sands underlain by gravelly layers. The northern terrace also consists mainly of coarse sands underlain by gravels.
Preliminary surveys showed that the Kibedya River is one of the tributaries with ground water of poor quality. This profile clearly demonstrates that salinity content of the ground water in the riverbed as well as in the terrace exceeds the maximum acceptable level of $200 \mathrm{mS} / \mathrm{m}$. Therefore this location in unsuitable for the construction of river or shallow wells.

Geo-electrical soundings indicate an increasing depth to basement in northern direction up to 11 metres.
The middle layer shows an interpreted resistivity of $24 \Omega \mathrm{~m}$ which probably indicates weathered basement material with saline ground water.

DA 5.6.2.10. Idibo, Magera River (165/1-10)
m-m--------
The cross-section is located about one kilometre east of the village, The western terrace consists of weathered basement material and prospects for the construction of shallow wells are poor. The eastern terrace on the other hand is obviously of fluvial origin and consists of clayey sands and medium coarse sands underlain by sandy clay. The ground water quality of this terrace as well as that of the riverbed is good (E.C. values about $70 \mathrm{mS} / \mathrm{m}$ ) and because a hydraulic connection exists prospects for shallow well construction are favourable.

Geo-electrical soundings show an increasing depth to basement in eastern direction, the maximum being about 20 metres below ground water level (165/1-10-a).
The resistivity value of the middle layer varies between 15 and $36 \Omega \mathrm{~m}$ which indicates fresh waterbearing sandy layers or weathered basement with saline ground water. The first possibility, however, seems the most probable one.

FIGURE DA 5.6. 2-9 Cross-section 165/1-9 Tribedya Valley at NOUYAMI


FIGURE DA 5.6. 2-10 Cross-section 165/1-10 Magera Valley at IDIBO


DA 5.6.2.11. Mbili, Mbili River (165/1-11)
This cross-section is located about 400 metres $S E$ of the village centre. The riverbed itself consists of coarse sands underlain by gravels and depth to bedrock varies from nil up to a few metres and therefore handdug holes can only be found in deeper filled up parts. Because of evaporation and lack of recharge, ground water quality is very poor ( $\mathrm{EC}>250 \mathrm{mS} / \mathrm{m}$ ) and unsuitable for exploitation, apart from the fact that most parts of the riverbed fall dry halfway through the dry season.

No geo-electrical soundings were carried out at this location because of bedrock outcrops.

DA 5.6.2.12. Mbili, Mbili River (165/1-12)
This profile is located about 2 km south of the village along the track to the main road Morogoro-Dodoma. The cross-section shows a situation like the previous one.
The riverbed is situated between bedrock outcrops and the profile covers a deeper filled up part. Only one drilling was carried out and bedrock was struck at a depth of 1.5 m .
The observation well fell dry in mid-October so that this location is unsuitable for the construction of a shallow or river well.

No geo-electrical soundings were carried out.
DA 5.6.2.13. Mwandi, Berega River (165/1-13)
The profile is situated between Mwandi and its subvillage Mugambo, about 1200 metres south of Mwandi.
The riverbed consists of coarse to medium coarse sand underlain by gravelly layers. The southern terrace consist of coarse sands overlain by finer material, probably weathered basement.
The northern terrace consists of thick, coarse sand and gravelly sand layers overlain by clayey deposits. Because of a hydraulic connection with the riverbed, prospects for the construction of shallow wells are good on this terrace, although salinity contents tend to be high (about $125 \Omega \mathrm{mS} / \mathrm{m}$ ).
Likewise cross-section 165/1-3, this profile clearly demonstrates inflow of saline water from the southern terrace into the riverbed.

No geo-electrical measurements were satisfactory due to lateral heterogeneity caused by bedrock outcrops in the vicinity of this profile.

FIGURE DA 5.6. 2-11 Cross-section 165/1-11 Mbili Valley at MBILI


FIGURE DA 5.6. 2-12 Cross-section 165/1-12 Mbili Valley at MBILI
(s)

s)

FIGURE DA 5.6. 2-13 Cross-section 165/1-13 Berega Valley at MWANDI


DA 5.6.2.14. Ibindo, Berega River (165/1-14)
The cross-section covers the riverbed of the Berega River and its terraces downstream.
The whole riverbed consists of medium coarse and coarse gravelly sands. Hand drillings had to be stopped due to hard layers, probably gravels. The eastern terrace consists entirely of weathered basement without a hydraulic connection and is thus an unsuitable location for the construction of shallow wells.
The western terrace, on the other hand, consists mainly of medium coarse sand layers of reasonable thickness and, in spite of higher E.C. values ( $125-150 \mathrm{mS} / \mathrm{m}$ ), prospects for the construction of shallow wells are satisfactory in comparison with ground water of the riverbed.

Geo-electrical soundings indicate a $V$-shaped filled up valley with a maximum depth to bedrock of about 20 metres ( $165 / 1-14-\mathrm{b}$ ). The middle layer of the interpreted model shows resistivities of about $45 \Omega \mathrm{~m}$ which probably indicates a weathered zone with fresh gound water.

DA 5.6.2.15. Leshata, Magera River (146/3-1)
-------------.............
The cross-section is located 400 metres east of the village centre. The western terrace consists of weathered basement material without a hydraulic connection with the present riverbed and is therefore unsuitable for the construction of shallow wells.
The riverbed itself consists of medium coarse to coarse sands underlain by clayey sand with some gravels. The eastern terrace is good but slightly decreasing in western direction, probably due to inflow of saline water from older deposits or weathered material.

According to geo-electrical sounding results, depth to bedrock is slightly increasing in eastern direction up to 28 metres approximately. Interpreted resistivities of the middle layer vary between 38 and $70 \Omega \mathrm{~m}$ which indicates weathered basement with fresh ground water.


DA 5.6.2.16. Nguyami, Ndogomi River (165/1-14)
The location of the profile is west of the village, where the Ndogomi River crosses the road to Chakwale. Five boreholes were drilled across the approximately 80 m wide alluvial valley. No geo-electrical soundings were carried out here. The alluvial deposits are mainly coarse, angular sands and clayey fine sands.
Because the salinity of the ground water is very high, E.C's up to $450 \mathrm{mS} / \mathrm{m}$ having been measured, it is not possible to exploit the ground water.

DA 5.6.2.17. Chakwale, Chakwale River (164/2-1)
The profile was drilled in the northern part of the village, along the road to Ndogomi.
Two boreholes were drilled in the riverbed of the Chakwale River and one on each of the two terrace deposits along the riverbed.
The profiles showed mainly coarse, gravelly sands. No ground water was present in this profile.

DA 5.6.2.18. Chakwale, Kibedja River (164/2-2)
The profile is situated south-west of the village, along the road to Kibedja.
Highly saline ground water, EC's up to $750 \mathrm{mS} / \mathrm{m}$ were found in both the riverbed deposits as well as the terraces along the river.

FIGURE DA 5.6.2-16 Cross-section 165/1-15 Ndogomi Valley at Nguyami
(E)

$: \underbrace{5: 8}_{i}$
.-
FIGURE DA 5.6.2-18 Cross-section 164/2-2 Chakwale Valley at Chakwale


FIGURE DA 5.6.2-17 Cross-section 164/2-1
(w) Kibedya Valley at Chakwale
(10)

DA 5.7. WATER POTENTIAL
DA 5.7.1. Ground water potential
It was not always possible to establish the exact thickness and extension of the riverbed aquifer. Drillings often had to be stopped at limited depth because of caving in of the boreholes due to loose sandy material, or because hard layer (gravels, stones) were encountered in the boreholes. Most ground water will flow through the riverbed aquifer because of its very high permeability. However, ground water also flows through the terrace deposits bordering the riverbed aquifer and even through the weathered zone of the basement in the valleys.

The calculation of the ground water flow through the different cross sections presented in Fig. DA 5.6.2-1 to -18 was based on the areas of the ground-water saturated cross-sections through the riverbed and coarse textural terrace deposits.
Because of the limitations mentioned above, these areas are minimum areas and therefore the calculated ground water flow through the crosssections, are minimum flows.

The areas of the cross-sections were determined for situations at the beginning of the dry season (ground water levels equal to the ground surface), the end of the dry season of 1978 (beginning of the rainy season in mid-November), and a hypothetical extremely dry year (beginning of the rainy season mid-December).

Ground water flow through the cross sections, presented in Table DA 5.7.1-1, was calculated according to the formulae:

```
Q = K.i.A
Q ground water flow m
i = hydraulic gradient
K = permeability m/day
A = minimum area of cross section m}\mp@subsup{m}{}{2
```

The hydraulic gradients was estimated, using Figure DA 5.3.4-1; the longitudinal profile of the Berega river and its tributaries and assuming that the hydraulic gradient is equal to the gradient of the surface of the riverbed.

The permeability of the cross sections have been estimated by examination of the drilling samples.

Assuming a water demand of $40 \mathrm{l} /$ day/capita, and an average growth rate of the population of $5 \%$ a year, the water demand per village has been determined for 1978, 1988 and 1998.
The water demand figures are give in Table DA 5.7.1-1, together with the water potential data.

The water demand of the twelve listed villages in 1978 is $862 \mathrm{~m}^{3} / \mathrm{day}$, $1296 \mathrm{~m}^{3} /$ day in 1988 and $1724 \mathrm{~m}^{3} /$ day in 1998.

The minium ground water flow through the cross section at Ibindo, which location is close to the junction of the Berega river with the Chogowale River, is $2055 \mathrm{~m}^{3} /$ day.
Hence the total flow through the riverbed aquifer is sufficient to meet the total demands up to the year 1998.

The ground water potential near the individual villages is ample; only the Mbili riverbed near Mbili offers no prospects for exploitation of ground water because the aquifer dries up during the dry season.

Table DA 5.7.1-1 - Ground water potential and water demand

| crosssection | village | population 1978 | population <br> dentand <br> 1978 <br> $\mathrm{ra}^{3} /$ day | $\begin{aligned} & \text { popalation } \\ & \text { dewand } \\ & 1988 \\ & \mathrm{~m}^{3} / \mathrm{day} \end{aligned}$ | ```population demand 1998 m}\mp@subsup{}{}{3}/\mathrm{ day``` | k m/day | i | $\mathrm{A}^{\mathrm{m}}{ }^{\text {2 }}$ |  |  |  | $\begin{aligned} & \mathrm{Q} \\ & \mathrm{~m}^{3} / \text { day } \end{aligned}$ |  |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  |  |  |  |  |  |  |  | BDS | EDS | E EDS | BDS | EDS | E <br> EDS |
| 165/1-1 | Berega | 2890 | 116 | 174 | 232 | 450 | . 014 | 340 | 300 | 290 | 2150 | 1880 | 1825 |
| 165/l-2 | Mabula | 1069 | 43 | 65 | 86 | 300 | . 015 | 225 | 170 | 160 | 1010 | 765 | 720 |
| 3 |  |  |  |  |  | 225 | . 014 | 360 | 320 | 310 | 1135 | 1000 | 875 |
| 165/1-4 | Magera | 1567 | 63 | 95 | 126 | 200 | . 014 | 260 | 145 | 125 | 725 | 406 | 350 |
| 5 |  |  |  |  |  | 250 | . 014 | 190 | 120 | 95 | 800 | 420 | 332 |
| 165/1-6 | Kinyolisi | 800 | 32 | 48 | 64 | 350 | . 015 | 72 | 50 | 40 | 378 | 263 | 210 |
| 165/1-7 | Ndogemi | 1434 | 57 | 86 | 114 | 275 | . 015 | 65 | 43 | 38 | 267 | 177 | 157 |
| 165/1-8 | Madege | 2050 | 82 | 123 | 164 | 300 | . 015 | 89 | 68 | 63 | 400 | 306 | 283 |
| 165/1-9 | Mguyami | 1792 | 72 | 108 | 144 | 350 | . 015 | 138 | 167 | 100 | 725 | 562 | 525 |
| 165/1-10 | Ibido | 2715 | 109 | 164 | 218 | 400 | . 015 | 110 | 80 | 60 | 660 | 480 | 360 |
| 165/1-11 | Mbili | 821 | 33 | 50 | 66 | 400 | . 026 | 10 | 7 | - | 104 | 72 | - |
| 12 |  |  |  |  |  | 300 | . 026 | 7 | - | - | 55 | - | - |
| 165/1-13 | Mawandi | 2558 | 102 | 153 | 204 | 400 | . 014 | 390 | 370 | 350 | 2185 | 2075 | 1960 |
| 165/1-14 | Ibindo | 1558 | 62 | 93 | 124 | 350 | . 014 | 475 | 435 | 420 | 2327 | 2130 | 2055 |
| 146/3-1 | Leshata | 2285 | 91 | 137 | 182 | 250 | . 016 | 150 | 105 | 90 | 600 | 420 | 360 |
| total. |  |  | 862 | 1296 | 1724 |  |  |  |  |  |  |  |  |

HDS beginaing dry scason
EDS end dry season
EEDS end extreme dry scason

DA 5.7.2. Ground water quality
Ground water salinity, expressed in EC values (mS/m), varies widely over the area; not only per tributary but also locally. In general ground water of the riverbeds appeared to be of a better quality than that of its terraces. It could be proved that inflow occurs of mineralized water from older deposits or weathered basement material into the present riverbed which results locally in an increasing salinity (DA 5.6.2.1, $-6.2 .3,-6.2 .13,-6.2 .14$ ). Salinity probably increases in these deposits because of evaporation and a longer detention time.
On the basis of the salt content a decision can be made whether the ground water is suitable for the construction of a shallow or river well (EC < $200 \mathrm{mS} / \mathrm{m}$ ).
The table below reviews of the suitable shallow or river well sites per cross-section as mentioned in DA 5.6.2; a distinction has been made whether the site can be considered a long term or a temporary solution.

Table DA 5.7.2-1 - Average EC values of the ground water in different cross sections

| village | river | wel1* <br> location | average <br> EC value <br> mS/m | classification** |
| :---: | :---: | :---: | :---: | :---: |
| Berega | Berega | T | 105 | fair, permissible |
| Mabula (2) | Kibedya | T | 150 | doubtful, temp. perm. |
| (3) | Berega | R | 150 | doubtful, temp. perm. |
| Magera (4) | Magera | T | 100 | fair, permissible |
| (5) | Magera | T | 85 | fair, permissible |
| Kinyolisi | Kinyolisi | T | 75 | good, permissible |
| Ndogomi | Ndogomi | T | 130 | doubtful, temp. perm. |
| Madege | Ndogomi | T | 150 | doubtful, temp. perm. |
| Nguyami | Kibedya | - | 300 | poor, not permissible |
| Idibo | Magera | T | 70 | good, permissible |
| Mbili (11) | Mbili | R | 250 | poor, not permissible |
| Mbili (12) | Mbili | R | 80 | doubtful |
| Mwandi | Berega | T | 76 | fair, permissible |
| Ibindo | Berega | T | 120 | fair, permissible |
| Leshata | Magera | T | 40 | good, permissible |


| $*$ | $=$ | terrace |
| :--- | :--- | :--- |
|  | T | $=$ |
| R | riverbed |  |
| $* *$ | temp. perm. | $=$ |
| temporarily permissible |  |  |

This table clearly demonstrates that the ground water quality of the Kibedya River and its tributaries, the Ndogomi and Mbili rivers is of poor quality (EC values $>150 \mathrm{mS} / \mathrm{m}$ ), while ground water quality of the Magera and Kinyolisi Rivers on the other hand are of reasonable quality. Ground water quality of the Berega River itself is reasonable, although considerable differences occur locally due to inflow of mineralized water from older deposits or weathered basement material.
The high salinity content of the Kibedya River may be due to recharge of highly mineralized water from the residual soils of weathered basement material of which its catchment area mainly consists. Hardly any terrace of fluviatile origin occurs along this river.

Only the locations in the Kibedya and Mbili Rivers near Nguyami and Mbili respectively, appeared to be unsuitable for the construction of shallow wells because of inadmissible salt contents, while four locations (Mabula $2 x$, Ndogomi and Madege) turned out to be suitable for the construction of shallow or river wells as a temporary solution.

Chemical analyses of ground water from 13 different locations are given in table DA 5.7.2-2.
A complete analysis was made of the two samples with an EC higher than $200 \mathrm{mS} / \mathrm{m}$.
Main constituents are chloride, bicarbonite, calcium and magnesium (sodium and potassium were not analyzed).
In the other samples, only iron, fluoride and nitrate were analyzed. Fluoride and iron contents are very low.
In Magera a high nitrate content of the ground water was found.
Bacteriological contamination of the ground water in the riverbed aquifers is undoubtly present, because of the abundance of cattle in the Berega area.
A bacteriological test, carried out on a surface water sample from the Berega River, showed serious contamination by Faecal Coliforms.
It is recommended therefore, that shallow wells be constructed on the terraces and as far as possible from the riverbed.
The filters in shallow wells in both the riverbed aquifers and on the terraces will have to be placed as deep as possible.

Table DA 5.7.2-2 - Chemical analyses of ground water

| location cross-section no date of sampling |  | Berega $165 / 1-1$ $24-10-78$ | Mabula $165 / I-2$ $23-10-78$ | Mabula $165 / 1-3$ $23-10-78$ | Magera 165/1-4 23-10-78 | Magera <br> 165/1-5 <br> 23-10-78 | $\begin{aligned} & \text { Kinyolisi } \\ & 165 / 1-6 \\ & 25-10-78 \end{aligned}$ | $\begin{aligned} & \text { Ndogomi } \\ & 165 / 1-7 \\ & 24-10-78 \end{aligned}$ | Madege $165 / 1-8$ $24-10-78$ | $\begin{aligned} & \text { Idiba } \\ & 165 / 1-10 \\ & 23-10-78 \end{aligned}$ | $\begin{aligned} & \text { Mbili } \\ & 165 / 1-11 \\ & 17-11-78 \end{aligned}$ | Mwandi $\begin{aligned} & 165 / 1-13 \\ & 18-11-78 \end{aligned}$ | fbindo $165 / 1-14$ $17-11-78$ | Chakwate $25-10-78$ |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| water analysis unit |  |  |  |  |  |  |  |  |  |  |  |  |  |  |
| E.C | mS/ın | 140 | 160 | 140 | 75 | 110 | 60 | 175 | 190 | 70 | 325 | 110 | 90 | 300 |
| pil |  | 8.0 | 8.3 | 7.8 | B. 1 | 8.1 | 8.0 | 7.8 | 7.7 | 7.8 | 7.5 | 7.2 | 7.3 | 7.8 |
| calcium | $\mathrm{mg} / 1 \mathrm{CA}^{+*}$ |  |  |  |  |  |  |  |  |  | 8.8 |  |  | 23.2 |
| magues ium | 田/7. $\mathrm{Hg}^{++}$ |  |  |  |  |  |  |  |  |  | 95 |  |  | 68 |
| total hardness | $\mathrm{mg} / \mathrm{l} \mathrm{CaCO}_{3}$ |  |  |  |  |  |  |  |  |  | 610 |  |  | 860 |
| íron | $\mathrm{mg} / \mathrm{lfe}$ | 0.03 | 0.03 | 0.04 | 0.04 | 0.05 | 0.04 | 0.16 | 0.07 | 0.23 | 0.9 | 0.31 |  | 0.06 |
| manganese | mg/1 Mn | - | - | - | 0.1 | - | 0.3 | 0.8 | 2.5 | 1.8 |  |  |  | 0.1 |
| fluoride | mg/1 F | 0.8 | 0.6 | 0.8 | 0.3 | 0.7 | 0.5 | 0.3 | 0.4 | 0.4 | 0.7 | 0.8 | 0.8 | 0.7 |
| chloride | nig/l $\mathrm{Cl}^{-}$ |  |  |  |  |  |  |  |  |  | 730 |  | . | 500 |
| bicarbonate | $\mathrm{mg} / \mathrm{LHCO}{ }_{3}^{-}$ |  |  |  |  |  |  |  |  |  | 586 |  |  | 488 |
| nitrate | $\mathrm{mg} / \mathrm{LNO} \mathrm{N}^{-}$ | 3.1 | 3.0 | 3.1 | 2.1 | 28.8 | 2.0 | 6.2 | 6.6 | 0.4 | 12.4 | 3.1 | 3.0 |  |
| sulfate | $\mathrm{mg} / 1 \mathrm{SO}_{4}^{--}$ |  |  |  |  |  |  |  |  |  | 68 |  |  |  |
| phosphate | $\mathrm{mg} / 1 \mathrm{PO}_{4}^{---}$ |  |  |  |  |  |  |  |  |  | 0.17 |  |  |  |

DA 5.7.3.
Summary of conclusions
The recommendations for an improved water supply by means of shallow or river wells within the investigated sub-area are summarized in the following table.

Table DA 5.7.3-1 - Recommendations for improved water supply by means of shallow wells

| village | temporary <br> solution | long term solution | possibilities of a pumped, piped water supply | no possi- <br> bilities |
| :---: | :---: | :---: | :---: | :---: |
| Berega <br> Mabula <br> Magera <br> Kinyolisi <br> Ndogomi <br> Madege <br> Nguyami <br> Idibo <br> Mbili <br> Mwando <br> Ibindo <br> Leshata | x x | $\mathbf{x}$ <br> x <br> x <br> x <br> x <br> x $\mathbf{x}$ |  | x <br> x |

Shallow or river wells may not be possible in Ndogomi because of an insufficient ground water potential of the riverbed aquifer, while the Kibedya River has inadmissibly saline ground water near Nguyami (the riverbed of the Ndogomi River west of the village was investigated by the MWCP siting crew; considerable salt contents up to $400 \mathrm{mS} / \mathrm{m}$ were found there).
The Mbili River near Mbili suffers from insufficient ground water potential as well as high salinity. Because of the salt content ( $>125 \mathrm{mS} / \mathrm{m}$ ) of the rivers near Mabula and Madege, shallow or river wells can serve only as a temporary solution.
The MWCP siting crew carried out some additional drillings and complete sitings in villages which were not involved in this study; the results are summarized in the following table.

Table DA 5.7.3-2 - Results of the MWCP hand drilling in the Berega area

| village | siting completed siting not yet sufficient sites completed | unsuccessful siting |  |
| :---: | :---: | :---: | :---: |
|  |  | insufficient | $\left\lvert\, \begin{aligned} & \mathrm{EC} \\ & >200 \mathrm{mS} / \mathrm{m} \end{aligned}\right.$ |
| Magubike Kiegeya | x | x | x |
| Italagwe | x |  |  |
| Makuyu | x |  |  |
| Iyogwe | x |  |  |
| Kilama |  | x |  |
| Chakwale |  |  | x |

Chakwale was surveyed in order to establish the possibility of shallow or river wells near the village which could serve as a temporary solution during the occasional breakdowns of the Gairo gravity scheme. It appeared, however, that salinity contents of the Kibedya River exceed the maximum level of $200 \mathrm{mS} / \mathrm{m}$ (even up to $400 \mathrm{mS} / \mathrm{m}$ ).

The water supply of the villages Kiegeya and Kilame cannot be improved by means of shallow or river wells, because no suitable sites could be found in the alluvial deposits near these villages. Riverbeds fell dry halfway through the dry season while near Kiegeya high salt contents were found ( $>500 \mathrm{mS} / \mathrm{m}$ ).
These villages may be served by shallow or river wells as a temporary solution during part of the dry season; otherwise resettlement will have to be considered, especially for Kilama, because of its remote location in case of an improved water supply by means of a gravity scheme. The domestic water supply of the villages Italagwe, Makuyu and Iyogwe can be improved by shallow wells. The MWCP siting crew has selected sufficient sites to serve these villages.

Improved domestic water supply for the villages within the other sub-area, mentioned in DA 5.2.1., is not possible by means of shallow wells, mainly because of the insufficient ground water potential of the riverbeds near by these villages during the dry season. Shallow or river wells may serve as a temporary solution during part of the dry season, however. The only possibilities for an improved domestic water supply are gravity schemes.
In a separate study, the surface hydrology section of the MDWSP established the quantity of surface water available for domestic use in this area (Volume III).
Part of the area can be served by an improved and extended Gairo gravity scheme, while new gravity schemes also appear to be possible.

In this case, the villages Mbili, Nguyami and Ndogomi would also be connected to such schemes, instead of having their supply improved by means of shallow or river wells, as initially suggested.

D ATA

DATA

DATA (tables and figures)

DD 1 Borehole data
DD 2 Sample logs of existing boreholes
DD 3 Chemical analyses of ground water from boreholes
DD 4 Hydrogeological field data of existing ground water supplies
DD 5 Summary of MWCP-hand drilled holes, approved for construction of shallow wells
DD 6 Data of geo-electrical soundings
DD 7 The geo-electrical soundings and their interpretations
DD 8 Time-drawdown and time-recovery graphs of pumping tests from MDWSP-boreholes
DD 9 Summary of hand drillings in the Ngerengere area
DD 10 Chemical analyses of ground water from hand drillings in the Ngerengere area
DD 11 The geo-electrical soundings in the Berega area and their interpretations
DD 12 Well loggings with their corresponding lithological profiles; Berega area
DD 13 Fluctuations of ground water levels and EC's measured in piezometers; Berega area

DATA DD 1

BOREHOLE DATA

DD 1 - Borehole data (see last sheet of data for explanatiutt of abhreviations)


EC*: EC values calculated from TDS values according to: EC $=$ TDS $\times 0.7$


Continued 1

| we 11 no. | date | location | drill | type <br> of well | elev. GL above SL - | elev. GL above GL a | depth |  | geolagical fomalions pierced | casing |  | screeil |  | tength | type |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  |  |  |  |  |  |  | $\begin{aligned} & d \mathrm{rij} \\ & \text { migh } \end{aligned}$ | finish m BGL |  | dia inches | depth - BCL | dia <br> inches | tepth <br> a BCL |  |  |
|  |  | Dams site |  |  |  |  |  |  |  |  |  |  |  |  |  |
| BH 21/51 | 18.09-'51 | Tungi |  | explor. |  |  | 12.43 |  | All.-P.C. |  | no casing left |  |  |  |  |
| BH 28/52 | 15-11-'52 | Mkata, Vet. Camp 3 | pers | exploit. | 415 | 0.30 | 35.36 |  | Q/T | 6 | 0-35.66 |  |  |  |  |
| RN1 37/52 | 17-11-'52 | Itkata, Vet. | perc | explor. | 415 |  | 40.84 |  | P.C. | 8 | 0-28.04 |  |  |  |  |
| ВН' 2/53 | 25-01-'53 | Gairo | pere | exploit. | 1260 |  | 73.76 |  | P.C. | 8 | 0-19.40 |  |  |  |  |
| B61 3/53 | 10-05-'53 | Minor Settlean. Ngerengere | perc <br> perc | exploit. | 180 |  | 117.39 |  | Jurassic | 8 | 0-2494 |  |  |  |  |
| B4 23/53 | 25-07-'53 | Kidugal ${ }^{\text {o }}$ | perc | exploit. |  |  | 121.92 |  | Jurassic | 6 | 0-30.48 | 6 | 30.48-35.66 | 5.18 | RPsp |
| BH 37/54 | 2-10-'54 | Lloyds ridge | perc | exploit. | 488 | 0.30 | 31.69 |  | Q/r | 8 | 0-31.39 |  |  |  |  |
| BH 41/54 | 4-10-'54 | Chakwale | perc | explait. | 1080 |  | 42.06 |  | P.C. | 8 | 0-13.72 |  |  |  |  |
| BH 43/54 | 16-11-'54 | Wami Prison | perc | exploit. | 324 |  | 47.50 | 43.50 | Q/T |  |  |  |  |  |  |
| Bf 43a/54 | 24-12-'54 | - Wami Prison | perc | exploit. | 384 |  | 46.55 | 44.81 | Q/T | ${ }^{8}$ | 0-44.81 |  |  |  |  |
| BH 30/55 | 15-09-'55 | Madoto S.E. | perc | exploit. | 430 |  | 49.07 |  | Q/T | B | 0-48.46 |  |  |  |  |
| В4 40/55 | 5-10-'55 | Madoto S.E. | perc | exploit. | 430 |  | 33.53 | 33.57 | Q/T | 8 |  |  |  |  |  |
| BH1 52/55 | 5-01-'55 | Kidegal lo | perc | exploit. |  |  | 39.62 |  | Jurassic |  | no casiug left |  |  |  |  |
| BH 2/56 | 22-05-'56 | Kidugallo | perc | exploit |  |  | 122.23 |  | Jurassic | 8 | 0-29.93 |  |  |  |  |
| BH 6/56 | 19-03-'56 | Kidugalio | pers | exploit. |  |  | 111.56 |  | Jurassic | 6 | 0-56.38 |  |  |  |  |
| BH 19/56 | 4-10-'56 | Antorion S.E. | pere | exploit. | 440 |  | 47.85 | 47.85 | Q/T | 8 |  |  |  |  |  |
| BH 21/56 | 11-08-'56 | Kihonda S.E. | perc | exploit. | 525 |  | 74.07 |  | P.C. | 6 | 0-6.40 |  |  |  |  |
| BH1 23/56 | 6-09-'56 | Kimamba town | perc | exploit. | 435 |  | 40.23 | 40.23 | Q/T | 6 | 0-40.23 |  |  |  |  |
| BH 34/56 | 6-11-'56 | Kimanha | perc | exploit. | 440 | 0.52 | 53.43 | 53.43 | Q/T | 8 | 0-53.43 |  |  |  |  |
| BH1 37/56 | 7-12-'57 | Kimamba | perc | exploit. |  |  | 52.73 | 35.97 | Q/T | 8 | 0-35.97 |  |  |  |  |
| BH $42 / 56$ | 24-01-'57 | Kimamba S.E. Semtari S.E. | pere perc | exploit. |  |  | 53.64 | 53.34 | 9/T | 8 | 0-53.34 |  |  |  |  |
| BH $2 / 57$ | 28-02-'57 | $K$ kimamba | perc | exploit. | 440 | 0.60 | 52.43 | 52.43 | Q/T | 8 | 10-52.43 |  |  |  |  |
| BH 10/57 | 29-05-'57 | Kimamba | perr | exploit. |  |  | 56.39 | 51.82 | g/'T | 8 | 0-22.25 | 8 | 22.25-51.82 | 29.57 | g1'sp |
| B13 19/57 | 9-67-'57 | Madot: S.E. <br> Kimamha | pers perc | exploit. | 435 | 0,71 | 64.92 | 64.61 | $0 / T$ | 10 | 0-12.19 | 8 | 12.19-64.63 | 52.42 | $8 \mathrm{SP} \cdot \mathrm{s}$ |



Continued 2

| wellno. | date | 1ocation | drill | type of well | elev. GL ahove SL | elev. <br> GI. <br> above <br> GL. <br> II | depth |  | geological fornations pierced | casing |  | screen |  |  |  |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  |  |  |  |  |  |  | drill | finish m BGL. |  | dia <br> inches | depth <br> m BGL | dia <br> inches | dejth <br> m BGL | length | t;pe |
| 明 23/57 | 28-08-'57 | Kilosa | perc | exploit. |  |  | 32.93 |  | Q/T | 8 |  |  |  |  |  |
|  |  | East Africa <br> Sisal plant | perc <br> perc |  |  |  |  |  |  |  |  |  |  |  |  |
| Hil $28 / 57$ | 7-09-'57 | Kilosa | pere | exploit. | 488 |  | 32.61 | 32.61 | Q/T | B | 0-32.61 |  |  |  |  |
| BH 31/57 | 18.09-'57 | Kilosa | perc | exploit. | 478 | 0.30 | 32.61 | 32.61 | Q/T | 6 | 0-32.61 |  |  |  |  |
| 8H 32/57 | 2-10-'57 | Kilosa | perc |  | 480 |  | 33.53 | 32.92 | Q/T | 8 | 0-32.92 |  |  |  |  |
| Bi 15/59 | 30-12-'59 | Mikumi | perc | exploit | 515 |  | 92.05 | 92.05 | Karroo | B | $\begin{array}{r} 0-64.92 \\ 64.92-92.05 \end{array}$ |  |  |  |  |
| BH 17/59 | 19-05-'59 | Kimamba town | perc | exploit. | 435 | 0.15 | 35.36 | 35.36 | Q/T | 6 | 0-35.36 |  |  |  |  |
| B4 18A/59 | 5-06-'59 | Pangawe | perc | explor. |  |  | 85.04 |  | P.c. |  | no casing left |  |  |  |  |
| $\text { BH } 18 B / 59$ | 24-08-'59 | Pangawe | perc perc | explor. exploit | 515 | 0.90 | 49.07 71.63 | 68.58 | P.C. | 6 | no casing left |  |  |  |  |
| B4 23/59 | 18-09-'59 | Pangawe | perc | exploit. | 535 |  | 56.08 |  | P.C. | 8 | 0-19.81 |  |  |  |  |
| B ${ }^{\text {B }}$ 15/60 | 17-09-'60 | Rubeho | perc | explor. | 1335 |  | 75.29 |  | P.C. | 8 | 0-19.20 |  |  |  |  |
| B ${ }^{\text {B }}$ 16/60 | 2-11-'60 | Meshugi | perc | explor. |  |  | 62.79 |  | P.C. | 6 | 0-26.82 | 6 | 26.82- | 35.97 | gpsp |
| BH 17/60 | 28-10-'60 | Magole | perc | exploit. | 425 |  | 48.77 | 46.33 | Q/T | 6 | $0-30.48$ | 6 | 30.48-46.33 | 15.85 | girsp |
| 明 31/60 | 3-04-'61 | Kimantra town | perc | exploit. | 440 |  | 69.50 |  | Q/T |  |  |  |  |  |  |
| Pf 34/60 | 2-12-'60 | Rubeho | pere |  | 1305 | 0.60 | 68.88 |  | P.c. | 6 | 0-68.88 |  |  |  |  |
| $\begin{aligned} & \text { BH } 37 / 60 \\ & \text { BH } 23 / 61 \end{aligned}$ | $\left\{\begin{array}{c} 22-01-61 \\ 2-10-161 \end{array}\right.$ | Kisitwe Kimamba towil | perc | explor. | 1350 | 0.2 | 67.97 189.59 | $\begin{array}{r} 67.97 \\ 185.32 \end{array}$ | $\begin{array}{\|l} \text { P.c. } \\ \text { Q/T } \end{array}$ | $\begin{aligned} & 6 \\ & 6 \end{aligned}$ | 0-67.97 |  |  |  |  |



Continued 3

| wellno. | date | location | drill | type <br> of well | elev. <br> GI. <br> above <br> SL | elev. <br> Gl. <br> alhove <br> GL <br> 輯 | depth |  | geological formations pierced | casing |  |  | screes |  |  |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  |  |  |  |  |  |  | $\begin{aligned} & \text { drill } \\ & \text { m BGL } \end{aligned}$ | $\begin{aligned} & \text { finish } \\ & \text { m BGI. } \end{aligned}$ |  | dia inches | depth m BGC | dia inches | depth <br> m BGL | length | t ype |
| BH 30/61 | 28-01-761 | Sentari S.E. Kimataba | perc | exploit. | 440 | 0.35 | 62.80 | 62.80 | Q/T | 6 | 0-37.80 | 6 | 37.80-62.80 | 25.00 | ripsp |
| BII 6/62 | 31-03-'62 | Mvomero | perc | exploit. | 410 | 0.50 | 62.48 | 61.56 | Q/T | 6 | ? | 6 | every other 3 m | 30.78 | ppsp |
| BH 7/62 | 8-12-62 | Saga Saga |  |  |  |  | 121.92 |  | Jurassic | 6 6 6 6 | $0-10.67$ $15.24-36.58$ $39.62-60.96$ $64.62-68.58$ | 6 6 6 | $10.67-15.24$ $36.58-39.62$ $60.96-64.62$ | $\begin{aligned} & 4.57 \\ & 3.64 \\ & 3.66 \end{aligned}$ | gpsp kpsp gpsp |
| BH $8 / 62$ | 8-04-'6. ${ }^{\text {7 }}$ | School Mkulazi |  |  | 165 |  | 107.90 |  | Jurassic |  | no casing left |  |  |  |  |
| BH 9/62 | 7-11-"62 | Gronero |  |  | 210 | 0.58 | 111.56 | 106.58 | Jurassic | 6 | $0-30.48$ | 6 | 30.48-106.68 | 76.2 | gpsp |
| 84 5/6.3 | 21-02-'63 | Higuzi Mgugu | perc |  | 435 |  | 37.19 |  | P.c. | 6 | 0-9.14 | 6 | 9.34-20.73 | 11.59 | g.psp |
| Bl\| 18/6.3 | 24-12-'63 | Wami Prison | pers | exploit. | 381 |  | 62.48 | 62.48 | Q/T | $\begin{aligned} & 6 \\ & 6 \\ & 6 \end{aligned}$ | $\begin{array}{r}\text { a-29.0 } \\ 35.7 \\ 61.6 \\ \hline 1.62 .5\end{array}$ | 6 | $29.0-35.7$ 51.5 | 6.7 10.1 | $\begin{aligned} & \text { Rpsp } \\ & \text { gpsp } \end{aligned}$ |
| B ${ }^{\text {d }}$ 4/64 | 30-05.'64 | Kikiboga Hotel | perc |  | 580 |  | 75.18 |  | Karroo | 8 | 0-33.83 | 8 | 33.83-47.55 | 13.72 | mpsp |
| R12 13/64 | 31-07-764 | Rigohoga | perc |  | 560 | 1.58 | 43.28 |  | Karroo | 6 | 0-31.39 | 6 | 31.39-44.80 | 13.42 | gpsp |
| 8H1 31/64 | 30-01-"65 | Ringolwira | pere | exploit. | 440 |  | 76.20 |  | P.C. | 8 | 0-13.41 | 8 | 13.41-23.16 | 9.75 | gpsp |
| BH 18/65 | 19-09-'65 | Magubike | perc | exploit. | 740 | 0.86 | 76.20 |  | P.c. |  |  | 6 | 0-61.57 | 61.57 | gpsp |
| Br 33/65 | 31-02-'65 | Gairo | perc | exploit. | 1270 |  | 90.22 | 90.22 | P.C. | 6 |  | 6 |  | 58.52 | ¢psp |
| Pl 14/66 | 28-08-'66 | Turiani | perc | exploit. | 380 |  | 36.60 | 36.60 | P.C. | 8 | 0-6.00 | 6 | 6.00-36.00 | 30.00 | kpsp |
| Bi $41 / 67$ | 1-11-'67 | Mikese | perc |  | 430 | 1.06 | 60.69 | 60.69 | P.C. |  | $\begin{array}{r} 0-2.74 \\ 12.50-22.56 \end{array}$ |  | 2.74-12.50 | 9.76 |  |
| BH 49/67 | 1967 | Pangawe | pere |  | 480 |  | 22.25 |  | P.C. |  |  |  |  |  |  |
| BH 17/68 | 1-07-768 | Ngerengere | pere | exploit. | 384 |  | 122.00 |  | Q/T | 6 |  | 6 |  |  |  |
| Bll $72 / 68$ | 28-01-'69 | Ki.losa |  | exploit. | 488 | 1.74 | 27.12 |  | Q/T | 6 | $\begin{array}{r} 0-21.03 \\ 24.08-27.12 \end{array}$ | 6 | 21.03-24.08 | 1.05 | misp |



Continued 4

| $\begin{aligned} & \text { well } \\ & \text { no. } \end{aligned}$ | date | location | drill | type <br> of well | elev. GL above SL m | elev. cl. ahove GL | depth |  | geological fornations pierced | casing |  |  | screen |  | type |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  |  |  |  |  |  |  | drill <br> m BGL | finish m BGL |  | dia <br> inches | depth <br> - BGL | dia <br> inctues | $\begin{aligned} & \text { drpth } \\ & \text { W BG. } \end{aligned}$ | length |  |
| BH 75/6B | 1968 | Kilosa | perc | exploit. | 488 |  | 38.00 |  | Q/T |  |  |  |  |  |  |
| BH 78/68 | 1968 | Kilosa | pere | exploit. |  |  |  |  | Q/T |  |  |  |  |  |  |
| BH 17/69 | 1969 | Htibua | pere | exploit. | 375 |  | 33.60 | 33.60 | Q/T | 6 | 0-9.70 | 6 | 9.70-33.60 | 23.90 | .ppsp |
| BH 26/69 | 1969 | Mtibwa | perc | exploil. | 375 | 0.66 | 47.90 |  | 0/T | 6 | 0-8.40 | 6 | 8.40-47.90 | 40.30 | ¢ 2 psp |
| B $46 / 69$ | 1969 | Mikumi | perc |  | 500 |  | 86.60 |  | Karroo |  |  |  |  |  |  |
| Bll $47 / 69$ | 1969 | Htibwa |  |  |  |  |  |  | Q/T |  |  |  |  |  |  |
| Вн 56/69 | 19-09-'69 | Wakwari | pere |  | 345 |  | 76.50 | 44.80 | Q/T | 8 | 0-35.10 | 8 | 35. 10-44.80 | 9.70 | gpsp |
| Bll 58/69 | 20-09-169 | Wakwari | perc |  | 350 |  | 71.30 | 71.30 | Q/T | 6 | 0-33.80 | 6 | 33.80-52.70 | 18.90 | $\mathrm{grsp}^{\text {sp }}$ |
|  |  |  |  |  |  |  |  |  | Q/T | 6 | 52.70-66.80 | 6 | 66.80-70.10 | 3.30 | gpsp |
| Bif 4/70 | 1970 | Mibigili Prison | pere | exploit. | 384 | 43.60 | 43.60 |  | Q/T | 6 | $\begin{array}{r} 0-35.10 \\ 39.60-43.60 \end{array}$ | 6 | 35.10-39.60 | 4.50 | ppsp |
| BII 21/70 | 1970 | Lomo, Kimamba Usagara Est. |  | exploit. | 478 |  | 24.00 |  | 0/T |  |  |  |  |  |  |
| BH $22 / 70$ | 21-01-'70 | Ulsagara Est. | perc | exploit. | 478 |  | 22.86 | 22.86 | Q/T |  |  |  |  |  |  |
| BII 97/69 | 1969 | Sentari Kilosa | perc | exploit. |  |  | 43.00 |  | Q/T |  |  |  |  |  |  |
| B11. $46 / 70$ | 1970 | 1 longa |  |  |  |  |  |  | Q/T |  |  |  |  |  |  |
| B11 46/61 | 14-07-'71 | Kidugallo | pere | exploit. | 255 |  | 201.78 |  | Jurassic |  |  |  | 64.62-70.72 | 63.70 | spsp |
| B ${ }^{\text {7 }} 73 / 71$ | 14-08-'71 | Mtibwa | pere | exploit. | 310 |  | 106.70 | 34.40 | 9/T | 8 | 0-28. 30 | 8 | 28.30-34.40 | 6.10 | 9psp |
| Bit 75/71 | 14-09-'71 | Hikumi | prere | exploit. | 520 | 0.91 | 76.20 |  | Karroo | 6 | 0-15.84 | 6 | 15.84-40.23 | 24.39. | gisp |
| Bil 111/71 | 1971 | Mikumi | perc |  | 560 |  |  |  |  |  |  |  |  |  |  |
| B ${ }^{\text {a }} 112 / 71$ | $8-1171$ | Mikumi | perc | exploit. | 550 |  | 46.93 |  | Karroo |  |  | 6 | 9.14-21.37 | 12.0 | gesp |
| BH 123/71 | 1971 | Kidatu S.E. | perc | exploit. |  |  | 69.80 |  | Q/T-Karroo |  |  |  |  |  |  |
| BH1 136/71 | 1971 | Mikumi | pere | exploit. |  |  | 62.48 |  | Q/T-Karroo | ${ }^{6}$ | 0-34.78 | 6 | 34.78-46.94 | 12.19 | ${ }^{3 \mu s p}$ |
| Bll 14/72 | 12-05-'72 | Mt ilwa | perc | exploit. | 375 |  | 198.10 | 36.60 | Q/T-P.C. | 8 | 0-9.10 | 8 | 9.10-21.30 | 12.2 | gpsp |
|  |  |  |  |  |  |  |  |  |  |  | $\begin{aligned} & 21.30-27.40 \\ & 35.50-36.60 \end{aligned}$ |  | 27.40-35.50 | 6.1 | gesp |
| BH1 38/72 | 1972 | Mapadu | perc |  | 540 |  | 135.33 |  | P.C. |  | no rasing left |  |  |  |  |
| B31 79/2 | 7-07-' 72 | Saha Saba Morogoro | pere |  | 520 |  | 16.80 |  |  |  |  |  |  |  |  |
| RH195/72 | 10-11.'73 | Kimanba town | perc |  | 440 |  | 137.16 |  | $\mathrm{Q} / \mathrm{T}$ | 8 | $0-210.00$ |  |  |  |  |
|  |  |  |  |  |  |  |  |  |  | 6 | 100.0-120.20 | 6 | 120.20-129.70 | 9.50 | J.s. |



Continued 5

P.C. $=$ Precamhian

Q/T $=$ Quarternary-Tertiair
All $=$
$S=$
SD $=$ Screen length
$\mathrm{KD}=$ Tramsmissivity
TOS $=$ Total dissolvel solids

GL $=$ Ground level
SL $=$ Sea level
RGL $=$ Below ground ievel
Swl $=$ Standing water Ipvel

## Q $=$ Discharge

Av.K $=$ Average hytraulis comilurtivit
$E C=$ Electrical conductivity


DATA DD 2

SAMPLE LOGS FROM EXISTING BOREHOLES











$\begin{array}{ll}\text { Borchole No.: 17/47 } & \text { Mapsheet: 183/2 } \\ \text { Location: Kigolwira } & \text { Co-ord. : }\end{array}$
Prison










































Location: Kimamba Co-ord. : Usagara S.E.


$-100$
















## DATA DD 3

CHEMICAL ANALYSES OF GROUND WATER FROM BOREHOLES

Table DD 3 - Chemical analyses of ground water from boreholes

| borehole no. | location | depth | date of sampling | date of analysis | spec. cond. at $25^{\circ} \mathrm{C}$ | ph | $\mathrm{Ca}^{2+}$ | $\mathrm{Hg}^{2+}$ | TII, | Fe | Mn | NH ${ }^{+}$ | $\mathrm{Ha}^{+}$ | $\mathbf{x}^{+}$ | $\mathbf{F}^{-}$ | $\mathrm{Cl}^{-}$ | HCO | H0 | N0 | $\mathrm{so}^{2-}$ | $\mathrm{PO}_{4}^{3-}$ |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  |  | * |  |  | ms/m |  | Pp ${ }^{\text {m }}$ |  | $\stackrel{\mathrm{CaCO}}{3}^{\mathrm{Cpm}}$ |  |  |  |  |  | part | per | illion |  |  |  |  |
| 2/53 | Gairo |  |  | 11-02-53 |  |  |  |  |  |  |  |  |  |  |  | 472 | 482 |  |  | 110 |  |
|  |  |  | 11-02-61 | 1-03-61 | 1982 | 7.9 | 72.8 | 123.3 | 696 |  |  | 0.4 |  |  | 0.1 | 520 | 296 |  |  |  |  |
| 3/53 | Ngerengere |  |  | 7.08.53 |  |  |  |  |  |  |  |  |  |  | 6.0 | 432 | 388 |  |  | 406 |  |
| 23/53 | Ridugallo |  |  | 12-08-53 |  |  |  |  |  |  |  |  |  |  | 0.6 | 330 | 802 |  |  | 136 |  |
| 37/54 | Kimamba |  |  | 12-10-54 |  | 8.0 | 63 | 107 |  |  |  |  |  |  | 1.2 | 65 | 960 |  |  | 360 |  |
| 43A/54 | Wani |  |  |  |  | 8.5 | 21 | 35 |  |  |  |  |  |  | 1.1 | 9 | 1240 |  |  | 3 |  |
| 30/55 | Hadoto |  | 27-09-55 | 25-10-55 |  | 7.5 | 43 | 53 |  |  |  |  |  |  | 0.3 | 32 |  |  |  | 86 |  |
| 40/55 | Kimamba |  | 19-10-55 | 2-11-55 |  | 6.5 | 55 | 32 |  |  |  |  |  |  | 0.2 | 7 |  |  |  |  |  |
| 2/56 | Kidagallo |  |  | 21-05-56 |  | 7.5 | 125 | 23 |  |  |  |  |  |  | nil | 83 |  |  |  | 37 |  |
| 6/56 | Kidugallo |  | 14-12-67 | $30-12-67$ $19-04-56$ | 119 | 8.1 7.5 | 522 | 67 | 334 |  |  | 日il | 98 | 4.2 | ${ }_{1}{ }_{1}$ | 84 708 | 304 | nil | nil | 458 |  |
| 19/56 | Kimamba |  |  | 26-10-56 |  | 7.5 | 93 | 56 |  |  |  |  |  |  | 1 | 94 | 430 |  |  | 458 |  |
| 21/56 | Kihonds |  |  | 22-08-56 |  | 7.5 | 15 | 9 |  |  |  |  |  |  | 0.2 | 29 | 100 |  |  | nil |  |
| 23/56 | Kimamba |  |  | 7-09-56 |  | 7.0 | 98 | 120 |  |  | - |  |  |  | 4 | 93 | 720 |  |  | 270 |  |
|  |  |  | 30-09-63 | 8-11-63 | 63 | 8.3 | 94 | nil | 236 |  |  |  |  |  | 0.4 | 28 | 252 |  |  |  |  |
| 34/56 | Kimamba |  |  | 20-11-56 |  | 7.5 | 46 | 47 |  |  |  |  |  |  | 1 | 33 | 430 |  |  | 72 |  |
| 42/56 | Kimambs |  |  | 11-02-57 |  | 7.5 | 69 | 34 |  |  |  |  | 65 | 3 | 1 | 28 | 49 |  |  | 89 |  |
| 2/57 | Kimamba |  |  | 11-03-57 |  | 7.5 | 58 | 39 |  |  |  |  | 65 | 3 | 1 | 22 | 427 |  |  | 77 |  |
| 10/57 | Kinaeba |  |  | 6-05-57 |  | 7.5 | 58 | 27 |  |  |  |  | 37 | 3 | 0.6 | 11 | 290 |  |  | 100 |  |
| 19/57 | Hadoto |  |  | 19-07-57 |  | 7.5 | 79 | 71 |  |  |  |  | 88 | 1.5 | 0.4 | 65 | 525 |  |  | 176. |  |
| 23/57 | Kiloss |  |  | 6-09-57 |  | 7.5 | 293 | 46 |  |  |  |  | 285 | 8 | 1 | 306 | 345 |  |  | 781 |  |
|  |  |  |  |  |  |  | 240 | 57 | 840 |  |  | 0.24 | 440 |  |  | 244 | 352 | nil | nil | 823 |  |
|  |  |  |  |  |  |  | 300 | 36 | 900 |  |  | nil | 340 |  |  | 292 | 256 | nil | nil | 809 |  |
|  |  |  |  |  |  |  | 288 | 50 | 930 |  |  | nil | 380 |  |  | 292 | 272 | nil | nil | 813 |  |
| 29/57 | Kilosa |  |  | 10-09-57 |  |  | 97 | 72 |  |  |  |  | 154 | 1 | 1 | 163 | 439 |  |  | 223 |  |
| 15/59 | Mikumi |  |  | 21-01-60 |  | 7.0 | 90 | 80 |  |  |  |  |  |  | 0.4 | 140 | 710 | nil |  | 76 |  |
|  |  |  |  | 21-01-60 |  | 7.0 | 120 | 90 |  |  |  |  |  |  | 0.3 | 210 | 780 |  |  | 104 |  |
|  |  |  | 20-05-63 | 25-05-63 | 23 | 6.8 | 24 | 9 | 96 |  |  | nil |  |  | 0.4 | 8 | 112 | nil | nil |  |  |
|  |  |  | 20-05-63 | 25-05-63 | 25 | 6.7 | 25 | 12 | 110 |  |  | mil |  |  | 0.4 | 4 | 120 | $n i 1$ |  |  |  |
| 17/59 | Kimamba |  |  | $\begin{array}{r} 21-05-59 \\ 4-07-59 \end{array}$ |  | 7.5 | 146 | 137 |  |  |  |  | 280 | 2 | 3 | 200 |  |  |  |  |  |
|  |  |  |  | 10-07-57 |  | 7.5 | 160 | 140 |  |  |  |  | 240 |  | 2 | 170 |  |  |  | 740 |  |
|  |  | 10-14 | 10-10-63 | 30-10-63 | 339 | 7.9 | 78 | 202 | 1036 |  |  | nil |  |  | 0,5 | 224 | 660 | nil | 21 |  |  |
|  |  | 16-20 | 10-10-63 | 31-10-63 | 317 | 7.9 | 73 | 202 | 1028 |  |  | nil |  |  | 0.5 | 224 | 660 | nil | 21 |  |  |
|  |  | 20-23 | 10-10-63 | 31-10-6.3 | 277 | 7.9 | 57 | 227 | 1072 |  |  | nil |  |  | 0.5 | 224 | 704 | nil | 21 |  |  |
|  |  | 23-26 | 10-10-63 | 31-10-63 | 317 | 7.9 | 72 | 217 | 1084 |  |  | nil |  |  | 0.5 | 224 | 784 | nil | 21 |  |  |
|  |  | 29-32 | 10-10-6.3 | 31-10-63 | 276 | 7.9 | 58 | 220 | 1056 |  |  | nil |  |  | 0.5 | 224 | 704 | Hil | 21 |  |  |



Continued 1

| borehole no. | location | depth | date of sampling | date of anslysis | spec. cond. at $25^{\circ} \mathrm{C}$ | pH | $\mathrm{Ca}^{2+}$ | $\mathrm{Hg}^{2+}$ | TH' | Fe | Mn | NH ${ }^{+}$ | $\mathrm{Ha}^{+}$ | $\mathbf{K}^{+}$ | $\mathrm{F}^{-}$ | $\mathrm{Cl}^{-}$ | HCO | NO | HO | $50^{2-}$ | $\mathrm{PO}_{4}^{3-}$ |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  |  | $\pm$ |  |  | ms/m |  | ppm |  | $\stackrel{\mathrm{Ppma}_{3}}{\mathrm{CaCO}_{3}}$ |  |  |  |  |  | part | per | million |  |  |  |  |
| 17/59 | Kimamba | 25-30 | 10-10-60 | 31-10-60 | 276 | 8.0 | 64 | 215 | 1056 |  |  | nil |  |  | 0.5 | 244 | 592 | nil | 21 |  |  |
|  |  | 15 | 30-09-63 | 8-11-63 | 294 | 7.9 | 104 | 194 | 1070 |  |  | nil |  |  | 0.3 | 228 | 664 | nil | 17 |  |  |
|  |  | 21 | 30-09-63 | 8-11-63 | 280 | 7.7 | 132 | 180 | 1080 |  |  | nil |  |  | 0.4 | 204 | 736 | nil | 18 |  |  |
|  |  | 30 | 30-09-63 | 8-11-63 | 266 | 7.8 | 109 | 170 | 980 |  |  | nil |  |  | 0.3 | 168 | B08 | nil | 10 |  |  |
| 19/59 | Mikumi |  | 8-10-59 | 14-10-59 | 151 | 7.3 | 87 | 76 | 536 | 0.1 |  | nil | 115 |  | 0.6 | 148 | 612 | nil | nil |  |  |
| $17 / 0$ | Magole |  |  | 17-11-60 | 72 | 8.2 | 22 | 33 | 192 | nil | nil |  |  |  | 0.3 | 44 | 356 |  |  |  |  |
| 31/60 | Kimama | 10 | 3-12-60 | 14-12-60 | 101 | 8.4 | 41 | 80 | 438 |  |  | nil |  |  | 9.0 | 108 | 288 | 4.0 | 2.2 |  |  |
|  |  | 45 52 | 11-01-61 | 27-01-61 | 58 | 8.1 | 43 | 36 | 256 |  |  | nil |  |  | 0.5 | 20 | 200 168 | nil | ${ }_{2.4}$ |  |  |
|  |  | 52 | 11-01-61 | 27-01-61 | 63 | 8.1 | 43 | 36 | 256 |  |  | nil |  |  | 0.4 | 64 | 168 | nil | 2.4 |  |  |
|  |  |  | 23-01-61 | 3-02-61 | 59 | 7.8 | 40 | 31 | 230 |  |  | nil |  |  | 0.8 | 40 | 1212 | nil | nil |  |  |
| $\begin{aligned} & 34 / 60 \\ & 37 / 60 \end{aligned}$ | Rubeho Kisitwe |  | 11-02-61 | 1-03-61 | 240 | 8.1 | 145 | 128 | 900 |  |  | nil |  |  | 0.4 | 668 | 188 | nil | nil |  |  |
| 23/61 | Kimamba | 10 |  | 23-06-61 |  | 8.4 |  |  | 210 |  |  | nil |  |  | 1.2 | 60 | 332 |  | nil |  |  |
|  |  | 17 |  | 23.06.61 |  | 8.1 |  |  | 720 |  |  |  | * |  | 0.8 | 208 | 34.4 | nil | nil |  |  |
|  |  | 34 |  | 23-01-61 |  | 8.1 |  |  | 374 |  |  | nil |  |  | 0.6 | 76 | 268 | nil | nil |  |  |
|  |  |  | 26-07-61 | 1-08-61 | 110 | 8.1 | 59 | 61 | 406 |  |  | ail |  |  | 0.4 | 104 | 204 | ni1 | nil |  |  |
|  |  |  | 26-07-61 | 1-08-61 | 70 | 8.1 | 44 | 36 | 256 |  |  | nil |  |  | 0.4 | 36 | 196 | nil | nil |  |  |
|  |  |  | 26-07-61 | 31-07-61 | 61 | 8.1 | 40 | 35 | 244 |  |  | ni1 |  |  | 0.4 | 44 | 172 | nil | nil |  |  |
|  |  |  | 9-08-61 | 21-08-61 | 122 | 8.2 | 61 | 74 | 462 |  |  | nil |  |  | 0.3 | 100 | 300 | nil | nil |  | . |
|  |  |  | 9-08-61 | 21-08-61 | 70 | 8.3 | 51 | 32 | 262 |  |  | nil |  |  | 0.3 | 44 | 228 | nil | nil |  |  |
| 30/61 | Kimamba |  | 17-01-62 | 27-01-62 |  | 8.0 | 99 | 137 | 822 |  |  | nil |  |  | 0.6 | 208 | 644 | nil | nil |  |  |
| 6/62 | Hvome ro |  | 16-03-62 | 20-03-61 | 62 | 7.7 | 51 | 20 | 248 | nil | nil | nil |  |  | 0.5 | 120 | 156 | nil | nil |  |  |
|  |  |  | 2-05-62 | 10-05-62 | 260 | 8.2 | 148 | 78 | 696 | nil | nit | nil |  |  | 0.2 | 392 | 128 | nil | 4.8 |  |  |
| 7/62 | Saga Saga |  | 7.06-62 | 16-06-62 | 1058 | 8.0 | 522 | 538 | 3620 |  |  | 0.5 |  |  | 0.6 | 872 | 444 | 0.1 | 4.2 |  |  |
|  |  |  | 5-07-62 | 16-07-62 | 995 | 8.0 | 552 | 576 | 3780 |  |  | nil |  |  | 0.6 | 964 | 348 | nil | nil |  |  |
|  |  |  | 26-07-62 | 2-08-62 | 312 | 8.4 | 128 | 109 | 176 |  |  |  |  |  |  |  |  |  |  |  |  |
|  |  |  | 17-08-62 | 24-08-62 | 581 | 8.2 | 230 | 222 | 1500 |  |  | 0.5 |  |  | 2.0 | 716 | 336 | njil | nil |  |  |
| 8/62 | Hkulazi | 102 | 13-02-62 | 20-02-62 | 227 | 8.3 | 47 | 85 | 480 |  |  | nil |  |  | nil | 392 | 412 | nil | nil |  |  |
| )762 | Gronero |  | 9-07-62 | 14-07-62 | 505 | 8.0 | 43 | 30 | 364 |  |  | nil |  |  | 0.6 | 396 | 148 | nil | 0.6 |  |  |
|  |  | 105 | 17-08-62 | 24-08-62 |  | 8.3 | 13 | 10 | 880 |  |  | nil |  |  | nil | 172 | 628 | nil | nil |  |  |
|  |  |  | 18-12-72 | $2-01-63$ $4-09-63$ | 120 | 8.3 | 14 | 15 | 100 |  |  | nil |  |  | 4.0 | 156 | 588 | mil | nil |  |  |
| 2/63 | Wami | 37 | $29-08-63$ $29-08-63$ | 4-09-63 | 277 273 | 8.6 8.6 | 7 | 15 | 160 | nil | nil | nil |  |  | 2.0 2.0 | 368 408 | 1040 | nil | nil |  |  |
|  |  | 49 | 29-08-63 | 4-09-63 | 275 | 8.6 | 6 | 14 | 148 | nil | nit | nit |  |  | 2.0 | 344 | 976 | nil | nil |  |  |


| borehole no. | cations <br> anions | TDS | $\mathrm{H}_{2} \mathrm{~S}$ | $\mathrm{O}_{2}$ | SS | turb | color | coliform | Cu | Zn | $\mathbf{P b}$ |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  |  |  | PPm |  | mg/1 | ftu | $\mathrm{Pt} / \mathrm{Co}$ | count $100 \mathrm{ml}$ |  | ppa |  |
| 17/59 |  | $\begin{aligned} & 2310 \\ & 1995 \\ & 1982 \\ & 1602 \end{aligned}$ |  |  |  |  |  |  |  |  |  |
| $\begin{aligned} & 19 / 59 \\ & 17 / 0 \\ & 31 / 60 \end{aligned}$ |  | 917 445 |  |  |  |  |  |  | $n i 1$ | nil | nil. |
|  |  | 760 |  |  |  |  |  |  |  |  |  |
|  |  | 364 |  |  |  | * |  |  |  |  |  |
|  |  | 455 378 |  |  |  |  |  |  |  |  |  |
| $\begin{aligned} & 34 / 60 \\ & 37 / 60 \\ & 23 / 61 \end{aligned}$ |  | 1848 |  |  |  |  |  |  |  |  |  |
|  |  | 760 |  |  |  |  |  |  |  |  |  |
|  |  | 1477 |  |  |  |  |  |  |  |  |  |
|  |  | 691 |  |  |  |  |  |  |  |  |  |
|  |  | 844 |  |  |  |  |  |  |  |  |  |
|  |  | 413 |  |  |  |  |  |  |  |  |  |
|  |  | 428 831 |  |  |  |  |  |  |  |  |  |
|  |  | 8381 |  |  |  |  |  |  |  |  |  |
| $\begin{aligned} & 30 / 61 \\ & 6 / 62 \end{aligned}$ |  | 1467 |  |  |  |  |  |  | nil | nil |  |
|  |  | 542 |  |  |  |  |  |  | $n \mathrm{nil}$ | nil | nil |
|  |  | 1650 |  |  |  |  |  |  | nil | nil | nid |
| 7/62 |  | 7008 |  |  |  |  |  |  |  |  |  |
|  |  | 6764 |  |  |  |  |  |  |  |  |  |
|  |  | 5442 |  |  |  |  |  |  |  |  |  |
| $\begin{aligned} & 8 / 62 \\ & 9 / 62 \end{aligned}$ |  | 1150 |  |  |  |  |  |  |  |  |  |
|  |  | 3751 |  |  |  |  |  |  |  |  |  |
|  |  | 1290 |  |  |  |  |  |  |  |  |  |
|  |  | 1245 |  |  |  |  |  |  |  |  |  |
| 2/63 |  | 1769 |  |  |  |  |  |  | nil | nil | nil |
|  |  | 1880 |  |  |  |  |  |  | nil | nil | nil |
|  |  | 1639 |  |  |  |  |  |  | nil | nil | nil |

Continued 2

| borehole no. | location | depth | date of sampling | date of analysis | spec. cond. pt $25^{\circ} \mathrm{C}$ | pH | $\mathrm{Ca}^{2+}$ | $\mathrm{Hg}^{2+}$ | TH | Fe | Ma | M ${ }^{+}$ | $\mathrm{Ha}^{+}$ | $\mathbf{K}^{+}$ | $\mathbf{F}^{-}$ | $\mathrm{Cl}^{-}$ | HCO | HO | \%0 | $\mathrm{so}^{2-}$ | $\mathrm{PO}_{4}^{3-}$ |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  |  | m |  |  | m5/m |  | Ppram |  | $\begin{aligned} & \mathrm{ppan} \\ & \mathrm{CaCO}_{3} \\ & \hline \end{aligned}$ |  |  |  |  |  | part | per | Aillio |  |  |  |  |
| 5/63 | Mgugu |  | 25-05-63 | 3-04-63 | 570 | 7.9 | 6.4 | 595 | 2640 |  |  | nil |  |  | 2.4 | 880 | 620 | nil | nil | 1756 |  |
|  |  | 27 | 29-08-63 | 4-09-63 | 562 | B. 0 | 144 | 192 | 2320 |  |  | ni1 |  |  | 2.0 | 892 | 488 | nil | nil |  |  |
|  |  | 37 | 29-08-63 | 9-09-63 | 559 | 7.9 | 150 | 184 | 2280 |  |  | nil |  |  | 2.0 | 860 | 576 | nil | nil |  |  |
| 18/63 | Wami |  | 22-11-63 | 29-11-63 |  | 8.4 | 9 | 23 | 158 | nil | nil | nil |  |  | 1.2 | 304 | 1364 | nil | nil | ni1 |  |
|  |  |  | 25-11-63 | 5-12-63 |  | 8.5 | 4 | 13 | 312 | nil | nil | nil |  |  | 2.0 | 172 | 1180 | nil | nil | nil |  |
| 4/64 | Kigoboga |  |  | 4-04-64 |  | 7.5 | 58 | 75 |  |  |  |  |  |  | 2.0 | 50 | nil | nil | $n \mathrm{nil}$ | 15 |  |
|  |  |  | 2-04-64 | 9-04-64 | 151 | 7.7 | 51 | 48 | 328 |  |  | nil |  |  | 0.9 | 4.0 | 404 | nil | nil |  |  |
|  |  |  | 5-05-64 | 15-05-64 | 113 | 7.9 | 30 | 53 | 296 |  |  |  |  |  | 0.6 | 148 | 548 | nil | nil |  |  |
| 13/64 | Kigoluga |  | 7-08-64 | 19-08-64 | 225 | 8.1 | 53 | 78 | 456 |  |  | nil |  |  | 1.0 | 60 | 468 | ail | nil |  |  |
| 18/65 | Magubike |  | 4-08-65 | 27-08-65 | 489 | 8.2 | 32 | 140 | 676 |  |  | $n i l$ |  |  | 2.0 | 944 | 736 | nil | nil |  |  |
|  |  |  | 15-11-65 | 27-01-66 | 443 | 8.3 | 22 | 170 | 764 |  |  | nil |  |  | 1.4 | 2024 | 376 | nil | nil |  |  |
|  |  |  | 10-08-67 | 28-08-67 |  | 8.5 | 8 | 234 | 999 |  |  | nil |  |  | $n i 1$ | 1124 | 452 | nil | nil |  |  |
| 33/65 | Gairo |  | 19-01-66 | 2-02-68 | 220 | 8.0 | 130 | 118 | 818 |  |  | nil |  |  | ail | 528 | 236 | nil | 2.3 |  |  |
|  |  |  | 31-05-71 | 24-06-71 | 285 | 7.5 | 182 | 145 | 1052 | 0.09 |  | 0.4 | 254 | 6.5 | 0.5 | 100 | 322 | ail | 1.6 |  |  |
| 14/66 | Turiani |  |  | $30-01-67$ $30-05-69$ |  |  | 34 48 | 29 23 | 206 |  |  |  |  |  | 0.4 2.0 0.0 | 68 57 |  |  |  |  |  |
|  |  |  |  | $30-05-69$ $14-10-69$ | 86 | 8.3 8.1 | 48 43 | 23 72 |  |  |  | 0.1 | 74 |  | 2.0 0.2 | 57 60 | 299 |  |  | 51 |  |
| 41/67 | Mikese |  | 31-10-67 | 17-11-67 | 265 | 7.1 | 1184 | 427 | 220 |  |  | nil | 387 | 8 | 2.0 | 304 | 220 | nil | 1.0 |  |  |
|  |  |  | 14-11-67 | 6-12-67 | 176 | 7.4 | 18 | 104 | 510 |  |  | nil | 232 | 7 | 1.6 | 332 | 400 | nil | nil |  |  |
| 68/68 | Wami |  | 6-10-68 | 10-10-68 |  | 8.0 | 35 | 21 |  |  |  |  | 300 |  | 0.1 | 57 | 903 |  |  |  |  |
| $72 / 68$ | Kilosa |  | 4-08-71 | 11-08-71 | 150 | 8.1 | 94 | 67 | 507 | 0.9 |  | nil | 161 | 2 | 1.3 | 211 | 560 | nil | 0.4 |  |  |
| 75/68 | Kiloba |  | 13-04-72 | 17-05-72 | 63 | 7.6 | 64 | 22 | 200 | 4.2 |  | nil | 114 | 6 | 0.2 | 21 | 166 | nil | nil | 123 |  |
|  |  |  | 27-07-72 | 29-08-72 | 47 | 6.7 | 38 | 20 | 178 | 4.2 |  | 0.3 | 14 | 5 | 0.2 | 19 | 133 | $n i 1$ | nil | 73 |  |
| 78/68 | Kilosa |  | 27-07-72 | 29-08-72 | 65 | 6.7 | 67 | 22 | 258 | 4.4 |  | 0.9 | 14 | 6 | 0.4 | 142 | 187 | nil | nil | 142 |  |
| 29/69 | Mtibwa |  |  |  |  | 8.3 | 43 | 26 | 580 |  |  |  | 34 |  | 0.4 | 36 |  |  |  |  |  |
| 56/69 | Kambala |  | 3-09-69 | 14-10-69 | 324 | 8.7 | 28 | 58 | 310 | nil | nil | nil | 150 | 5 | 0.8 | 164 | 600 | nil | nil |  |  |
| 58/69 | Kambala | 70 |  | 29-11-69 | 153 | 8.4 | 54 | 28 | 250 | 1.2 | 0.1 | 0.3 | 380 | 2 | 1.0 | 259 | 598 | $n i 1$ | 0.1 | 268 |  |
| 4/70 | Mbigili |  |  |  |  | 7.9 | 23 | 18 | 130 |  |  |  | 6 | 4 | 0.5 | 42 | 175 |  |  | 25 |  |
| 46/71 | Kidugal1o | 60 | 12-08-71 | 18-08-71 | 114 | 7.8 | E16 | 30 | 295 |  |  |  | 126 | 3 | 1.0 | 142 | 496 | 4.5 |  |  |  |
|  |  | 60 | 7-07-71 | 15-07-71 | 150 | 7.1 | 103 | 30 | 381 |  |  | nil | 201 | 6 | 1.1 | 177 | 560 | nil | 2.8 |  |  |
|  |  | 60 | 21-07-71 | 29-07-71 | 150 | 7.4 | 140 | 9 | 386 | 0.4 |  | 0.4 | 115 | 13 | 1.4 | 199 | 524 | nil | 15.2 |  |  |
| 73/71 | Mtibwa |  |  |  |  | 6.9 | 56 | 19 | 810 | 0.9 |  |  | 98 | 2 | 0.4 | 142 |  |  |  | 37 |  |
| 14/72 | Heihwa |  |  |  |  | 8.5 | 29 | 19 | 443 | 1.9 |  |  | 57 | 3 |  | 61 |  |  |  | 21 |  |
| 38/72 | Magadu | $\begin{aligned} & 24 \\ & 36 \end{aligned}$ | 24-04-72 $17-06-72$ | $19-05-72$ $10-08-72$ | 650 270 | 7.4 7.4 | 169 | 152 142 | 1049 952 | 3.2 0.1 |  | nil | 936 346 | 6 5 | 0.2 0.2 | 2201 1164 | 362 121 | nil nil | ${ }_{\text {ni] }}{ }^{2}$ | 95 55 |  |


| bore- <br> hole no. | cations <br> unions | TDS | $\mathrm{H}_{2} \mathrm{~S}$ | $\mathrm{O}_{2}$ | 88 | turb | color | collform | cu | 2 n | Pb |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  |  |  | ppm |  | mg/1 | FTU | Pt/Co | count <br> 100 ml |  | ppm |  |
| 5/63 |  | 5055 4810 4549 |  |  |  |  |  |  |  |  |  |
| 18/63 |  | 2007 | , |  |  |  |  |  | nil | nil |  |
|  |  | 604 |  |  |  |  |  |  |  |  |  |
|  |  | 816 |  |  |  |  |  |  |  |  |  |
| 13/64 |  | 1232 |  |  |  |  |  |  |  |  |  |
| 18/65 |  | 2935 |  |  |  |  |  |  |  |  |  |
|  |  | 2680 |  |  |  |  |  |  |  |  |  |
| 33/65 |  | 3331 |  |  |  |  |  |  |  | 1.6 |  |
|  |  | 1800 |  |  |  |  |  |  |  |  |  |
| 14/66 |  | 484 |  |  |  |  |  |  |  |  |  |
|  |  | 468 |  |  |  |  |  |  |  |  |  |
| 41/67 |  | 6922 |  |  |  |  |  |  |  |  |  |
|  |  | 1051 |  |  |  |  |  |  |  |  |  |
| 68/68 |  | 886 |  |  |  |  |  |  |  |  |  |
| 72/68 | - | 1100 |  |  |  |  |  |  |  |  |  |
| 75/68 |  | 400 |  |  |  |  |  |  |  |  |  |
|  |  | 510 |  |  |  |  |  |  |  |  |  |
| $78 / 68$ $29 / 69$ |  | 425 330 |  |  |  |  |  |  |  |  |  |
| 29/69 $56 / 69$ |  | $\begin{array}{r}330 \\ 1564 \\ \hline\end{array}$ |  |  |  |  |  |  |  |  |  |
| 58/69 |  | 1618 |  |  |  |  |  |  | nil | $n i$ | ni1 |
| $4 / 70$ |  | 290 |  |  |  |  |  |  |  |  |  |
| 46/71 |  | 750 |  |  |  |  |  |  |  |  |  |
|  |  | 1000 |  |  |  |  |  |  |  |  |  |
|  |  | 1370 |  |  |  |  |  |  |  |  |  |
| $73 / 71$ $14 / 72$ |  | 630 |  |  |  |  |  |  |  |  |  |
| $14 / 72$ <br> $38 / 72$ |  | 310 4250 |  |  |  |  |  |  |  |  |  |
| 3672 |  | 4250 1650 |  |  |  |  |  |  |  |  |  |
|  |  |  |  |  |  |  |  |  |  |  |  |

Continued 3

| borehole no. | location | depth | date of sampling | date of analysis | spec. cond. at $25^{\circ} \mathrm{C}$ | pH | $\mathrm{Ca}^{2+}$ | $\mathrm{Hg}^{\mathbf{2 +}}$ | TH | Fe | Hn | Nil ${ }^{+}$ | $\mathrm{Ha}^{+}$ | $\mathrm{K}^{+}$ | $\mathbf{F}^{*}$ | $\mathrm{Cl}^{-}$ | HCO | NO | HO | $\mathrm{so}^{2-}$ | $\mathrm{PO}_{4}^{3-}$ |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  |  | $\cdots$ |  |  | ms/r |  | ppa |  | $\stackrel{\mathrm{Ppman}_{\mathrm{CaCO}}^{3}}{ }$ |  |  |  |  |  | part | per | illion |  |  |  |  |
| 95/72 | Kimamba | 18 | 18-08-72 | 24-08-72 |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |
|  |  |  | 14-09-72 | 27-10-72 | 150 | 8.1 | 51 | 62 | 380 | 75 | nil | nil | 115 | 1 | 0.1 | 102 | 278 | nil | ni1 | 277 |  |
|  |  | 129 | 39-11-72 | 9-12-72 |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |
|  |  | 123 | 28-11-72 | 19-12-72 | 58 | 7.8 | 35 | 17 | 157 | 0.4 |  | nil | 66 | 4 | 0.2 | 29 | 255 | ni1 | 1.7 | 27 |  |
|  |  | 100-120 | 15-12-72 | 19-12-72 | 60 | 7.7 | 35 | 18 | 159 | 0.4 | 0.6 | nil | 65 | 6 | 0.2 | 30 | 234 | nil | 1.6 |  |  |
|  |  | 125 | 27-12-72 | 16-01-73 | 58 | 7.9 | 56 | 19 | 160 | nil | 0.5 | nil | 103 | 3 | 0.3 | 26 | 236 | nil | nil | 50 |  |
| 175/77 | Msowero |  |  |  | 155 | 8.6 | 90 | 48 | 425 | 2.5 |  | nil | 215 | 11 | 1.0 | 173 | 559 | ail | 0.5 |  |  |
| 226/77 | Dumila |  |  |  | 102 | 8.7 | 14 | 8 | 68 |  | 0.7 |  | 82 | 43 | 3.5 | 125 | 396 | nil | nil |  |  |


| borehole no. | cations <br> anions | TDS | $\mathrm{H}_{2} \mathrm{~S}$ | $\mathrm{O}_{2}$ | SS | turb | color | coll form | Cu | 2 n | Pb |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  |  |  | ppm |  | mg/1 | FTU | $\mathrm{Pt} / \mathrm{Co}$ | count <br> 100 ml |  | ppm |  |
| $\left\{\begin{array}{c} 95 / 72 \\ \\ 175 / 77 \\ 27^{\prime} / 777 \end{array}\right.$ |  | $\begin{array}{r} 4680 \\ 1125 \\ 800 \\ 525 \\ 350 \\ 360 \\ 390 \\ 1040 \end{array}$ |  |  |  |  |  |  |  |  |  |

Data DD 4
Hydrogeological field data of existing ground water supplies

DD
4 - Hydrogeological field data of existing ground water supplies



| village | nos. and types of ground water supplies (Nov. 1978)* |  |  |  |  |  | type of supply | total depth (m-GL) | aquif. depth (m-GL) | aquif. thickn. ( $\mathrm{m}-\mathrm{GL}$ ) | aquif. type | water Level (m-GL) | $\begin{aligned} & \mathrm{EC} \\ & \mathrm{mS} / \mathrm{m} \end{aligned}$ | date | regular measurement programme |  |  |  | remarks |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  |  |  |  |  |  |  | min <br> WL <br> ( $\mathrm{m}_{\mathrm{B}}$-GL) |  |  |  |  |  |  |  | $\begin{aligned} & \text { max. } \\ & \text { WL } \\ & (\mathrm{m}-\mathrm{GL}) \end{aligned}$ | min. <br> E.C <br> mS/m | $\max$. <br> EC <br> mS/m |  |
|  | hh | bhrb | sw | Psw | pbh | s |  |  |  |  |  |  |  |  |  |  |  |  |
| Kwaba |  | 4 |  |  |  |  | thirb | 0.4 |  |  | coarse sand | 0.4 | 125 | 18-10-78 | 1.0 |  |  |  |  |
| Leshata |  | 6 |  |  |  |  | hhirb | 0.4 |  |  |  | 0.2 | 40 | 29-06-78 | 1.4 | 0.0 | 40 | 50 |  |
| Lubungu | 2 |  | 2 |  |  |  | sw | 4.1 |  |  | sandy loam | 3.7 | 53 | 29-09-78 |  |  |  |  | dry out |
|  |  |  |  |  |  |  | sw | 3.4 |  |  |  | 3.1 | 82 |  |  |  |  |  |  |
| Euhindo | no | ound w | ater | supp | lies |  |  |  |  |  |  |  |  |  |  |  |  |  |  |
| Lukenge | 2 |  |  |  |  |  | swrb | 1.5 |  |  |  | 1.3 | 20 | 11-08-78 |  |  |  |  | pumped supply |
|  |  |  |  |  |  |  | hh | 1.4 |  |  | basement | 1.4 | - 8 |  |  |  |  |  | out of order |
|  |  |  |  |  |  |  | hh | 2.5 |  |  | cemented sand | 2.4 | 7 |  |  |  |  |  |  |
| Lukobe | 3 |  | 1 |  |  |  | sw | 3.5 |  |  |  | 3.1 | 142 | 28-09-78 |  | 0.0 |  |  | bedrock at SM |
|  |  |  |  |  |  |  | hh | 2.5 |  |  | sand | 2.5 | 12 |  |  |  |  |  |  |
| Luhulunge | no | and w | ater | supp | lies |  |  |  |  |  |  |  |  |  |  |  |  |  |  |
| Luloengwe | 2 | 2 |  |  |  |  | hh | 3.0 | 1.5-7 |  | middle sand | 3.0 | 142 | 18-10-78 | 4.0 |  |  |  |  |
|  |  |  |  |  |  |  | ht | 2.5 |  |  | sand | 2.5 | 130 |  | 4.0 |  |  |  |  |
|  |  |  |  |  |  |  | bhrb | 2.0 |  |  | coarse sand | 2.0 | 100 |  |  |  |  |  |  |
| Mabula |  | 8 |  |  |  |  | harb | 0.3 |  |  | coarse sand | 0.3 | 160 | 7-07-78 | 0.4 | 0.0 | 85 | 150 | Kibedya river |
|  |  |  |  |  |  |  | thrb | 0.1 |  |  | coarse sand | 0.1 | 85 |  |  |  |  |  | Berega river |
| Machatu | no | ound w | ater | supp | lies |  |  |  |  |  |  |  |  |  |  |  |  |  |  |
| Madege |  | 5 |  |  |  |  | hhrb | 0.9 |  |  | coarse sand | 0.9 | 200 | 29-06-78 | 1.2 | 0.0 | 140 | 200 |  |
| Madizini. | 20 |  |  |  |  |  | hh | 4.1 |  |  |  | 3.0 | 30 | 24-08-78 | 4.1 | 2.8 | 24 | 30 |  |
| Hadoto |  |  |  |  | 1 |  | pbh |  |  |  |  |  | 75 | 25-07-78 | - |  |  |  | borehole from sisal estate |
| Madudu | 5 |  | 2 |  |  |  | sw | 7.2 |  |  |  | 3.9 | 60 | 15-08-78 | 6.0 | 3.0 | 55 | 60 | never dry |
|  |  |  |  |  |  |  | sw | 3.2 |  |  |  | 3.2 | 64 |  |  |  |  | - | ) dry in dry |
|  |  |  |  |  |  |  | hh | 3.0 |  |  |  | 2.4 | 49 |  | 3.0 | 2.0 | 39 | 60 | ) season |
| Magera |  | 10 |  |  |  |  | hhrb | 0.3 |  |  | sand + gravel | 0.3 | 80 | 7-07-78 | 1.8 | 0.0 | 60 | 110 |  |
| Magole |  |  |  |  | 1 |  | pbh |  |  |  |  |  | 72 | 16-07-78 |  |  |  |  | often out of order |
| Magomeni | 4 |  | 3 |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |
| Magubike |  | 8 |  |  | 1 |  | pbh |  |  |  |  |  | 490 | 10-10-78 |  |  |  |  | abandoned |
|  |  |  |  |  |  |  | hhrb | 2.0 |  |  |  | 2.0 | 84 |  |  |  |  |  |  |
| Maguha |  |  |  |  |  | 1 |  |  |  |  |  |  | 49 | 4-07-78 |  |  |  |  |  |
| Maharaka |  | 4 |  |  |  |  | bhirb | 1.0 |  |  | sandy loam | 1.0 | 140 | 19-10-78 |  |  |  |  |  |
|  |  |  |  |  |  |  | hincb | 1.0 |  |  | sand |  | 130 |  |  |  |  |  |  |
| Makyuy |  | 6 |  |  |  |  | hheb | 2.2 |  |  |  | 2.1 | 28 | 6-07-78 |  |  |  |  |  |
| Makuyu (Tur) | 3 |  | 2 |  |  |  | sw | 10.0 |  |  |  | 4.5 | 26 | 15-08-78 | 7.0 | 4.0 | 26 | 28 |  |
|  |  |  |  |  |  |  | sw | 4.5 |  |  | middle sand | 2.6 | 61 |  | 3.7 | 2.0 | 49 | 65 |  |
|  |  |  |  |  |  |  | hh | 11.5 |  |  | sand | 11.0 | 64 |  | 11.5 | 10.5 | 60 | 80 |  |
|  |  |  |  |  |  |  | hh | 13.0 |  |  | loam + stores | 12.9 | 102 |  | 12.9 | 12.2 | 100 | 115 |  |
| Malangalo | 10 |  |  |  |  |  | hh | 3.0 |  |  | coarse sand | 1.6 | 68 | 12-10-78 |  |  |  |  |  |
|  |  |  |  |  |  |  | hh | 2.8 |  |  |  | 2.4 | 84 |  | 5.0 | 2.0 | 70 | 90 |  |




hh $=$ hand dug hole
hhrb $=$ hand dug hole in river bed
$\mathrm{sw}=$ shallow well
psw $=\quad$ pumped supply from shallow well
ph $=$ pumped supply from borehole
$s=$ spring

## DATA DD 5

## SUMMARY OF MWCP-HAND DRILLED HOLES APPROVED FOR CONSTRUCTION OF SHALLOW WELLS

DD 5 - Sumary of MWCP hand drilled holes, approved for construction of shallow wells


| village | hand <br> drilling <br> no. | total depth (m-GL) | aquifer depth <br> from - to ( $\mathrm{m}-\mathrm{GL}$ ) | aquifer thickness (m) | aquifer type | water level during drilling ( $\mathrm{m}-\mathrm{GL}$ ) | tested yield ( $1 / \mathrm{FR}$ ) | maximum drawdown | $\underset{\left(\mathrm{m}^{2} / \text { day }\right)}{ }$ | $\begin{aligned} & \mathrm{X} \\ & (\mathrm{~m} / \text { day }) \end{aligned}$ | EC during remarks drilling <br> (ms/m) |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| Dihombo | 166/3-7-1 | 8.5 | 5.0-8.0 | 3.0 | fine sand | 1.3 | 1200 | 0.4 | 90 | 30 | 30 |
|  | 166/3-6-1 | 8.0 | 4.2-8.0 | 3.8 | middle sand | 1.7 | 1200 | 2.8 | 13 | 3 | 20 |
|  | 166/3-5-1 | 9.0 | 5.0-9.0 | 4.0 | fine sand | 2.1 | 1400 | 0.4 | 105 | 26 | 20 |
| $\begin{aligned} & \text { Bumila } \\ & \text { Hembet } \mathrm{i} \end{aligned}$ | 165/4-8-2 | 7.0 | 4.0-6.0 | 2.0 | middle sand | 3.8 | 1400 | 0.3 | 140 | 70 | 95 |
|  | 165/4-17-1 | 9.2 | 8.0-9.2 | 1.2 | coarse sand | 5.8 | 1200 | 0.4 | 90 | 75 | 15 |
|  | $166 / 3-4-1$ | 9.0 | 2.0-2.5 | 0.5 | coarse sand | 1.4 | 1380 | 0.3 | 138 | $276$ | 60 |
|  |  |  | 6.0-9.0 | 3.0 | coarse sand |  |  |  |  |  |  |
|  | 166/3-3-1 | 7.0 | 2.5-6.8 | 4.3 | fine sand | 2.0 | 1400 | 0.7 | 60 | 14 | 70 |
|  | $166 / 3-1-1$ | 7.5 | 3.5-7.5 | 4.0 | middle sand | 1.9 | 1200 | 0.6 | 60 | 15 | 20 |
| Kibungo | 183/4-12-3 | 3.5 | 0.8-3.5 | 2.7 | middle sand | 0.8 | 600 | 0.7 | 26 | 10 | 45 |
|  | 183/4-11-1 | 5.2 | 1.0-4.8 | 3.8 | coarse sand | 0.8 | 1000 | 0.7 | 43 | 11 | 38 |
| Kikundi | 183/4-10-7 | 4.7 | 1.0-3.5 | 2.5 | middle sand | 0.3 | 600 | 3.3 | 5 | 2 | 170 |
|  | 183/4-9-5 | 5.5 | 2.0-5.5 | 3.4 | sand | 0.3 | 1300 | 0.3 | 130 | 37 | 76 |
|  | 184/1-4-1 | 6.8 | 5.2-6.8 | 1.6 | fine sand | 1.7 | 1260 |  |  |  | 150 |
| Kiroka | 183/4-3-2 | 6.5 | 2.0-4.5 | 2.5 | clay sand | 1.0 | 1000 | 0.1 | 300 | 120 | 68 |
|  | 183/4-2-3 | 3.0 | 1.4-3.0 | 1.6 | sand | 0.5 | 500 | 2.5 | 6 | 4 | 70 |
|  | 183/4-1-1 | 3.2 | 2.5-3.2 | 0.7 | clay/stones | 0.7 | 1000 | 0.3 | 100 | 143 | 68 |
| Kiziwa | 183/8-3 | 3.3 | 2.0-3.3 | 1.3 | middle sand | 1.9 | 1000 | 0.9 | 33 | 25 | 58 |
|  | 183/4-6-1 | 4.2 | 2.0-4.2 | 2.2 | middle sand | 1.0 | 1200 | 0.2 | 180 | 82 | 88 |
|  | 183/4-5-3 | 5.0 | 0.5-1.5 | 1.0 | fine sand | 0.6 | 700 | 3.4 | 6 | 6 | 74 |
|  |  |  | 3.0-4.5 | 1.5 | fine sand |  |  |  |  |  |  |
| Kondoa | $182 / 3-13$ | 10.0 | 2.3-7.2 | 4.9 | fine sand | 2.43 | $1505$ |  |  |  | $50$ |
|  | 182/3-14 | 10.0 | 1.8-5.7 | 3.9 | middle sand | 1.93 | $1224$ |  |  |  | $58$ |
|  |  |  | 6.4-7.0 | 0.6 | fine sand |  |  |  |  |  |  |
|  |  |  | 7.9-8.8 | 0.9 | fine sand |  |  |  |  |  |  |
|  |  |  | 9.8-10.0 | 0.2 | fine sand |  |  |  |  |  |  |
|  | 182/3-15 | 10.0 | 1.2-1.5 | 0.3 | fine sand | 1.37 | 1479 |  |  |  | 55 |
|  |  |  | 3.0-4.9 | 1.9 | fine sand |  |  |  |  |  |  |
|  |  |  | 7.7-9.0 | 1.3 | fine sand |  |  |  |  | . |  |
|  | 182/3-16 | 10.0 | 2.0-2.3 | 0.3 | fine sand | 2.16 | 1666 |  |  |  | 37 |
|  |  |  | 2.5-10.0 | 7.5 | fine sand |  |  |  |  |  |  |
|  | 182/3-17 | 10.0 | 3.0-3.8 | 0.8 | fine sand | 2.20 | 1241 |  |  |  | 38 |
|  |  |  | 4.3-9.0 | 4.7 | fine sand |  |  |  |  |  |  |
|  | $182 / 3-19$ | 11.5 | 2.0-7.0 | 5.0 | fine sand | 1.31 | 1139 |  |  |  | 57 |
|  | 182/3-20 | 11.0 | 0.5-5.8 | 5.3 | fine sand | 0.50 | 1632 |  |  |  | 100 |
| Konga | 183/3-11-1 | 6.7 | 3.5-6.5 | 3.0 | loan/stones | 3.5 | 500 | 2.5 | 6 | 2 | 50 |
|  | 183/3-9-1 | 9.0 | 5.5-6.5 | 1.0 | fine sand | 4.8 | 1000 |  |  |  | 15 |
| Kwamtonga | 166/1-17-1 | 9.0 | 5.7-5.8 | 0.1 | fine sand | 1.9 | 1000 | 3.6 | 8 | 80 | 28 |
|  | 166/1-15-1 | 8.5 | 3.5-5.5 | 2.0 | loamy sand | 0.8 |  |  |  |  |  |
| Madoto | 182/1-2-1 | 11.0 | 9.3-9.7 | 0.4 | clay sand | 3.05 | 629 |  |  |  | 185 |
|  |  |  | 9.8-9.9 | 0.1 | fine sand |  |  |  |  |  |  |
|  | 182/1-5-1a | 9.6 | $3.6-4.2$ | $0.6$ | fine sand | 1.80 | 1020 |  |  |  | 70 |
|  |  |  | 9.4-9.6 | 0.2 | fine sand |  |  |  |  |  |  |


| village | hand drilling no. | total depth (a-GL) | aquifer depth from - to ( $\quad$-GL) | aquifer thickness <br> (m) | aquifer type | water level <br> during drilling $(m-G L)$ | tested yield (1/HR) | maximum drawdown | ${ }_{\left(\mathrm{m}^{2} / \text { day }\right)}^{\mathrm{KD}}$ | $\begin{aligned} & \mathrm{K} \\ & (\mathrm{~m} / \text { day }) \end{aligned}$ | EC during remarks drilling (mS/m) |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  | 182/1-5-1b | 8.8 | 4.5-5.8 | 1.3 | middle sand | 2.90 | 1020 |  |  |  | 52 |
|  |  |  | 8.2-8.8 | 0.6 | middle sand |  |  |  |  |  |  |
| Magubike | 165/1-16-4 | 4.6 | 1.0-1.3 | 0.3 | sand | 0.5 | 1440 | 0.4 | 108 | 600 | 88 |
|  | 165/1-17-3 | 7.0 | 5.0-7.0 | 2.0 | coarse sand | 2.0 | 1275 | 0.2 | 191 | 96 | 175 |
| Maharaka | 200/2-8-2 | 6.5 | 3.5-6.2 | 2.7 | middle sand | 1.6 | 1200 | 0.4 | 90 | 33 | 100 |
|  | 200/2-7-2 | 4.8 | 0.5-2.2 | 1.7 | fine sand | 0.5 | 1200 | 1.0 | 36 | 21 | 155 |
|  |  |  | 3.5-4.8 | 1.3 | weathered basement |  |  |  |  |  |  |
|  | 200/2-6-1 | 9.0 | 4.5-5.2 | 0.7 | fine sand | 2.9 | 1000 | 0.1 | 300 | 429 | 150 |
|  |  |  | 6.5-8.2 | 1.7 | middle sand |  |  |  |  |  |  |
| Makuyu | 165/1-24-3 | 7.0 | 4.0-6.3 | 2.3 | fine sand | 2.7 | 1530 | 0.2 | 230 | 100 | 35 |
|  | 165/1-25-2 | 8.8 | 5.0-7.0 | 2.0 | fine sand | 4.3 | 1530 | 0.35 | 131 | 66 | 54 |
|  | 165/1-26-2 | 7.5 | 6.0-7.0 | 1.0 | aiddle sand | 1.3 | 1200 |  |  |  | 120 |
|  | 165/1-21-4 | 6.5 | 2.5-6.0 | 3.5 | fine sand | 1.7 | 990 | 3.8 | 8 | 2 | 77 |
|  | 165/4-34-1 | 9.0 | 0.8-2.0 | 1.2 | sand | 0.4 | 1150 | 0.4 | 86 | 72 | 110 |
| Mandela | 165/4-47-2 | 9.0 | 8.2-9.0 | 0.8 | coarse sand | 5.70 | 950 | 1.7 | 17 | 21 | 130 |
|  | 165/4-45-1 | 8.2 | 6.5-8.2 | 1.7 | middle sand | 6.5 | 990 |  |  |  | 55 |
| Manza | 201/1-3-2 | 9.0 | 2.0-6.5 | 4.0 | middle sand | 1.8 | 1280 | 0.3 | 128 | 32 | 70 |
|  | 201/1-2-1 | 5.8 | 4.5-5.8 | 1.3 | stones | 1.8 | 820 | 0.8 | 31 | 24 | 130 |
|  | 201/1-1-2 | 5.2 | 3.0-4.5 | 1.5 | weathered basement | 0.2 | 1170 | 1.0 | 35 | 23 | 130 |
| Mbogo | 166/1-14-1 | 9.0 | 5.2-8.5 | 3.2 | middle sand | 1.6 |  |  |  |  | 10 |
|  | 166/1-13-1 | 9.0 | 2.0-3.0 | 1.0 | coarse sand | 2.0 | 1200 | 0.4 | 90 | 90 | 19 |
|  |  |  | 5.7-6.0 | 0.3 | fine sand |  |  |  |  |  |  |
| Mikese | 183/4-15-1 | 7.0 | 1.8-6.0 | 3.2 | fine sand | 1.1 | 900 | 4.2 | 6 | 2 | 120 |
|  | 183/4-16-1 | 6.4 | 2.2-6.4 | 4.2 | fine sand | 0.4 | 1000 |  |  |  | 61 |
|  | 183/4-20-1 | 6.0 | 4.2-4.5 | 0.3 | middle sand | 0.8 | 1100 |  |  |  | 140 |
|  |  |  | 5.8-6.0 | 0.2 | middle sand |  |  |  |  |  |  |
|  |  | 7.0 | 2.0-3.2 | 1.2 | fine sand | 0.3 | 1140 | 0.5 |  |  | 89 |
| Hirama | 165/4-43-1 | 7.3 | 1.0-7.0 | 6.0 | middle sand | 0.9 | 1430 | 1.2 | 36 | 6 | 50 |
|  | 165/4-44-3 | 10.0 | 2.7-5.8 | 3.1 | middle sand | 2.4 | 1300 | 0.1 | 390 | 126 | 20 |
|  | 165/4-40-3 | 11.0 | $0.8-1.2$ | 0.3 | coarse sand | 0.7 | 900 |  |  |  | 150 |
|  |  |  | 8.2-9.1 | 0.9 | middle sand |  |  |  |  |  |  |
|  |  |  | 10.2-11.0 | 0.8 | middle sand |  |  |  |  |  |  |
| Mkindo | 166/1-4-1 | 9.0 | 5.0-5.8 | 0.8 | fine sand | 3.7 | 1240 | 0.7 | 53 | 66 | 35 |
|  |  |  | 6.5-8.0 | 1.5 | middle sand |  |  |  |  |  |  |
|  | 166/1-3-1 | 9.0 | 4.0-5.5 | 1.5 | fine sand | 4.1 | 1200 | 0.8 | 45 | 30 | 15 |
|  |  |  | 6.5-9.0 | 2.5 | loamy sand |  |  |  |  |  |  |
|  | 166/1-2-1 | 9.0 | 6.0-9.0 | 3.0 | coarse sand | 5.4 | 1200 | 0.7 | 51 | 17 | 75 |
|  | 166/1-1-1 | 8.0 | 5.0-6.5 | 1.5 | voarse sand | 3.6 | 600 | 2.0 | 6 | 4 | 120 |
|  | 166/3-91-1 | 8.5 | $3.5-5.5$ | $2.0$ | loamy sand | 2.0 | 1100 | 2.0 | 11 | 6 | 20 |
|  |  |  | 6.5-8.5 | 2.0 | coarse sand |  |  |  |  |  |  |


| village | hand drilling no. | total depth ( $\mathrm{m}-\mathrm{GL}$ ) | aquifer depth <br> from - to ( $\mathrm{m}-\mathrm{GL}$ ) | aquifer thickness (m) | aquifer type | water level <br> during drilling $(\mathrm{m}-\mathrm{GL})$ | tested yield (1/HR) | maximum drawdown | $\begin{aligned} & \mathrm{KD} \\ & \left(\mathrm{~m}^{2} / \text { day }\right) \end{aligned}$ | $\begin{aligned} & \mathrm{K} \\ & (\mathrm{~m} / \text { day }) \end{aligned}$ | EC during remarks drilling ( $\mathrm{mS} / \mathrm{m}$ ) |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| Msongozi. | 200/2-4-1 | 9.5 | 6.0-9.0 | 3.0 | middle sand | 3.6 | 1000 | 0.6 | 50 | 17 | 175 |
|  | 200/2-3-1 | 7.0 | 2.0-7.0 | 5.0 | loam/stones | 1.9 | 1300 | 0.1 | 390 | 78 | 150 |
|  | 200/2-1-3 | 7.0 | 5.0-7.0 | 2.0 | coarse sand | 1.2 | 1300 | 0.1 | 390 | 195 | 180 |
| Msufini | 165/4-16-1 | 7.0 | 0.8-1.3 | 0.5 | fine sand | 0.8 | 1480 | 4.5 | 10 | 20 | 70 |
|  |  |  | 4.0-5.2 | 1.2 | fine sand |  |  |  |  |  |  |
|  | 165/4-15-1 | 7.5 |  | 0.5 | middle sand | 1.8 | 1400 | 0.4 | 105 | 210 | 22 |
|  |  |  | 4.2-4.8 | 0.6 | loam |  |  |  |  |  |  |
|  | 165/4-15-1 | 5.0 | 2.3-4.0 | 1.7 | fine sand | 2.0 | 1100 | 2.5 | 13 | 8 | 62 |
| Mugudeni | 165/4-7-1 | 8.2 | 5.8-8.0 | 2.2 | coarse sand | 5.3 | 1300 | 0.2 | 195 | 89 | 80 |
|  | 165/4-6-1 | 7.5 | 3.5-6.5 | 3.0 | fine sand | 1.1 | 1240 | 3.0 | 12 | 4 | 175 |
|  | 165/4-3-1 | 8.0 | 5.5-6.5 | 1.0 | fine sand | 2.5 | 1200 | 0.2 | 180 | 180 | 115 |
|  | 165/4-2-1 | 9.0 | 3.5-4.5 | 1.0 | fine sand | 1.4 | 1100 | 0.5 | 66 | 66 | 150 |
|  |  |  | 5.0-8.0 | 3.0 | fine sand |  |  |  |  |  |  |
|  | 165/4-2-2 | 6.5 | 3.5-5.5 | 2.0 | fine sand | 2.4 | 1100 | 0.2 | 165 | 83 | 150 |
| Mvomero | 165/4-31-1 | 9.0 | 5.5-8.6 | 2.6 | middle sand | 4.7 | 950 | 1.5 | 19 | 7 | 120 |
|  | 165/4-25-1 | 9.5 | 6.8-9.5 | 2.7 | middle sand | 7.1 | 1200 |  |  |  | 35 |
|  | 165/4-26-1 | 9.0 | 6.5-9.0 | 2.5 | fine sand | 1.9 | 1200 |  |  |  | 95 |
|  | 165/4-24-1 | 9.5 | $8.0-9.5$ | 1.5 | loam | 6.9 | 650 |  |  |  |  |
| Ngerengere <br> Darajani | 184/1-2-2 | 6.1 | 4.0-6.1 | 2.1 | fine sand | 3.1 | 1330 |  |  |  | 170 |
|  |  |  |  |  |  |  |  |  |  |  |  |
|  | 182/1-7 | 15.0 | 9.8-11.2 | 1.4 | clay sand | 4.3 | 1071 |  |  |  | 42 |
| Batini | 182/1-8-1 | 10.0 | 3.7-4.1 | 0.4 | fine sand | 5.0 | 646 |  | - |  | 45 |
|  | 182/1-9 | 10.0 | 4.9-6.2 | 1.3 | fine sand | 5.0 | 459 |  |  |  | 70 |
|  | 182/1-10 | 10.0 | 4.8-5.0 | 0.2 | coarse sand | 5.3 | 1224 |  |  |  | 35 |
|  | 182/1-11- | 10.0 | 5.0-6.0 | 1.0 | loam sand | 4.48 | 1462 |  |  |  | 24 |
|  | 182/1-13-2 | 11.0 | 6.8-9.5 | 2.7 | fine sand | 7.25 | 1309 |  |  |  | 14 |
|  | 182/1-14 | 10.0 | 5.7-9.8 | 4.1 | coarse sand | 5.6 | 1224 |  |  |  | 32 |
|  | 182/1-15-2 | 10.0 | 6.2-7.5 | 1.3 | fine sand | 4.8 | 1207 |  |  |  | 36 |
|  | 182/1-16-1 | 10.0 | 4.1-5.0 | 0.9 | fine sand | 4.0 | 1275 |  |  |  | 22 |
|  |  |  | 6.0-6.6 | 0.6 | fine sand |  |  |  |  |  |  |
|  | 182/1-17 | 9.2 |  | 1.1 |  | 5.0 | 1241 |  |  |  | 30 |
|  |  |  | 7.2-9.0 | 1.8 | fine sand |  |  |  |  |  |  |
|  | 182/1-18 | 12.0 | 5.0- 5.7 | 0.7 | coarse sand | 5.0 | 1530 |  |  |  | 64 |
|  | 182/1-20 | 11.0 | 6.5-7.2 | 0.7 | middle sand | 5.68 | 1000 |  |  |  | 44 |
|  | 182/1-37 | 12.0 | 9.8-10.8 | 1.0 | fine sand | 5.86 | 540 |  |  |  | 32 |
| Rudewa | 182/1-31-2 | 8.0 | 0.6-3.2 | 2.6 | middle sand | 0.4 | 1173 |  |  |  | 20 |
| Congoni | 182/1-32-1 | 10.0 | 3.5-5.3 | 1.8 | middle sand | 1.55 | 561 |  |  |  | 87 |
| Rudewa | 182/1-12 | 10.0 | 5.1-5.8 | 0.7 | middle sand | 5.3 | 826 |  |  |  | 24 |
| Mbuyuni |  |  | 9.4-9.7 | 0.3 | fine sand |  |  |  |  |  |  |
|  |  |  | 6.0-6.5 | 0.5 | fine sald | 4.7 | 1676 |  |  |  | 28 |
|  | 182/1-21-1 | 11.5 | 4.0-4.3 | 0.3 | coarse sand | 4.0 | 1207 |  |  |  | 25 |
|  |  |  | 6.0- 7.4 | 1.3 | fine sand |  |  |  |  |  |  |
|  |  |  | 7.6-8.4 | 0.8 | fine sand |  |  |  |  |  |  |
|  | 182/1-22-1 | 9.1 | 4.5-5.5 | 1.0 | middle sand | 4.57 | 1530 |  |  |  | 26 |


| village | $\begin{aligned} & \text { hand } \\ & \text { drilling } \end{aligned}$ no. | $\begin{aligned} & \text { total } \\ & \text { depth } \\ & (\mathbf{m}-G L) \end{aligned}$ | aquifer depth <br> from - to <br> ( m -GL) | aquifer thickness (m) | aquifer type | water level during drilling ( m -GL) | $\begin{aligned} & \text { tested } \\ & \text { yield } \\ & (1 / \mathrm{HR}) \end{aligned}$ | maximum <br> drawdown | $\mathrm{KD}_{\left(\mathrm{m}^{2} / \mathrm{day}\right)}$ | $\underset{(\Phi / \text { day })}{\mathrm{K}}$ | EC during remarks drilling <br> ( $\mathrm{mS} / \mathrm{m}$ ) |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  |  |  | 5.9-9.10 | 3.2 | fine sand |  |  |  |  |  |  |
|  | 182/1-23 | 10.0 | 3.9-4.5 | 0.6 | middle sand | 4.1 | 1003 |  |  |  | 30 |
|  | 182/1-25-1 | 13.0 | 11.8-13.0 | 1.2 | fine sand | 5.0 | 1105 |  |  |  | 130 |
|  | 182/1-26-1 | 11.2 | 10.5-11.0 | 0.5 | middle sand | 3.5 | 969 |  |  |  | 74 |
|  | 182/1-27-1 | 11.0 | 4.0-6.2 | 2.2 | fine sand | 3.75 | 1088 |  |  |  | 110 |
|  |  |  | 10.1-11.0 | 0.9 | fine sand |  |  |  |  |  |  |
|  | 182/1-28-1 | 13.5 | 13.2-13.5 | 0.3 | fine sand | 3.7 | 935 |  |  |  | 62 |
|  | 182/1-29-2 | 10.0 | 9.0-9.7 | 0.7 | fine sand | 3.15 | 986 |  |  |  | 57 |
|  | 182/1-30-1 | 8.2 | 3.0-3.2 | 0.2 | conrse sand | 2.9 | 1173 |  |  |  | 31 |
|  |  |  | 5.1-8.0 | 2.9 | middle sand |  |  |  |  |  |  |

## DATA DD 6

data on geo-electrical soundings

Explanation of letters used in following tables:
a) The soundings have been successively numbered in chronological order per 1:50,000 topographical mapsheets.
b) Some location names are only represented on the 1:50,000 mapsheet. SE = sisal estate
c) The midpoint of the soundings are represented in kilometer coordinates in accordance to the $1: 50,000$ mapsheets.

For the investigated area these co-ordinates can be converted into length and latitude degrees with the following equations:
length: $X_{\text {degree }}=9,05 \times 10^{-3} \mathrm{x}_{\mathrm{km}}+34,5$
latitude: $Y_{\text {degree }}=-9,03 \times 10^{-3} Y_{k m l}+90,3$
d) Elevation in metres above sea level.
e) The azimuth of the soundings is given in the $360^{\circ}$ system.

| mapsheet | number <br> a) | $\begin{aligned} & \text { date } \\ & 1978 \end{aligned}$ | location <br> b) | co-ordinates <br> c) | eleva- <br> tion <br> d) | $\mathrm{L} / 2$ <br> $\max$ | azimuth <br> e) |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| 165/3 | 1 | 14-7 | Mbugani | 306,0-9281,8 | 230 | 200 | 45 |
| 165/4 | 1 | 31-5 | - | 334,3-9390,7 | 375 | 175 | 130 |
|  | 2 | 31-5 | - | 329,9-9393,8 | 385 | 125 | 135 |
|  | 3 | 1-6 | - | 308,4-9391,3 | 455 | 200 | 70 |
|  | 4 | 1-6 | Madudu | 311,8-9392,3 | 455 | 200 | 85 |
|  | 5 | 1-6 | Dumila | 316,4-9394,2 | 430 | 200 | 60 |
|  | 6 | 1-6 | Magole | 319,3-9395,0 | 425 | 200 | 80 |
|  | 7 | 2-6 | Mandela | 321,9-9397,1 | 430 | 200 | 40 |
|  | 8 | 2-6 | - | 323,6-9398,8 | 430 | 200 | 60 |
|  | 9 | 2-6 | Mugudeni | 326,1-9300,6 | 420 | 200 | 45 |
|  | 10 | 5-6 | Mvomero | 327,6-9302,3 | 410 | 300 | 170 |
|  | 11 | 5-6 | - | 327,2-9303,4 | 415 | 200 | 90 |
|  | 12 | 5-6 | Mirama | 328,2-9395,7 | 395 | 200 | 135 |
|  | 13 | 6-6 | Mugudeni | 325,9-9299,7 | 420 | 150 | 65 |
|  | 14 | 6-6 | - | 326,4-9299,1 | 410 | 150 | 160 |
|  | 15 | 6-6 | - | 327, 0-9297,2 | 400 | 175 | 155 |
|  | 16 | 6-6 | - | 331,5-9292,2 | 380 | 150 | 130 |
|  | 17 | 6-6 | - | 327,0-9301,5 | 410 | 175 | 50 |
|  | 18 | 12-6 | - | 326,0-9303,4 | 420 | 175 | 90 |
|  | 19 | 12-6 | - | 324,0-9302,9 | 440 | 175 | 70 |
|  | 20 | 13-6 | - | 329,2-9303,9 | 410 | 200 | 60 |
|  | 21 | 13-6 | Kigugu | 330,8-9304,8 | 390 | 100 | 60 |
|  | 22 | 13-6 | Msufini | 332,1-9305,7 | 390 | 300 | 40 |
|  | 23 | 13-6 | - | 333,5-9306,5 | 390 | 100 | 100 |
|  | 24 | 23-6 | - | 328,6-9302,8 | 400 | 200 | 80 |
|  | 25 | 26-6 | - | 329,4-9302,9 | 400 | 500 | 100 |
|  | 26 | 26-6 | Dibamba | 330,9-9303,0 | 390 | 600 | 60 |
|  | 27 | 27-6 | - | 308,6-9291,6 | 455 | 150 | 80 |
|  | 28 | 27-6 | - | 309,7-9291,8 | 450 | 250 | 55 |
|  | 29 | 27-6 | - | 310,6-9292,1 | 450 | 250 | 75 |
|  | 30 | 27-6 | Madudu | 313,1-9292,6 | 440 | 175 | 75 |
|  | 31 | 28-6 | - | 315,5-9293,8 | 435 | 250 | 75 |
|  | 32 | 28-6 | - | 314,3-9293,3 | 440 | 175 | 75 |
|  | 33 | 28-6 | - | 317,2-9294,6 | 425 | 200 | 70 |
|  | 34 | 28-6 | Magole | 319,5-9254,6 | 420 | 175 | 135 |
|  | 35 | 28-6 | - | 320,9-9296,1 | 430 | 200 | 45 |
|  | 36 | 29-6 | - | 326,7-9298,2 | 405 | 175 | 160 |
|  | 37 | 29-6 | - | 332,6-9291,3 | 375 | 150 | 130 |
|  | 38 | 29-6 | - | 333,9-9290,3 | 375 | 125 | 115 |
|  | 39 | 29-6 | - | 329,3-9294,4 | 390 | 125 | 135 |
|  | 40 | 29-6 | - | 328,7-9295,1 | 390 | 150 | 135 |


| mapsheet | number <br> a) | date 1978 | location <br> b) | co-ordinates <br> c) | elevation <br> d) | L/2 <br> max | azimuth <br> e) |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| 165/4 | 41 | 29-6 | Mandela | 322,9-9298,3 | 430 | 150 | 50 |
|  | 42 | 30-6 | - | 332,1-9303.6 | 380 | 500 | 65 |
|  | 43 | 5-7 | Mbugani SE | 306,7-9283,4 | 425 | 300 | 20 |
|  | 44 | 5-7 | Kitete SE | 307,4-9284,9 | 425 | 250 | 20 |
|  | 45 | 6-7 | Kitete SE | 307,8-9286,4 | 430 | 300 | 10 |
|  | 46 | 6-7 | Kitete SE | 308, 1-9288,2 | 435 | 250 | 0 |
|  | 47 | 6-7 | Kitete | 307,9-9290,4 | 435 | 250 | 150 |
|  | 48 | 14-7 | Kitete SE | 308,0-9289,6 | 430 | 350 | 175 |
|  | 49 | 7-8 | - | 324,2-9282,4 | 370 | 100 | 90 |
|  | 50 | 7-8 | - | 324,4-9286,4 | 385 | 200 | 45 |
|  | 51 | 7-8 | Magole farms | 324,2-9288,3 | 390 | 200 | 50 |
|  | 52 | 14-8 | Magole farms | 323,6-9295,0 | 415 | 200 | 40 |
|  | 53 | 14-8 | Magole farms | 323,6-9296,2 | 420 | 200 | 60 |
|  | 54 | 14-8 | Mandela | 321,3-9297,7 | 430 | 400 | 65 |
|  | 55 | 15-8 | Mandela | 321,2-9298,3 | 440 | 500 | 0 |
|  | 56 | 15-8 | Mandela | 320,8-9297,9 | 440 | 500 | 170 |
|  | 57 | 15-8 | Mandela | 322,1-9296,7 | 430 | 175 | 150 |
|  | 58 | 18-8 | Dibamba | $329-9302$ | 400 | 300 | 90 |
|  | 59 | 18-8 | Dibamba | 330,1-9302,8 | 400 | 600 | 90 |
|  | 60 | 18-8 | Dibamba | $331-9302$ | 395 | 300 | 80 |
|  | 61 | 21-8 | Mkundi | $315-9299$ | 470 | 200 | 0 |
|  | 62 | 21-8 | Mkundi | $315-9298$ | 460 | 200 | 170 |
|  | 63 | 21-8 | Mkundi | 315-9298 | 470 | 250 | 90 |
|  | 64 | 21-8 | Mkundi | $316-9297$ | 460 | 175 | 170 |
|  | 65 | 21-8 | Mkundi | $316-9298$ | 460 | 150 | 90 |
|  | 66 | 22-8 | Madudu | 311,6-9292,1 | 440 | 250 | 85 |
|  | 67 | 22-8 | - | 311,8-9290,7 | 430 | 250 | 85 |
|  | 68 | 22-8 | Peyapeya SE | 311,9-9289,3 | 430 | 200 | 85 |
|  | 69 | 22-8 | Peyapeya SE | 311,9-9287,8 | 425 | 175 | 85 |
|  | 70 | 23-8 | - | 310,9-9293,2 | 460 | 300 | 90 |
|  | 71 | 23-8 | - | 310,6-9293,9 | 480 | 300 | 90 |
|  | 72 | 23-8 | - | 312,8-9293,5 | 455 | 250 | 90 |
|  | 73 | 23-8 | - | 311,6-9293,4 | 455 | 300 | 165 |
|  | 74 | 23-8 | - | 313,1-9286,3 | 420 | 250 | 100 |
|  | 75 | 19-9 | Mbigili SE | 322,7-9289,1 | 395 | 150 | 50 |
|  | 76 | 19-9 | Mbigili SE | 322,0-9290,3 | 400 | 200 | 50 |
|  | 77 | 19-9 | Mbigili SE | 321,1-9291,6 | 405 | 150 | 55 |
|  | 78 | 19-9 | Mbigili SE | 320,3-9292,3 | 410 | 175 | 55 |
|  | 79 | 19-9 | Mbigili SE | 319,8-9289,7 | 405 | 200 | 140 |
|  | 80 | 20-9 | Mbigili SE | 320,7-9287,2 | 390 | 200 | 145 |
|  | 81 | 20-9 | - | 326,6-9287,0 | 380 | 100 | 20 |
|  | 82 | 20-9 | - | 329,3-9286,2 | 370 | 75 | 145 |


| mapsheet | number <br> a) | $\begin{aligned} & \text { date } \\ & 1978 \end{aligned}$ | location <br> b) | co-ordinates <br> c) | elevation <br> d) | $\mathrm{L} / 2$ <br> $\max$ | azimuth <br> e) |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| 165/4 | 83 | 20-9 | Wami Prison | 330,2-9289,7 | 380 | 125 | 45 |
|  | 84 | 3-10 | Wami Prison | 330,7-9292,8 | 385 | 125 | 140 |
|  | 85 | 9-10 | - | 320,7-9299,7 | 440 | 250 | 0 |
|  | 86 | 9-10 | Magole farms | 323,8-9293,6 | 405 | 200 | 55 |
|  | 87 | 9-10 | Magole farms | 323,9-9292,0 | 400 | 200 | 50 |
|  | 88 | 9-10 | Magole farms | 324,1-9290,3 | 390 | 100 | 50 |
|  | 89 | 10-10 | - | 327,7-9295,8 | 395 | 125 | 55 |
|  | 90 | 10-10 | - | 326,1-9296,3 | 405 | 175 | 55 |
|  | 91 | 10-10 | Mirama | 327,9-9296,2 | 395 | 200 | 135 |
|  | 92 | 10-10 | - | 328,8-9295,5 | 390 | 150 | 40 |
|  | 93 | 18-10 | Makuyu | 320,7-9301,5 | 450 | 75 | 70 |
|  | 94 | 18-10 | Makuyu | 320,7-9300,9 | 440 | 350 | 90 |
|  | 95 | 18-10 |  | 320,7-9299,9 | 440 | 300 | 0 |
|  | 96 | 7-11 | - | 324,9-9302,7 | 420 | 300 | 60 |
|  | 97 | 7-11 | - | 324,7-9201,5 | 420 | 250 | 60 |
|  | 98 | 7-11 | - | 327,7-9297,5 | 400 | 150 | 160 |
|  | 99 | 8-11 | - | 329,3-9296,6 | 390 | 175 | 140 |
|  | 100 | 8-11 | - | 328,7-9297,0 | 395 | 290 | 140 |
|  | 101 | 8-11 | - | $325-9300$ | 420 | 400 | 55 |
|  | 102 | 14-11 | - | 324, 9-9302, 3 | 420 | 500 | 55 |
|  | 103 | 14-11 | Mirama | 329,4-9295,9 | 390 | 250 | 135 |
| 166/1 | 1 | 28-8 | Mtibwa | 351,8-9319,8 | 370 | 150 | 55 |
|  | 2 | 28-8 | Mtibwa | 351,8-9320,7 | 370 | 150 | 60 |
|  | 3 | 29-8 | Mtibwa | 351,2-9321,5 | 380 | 200 | 35 |
|  | 4 | 29-8 | Mtibwa | 352,7-9319,4 | 370 | 125 | 55 |
|  | 5 | 29-8 | - | 356-9316 | 360 | 100 | 175 |
|  | 6 | 29-8 | Lukenge | 348,8-9309,8 | 360 | 100 | 45 |
|  | 7 | 29-8 | Lukenge | 348,0-9310,8 | 360 | 100 | 5 |
|  | 8 | 29-8 | Lukenge | 349,1-9313,0 | 360 | 150 | 90 |
|  | 9 | 30-8 | - | 351,2-9312,8 | 360 | 300 | 90 |
|  | 10 | 30-8 | - | 354,9-9313,3 | 350 | 150 | 90 |
|  | 11 | 30-8 | - | 349,1-9313,6 | 360 | 100 | 95 |
|  | 12 | 30-8 | - | 349,3-9315,2 | 360 | 125 | 95 |
|  | 13 | 30-8 | Kidudwe | 354,2-9316,3 | 360 | 100 | 90 |
|  | 14 | 30-8 | Mtibwa | 439,6-9319,4 | 370 | 200 | 145 |
|  | 15 | 31-8 | Kigugu | 343,0-9310,6 | 355 | 200 | 40 |
|  | 16 | 31-8 | Kigugu | 343,6-9311,5 | 355 | 150 | 0 |
|  | 17 | 31-8 | Kigugu | 342,2-9311,5 | 360 | 175 | 20 |
|  | 18 | 31-8 | Kigugu | 341,8-9310,9 | 360 | 300 | 40 |
|  | 19 | 31-8 | - | 349,7-9321,9 | 380 | 175 | 30 |
|  | 20 | 1-9 | Kidudwe | 354,7-9317,3 | 360 | 200 | 20 |


| mapsheet | number <br> a) | $\begin{aligned} & \text { date } \\ & 1978 \end{aligned}$ | location <br> b) | co-ordinates <br> c) | eleva- <br> tion <br> d) | $\begin{aligned} & \mathrm{L} / 2 \\ & \max \end{aligned}$ | azimuth <br> e) |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| 166/1 | 21 | 1-9 | - | 349,7-9316,7 | 360 | 250 | 90 |
|  | 22 | 19-10 | Kigugu | 342,2-9312,3 | 360 | 150 | 20 |
|  | 23 | 19-10 | Kigugu | 342,3-9310,7 | 360 | 250 | 40 |
|  | 24 | 19-10 | Kigugu | 343,0-9311,0 | 360 | 175 | 40 |
|  | 25 | 19-10 | Kigugu | 343,6-9310,8 | 360 | 150 | 50 |
| 166/3 | 1 | 7-6 | - | 334,9-9289,8 | 370 | 150 | 110 |
|  | 2 | 7-6 | - | 336,4-9289,2 | 365 | 125 | 110 |
|  | 3 | 7-6 | Dakawa | 337,7-9288,2 | 360 | 150 | 155 |
|  | 4 | 7-6 | Wami SE | 339,7-9286,8 | 360 | 100 | 60 |
|  | 5 | 8-6 | - | 337,8-9289,6 | 360 | 125 | 0 |
|  | 6 | 8-6 | - | 338,1-9291,4 | 360 | 125 | 5 |
|  | 7 | 8-6 | - | 338,5-9293,3 | 360 | 250 | 15 |
|  | 8 | 8-6 | - | 338,9-9295,1 | 360 | 350 | 5 |
|  | 9 | 8-6 | Wami SE | 339,0-9286,4 | 360 | 60 | 155 |
|  | 10 | 9-6 | - | 338,9-9296,7 | 360 | 250 | 0 |
|  | 11 | 9-6 | - | 338,9-9298,3 | 360 | 125 | 5 |
|  | 12 | 9-6 | - | 338,8-9299,9 | 360 | 125 | 170 |
|  | 13 | 9-6 | - | 338,8-9301,2 | 360 | 150 | 5 |
|  | 14 | 14-6 | Dihombo | 338,5-9307,9 | 360 | 125 | 55 |
|  | 15 | 14-6 | - | 339,4-9308,9 | 360 | 125 | 35 |
|  | 16 | 14-6 | Hembeti SE | 337,0-9307,1 | 365 | 200 | 55 |
|  | 17 | 14-6 | Hembeti | 335,5-9306,9 | 375 | 200 | 90 |
|  | 18 | 15-6 | - | 339,7-9285,3 | 360 | 100 | 155 |
|  | 19 | 15-6 | Luhindo | 340,2-9283,8 | 360 | 125 | 0 |
|  | 20 | 15-6 | - | 340,5-9282,4 | 370 | 100 | 160 |
|  | 21 | 20-6 | Wami SE | 339,4-9285,2 | 360 | 100 | 60 |
|  | 22 | 20-6 | Wami SE | 340,2-9282,5 | 410 | 150 | 55 |
|  | 23 | 22-6 | Dihombo | 338,5-9307,9 | 360 | 30 | 55 |
|  | 24 | 22-6 | - | 338,9-9302,4 | 360 | 200 | 10 |
|  | 25 | 22-6 | - | 339,0-9303,5 | 360 | 175 | 5 |
|  | 26 | 22-6 | - | 339,2-9304,6 | 360 | 125 | 5 |
|  | 27 | 23-6 | - | 339,2-9305,9 | 360 | 125 | 145 |
|  | 28 | 23-6 | - | 338,9-9306,8 | 360 | 200 | 160 |
|  | 29 | 23-6 | Dihombo | 338,6-9307,6 | 360 | 175 | 155 |
|  | 30 | 23-6 | Hembeti | 337, 1-9307, 3 | 365 | 200 | 55 |
|  | 31 | 16-8 | Wami SE | 338,9-9285,0 | 360 | 200 | 60 |
|  | 32 | 16-8 | Wami SE | 338,4-9284,8 | 360 | 300 | 150 |
|  | 33 | 16-8 | - | 340,0-9286,0 | 360 | 200 | 55 |
|  | 34 | 16-8 | - | 339,3-9285,9 | 360 | 200 | 150 |
|  | 35 | 16-8 | - | 339,9-9284,8 | 360 | 100 | 150 |
|  | 36 | 17-8 | - | 338,5-9286,9 | 350 | 100 | 135 |
|  | 37 | 17-8 | Dakawa | 337,7-9288,8 | 360 | 100 | 20 |


| map- <br> sheet | number <br> a) | $\begin{aligned} & \text { date } \\ & 1978 \end{aligned}$ | location <br> b) | co-ordinates <br> c) | eleva- <br> tion <br> d) | $\begin{aligned} & \mathrm{L} / 2 \\ & \max \end{aligned}$ | azimuth <br> e) |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| 166/3 | 38 | 17-8 | - | 337,5-9288,6 | 360 | 200 | 150 |
|  | 39 | 17-8 | - | 336,5-9288,8 | 360 | 150 | 75 |
|  | 40 | 17-8 | - | 336,7-9288,0 | 360 | 175 | 45 |
|  | 41 | 17-8 | - | 337,4-9297,4 | 360 | 100 | 0 |
| 181/4 | 1 | 19-7 | Kilosa | 276,3-9245,3 | 490 | 400 | 75 |
|  | 2 | 2-8 | - | 276,8-924 ${ }^{2}$, 4 | 490 | 350 | 15 |
| 182/1 | 1 | 4-7 | Msowero | 302,9-9278,4 | 435 | 600 | 45 |
|  | 2 | 5-7 | - | 303,9-9279,5 | 420 | 500 | 35 |
|  | 3 | 5-7 | Mbugani | 304,9-9280,7 | 425 | 500 | 45 |
|  | 4 | 10-7 | Msowero | 302,1-9277,6 | 435 | 250 | 10 |
|  | 5 | 10.7 | Msowero SE | 301,5-9276,0 | 420 | 250 | 35 |
|  | 6 | 10-7 | Msowero SE | 300,6-9274,7 | 420 | 250 | 35 |
|  | 7 | 11-7 | Msowero SE | 299,8-9273,3 | 420 | 600 | 30 |
|  | 8 | 11-7 | - | 297,2-9269,7 | 235 | 125 | 20 |
|  | 9 | 11-7 | Tanganyika SE | 206,6-9268,1 | 235 | 150 | 15 |
|  | 10 | 11-7 | Tanganyika SE | 295,9-9266,9 | 440 | 150 | 30 |
|  | 11 | 11-7 | Tanganyika SE | 294,8-9265,1 | 440 | 100 | 30 |
|  | 12 | 11-7 | Rudewa-Gongoni | 293,8-9263,6 | 440 | 125 | 30 |
|  | 13 | 12-7 | Rudewa-Batini | 292,6-9260,9 | 235 | 600 | 5 |
|  | 14 | 13-7 | Rudewa-Batini | 292,4-9261,2 | 235 | 500 | 30 |
|  | 15 | 13-7 | - | 291,3-9260,1 | 435 | 600 | 60 |
|  | 16 | 20-7 | - | 289,8-9259,1. | 440 | 300 | 40 |
|  | 17 | 21-7 | Ilonga SE | 288,5-9257,8 | 445 | 150 | 40 |
|  | 18 | 24-7 | Msimba | 287,4-9256,5 | 460 | 175 | 40 |
|  | 19 | 24-7 | Msimba | 286,3-9255,3 | 470 | 400 | 40 |
|  | 20 | 24-7 | Ilonga SE | 285,2-9254,0 | 490 | 450 | 40 |
|  | 21 | 26-7 | Rudewa SE | 292,4-9259,2 | 440 | 500 | 40 |
|  | 22 | 27-7 | Madoto | 292,9-9254,3 | 440 | 700 | 150 |
|  | 23 | 28-7 | - | 292,4-9256,3 | 440 | 600 | 150 |
|  | 24 | 3-8 | Tanganyika SE | 296,8-9267,7 | 440 | 200 | 30 |
|  | 25 | 3-8 | Tanganyika SE | 297,0-9266,7 | 425 | 800 | 25 |
|  | 26 | 3-8 | - | 297,4-9264,6 | 420 | 600 | 25 |
|  | 27 | 24-8 | Msowero SE | 304,5-9276,2 | 420 | 600 | 15 |
|  | 28 | 24-8 | Msowero SE | 302,3-9277,4 | 430 | 400 | 10 |
|  | 29 | 24-8 | - | 305,7-9276,1 | 420 | 400 | 100 |
|  | 30 | 25-8 | Msowero SE | 302,9-9276,8 | 425 | 500 | 15 |
|  | 31 | 25-8 | Dundumwa | 304,9-9277,1 | 420 | 250 | 115 |
|  | 32 | 25-8 | - | 304,1-9278,5 | 425 | 175 | 80 |
|  | 33 | 26-10 | - | 294,2-9263,4 | 440 | 60 | 30 |
|  | 34 | 26-10 | - | 294,9-9263,4 | 440 | 175 | 20 |
|  | 35 | 26-10 | - | 297,3-9263,3 | 420 | 300 | 30 |


| map- | number | date | location | co-ordinates | eleva- | L/2 | azimuth |
| :--- | :--- | :--- | :--- | :--- | :--- | :--- | :--- |
| sheet |  |  |  |  |  |  |  |$\quad$| tion | max |  |  |  |  |
| :--- | :--- | :--- | :--- | :--- | :--- |
|  | a) |  | b) | c) | d) |


| 182/1 | 36 | 27-10 | Kimamba SE | 294,1-9255,9 | 430 | 500 | 175 |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  | 37 | 27-10 | Madoto SE | 295,4-9253,6 | 430 | 700 | 170 |
|  | 38 | 27-10 | - | 290,7-9258,4 | 435 | 300 | 70 |
|  | 39 | 9-11 | - | 294,6-9263,3 | 435 | 250 | 45 |
|  | 40 | 9-11 | - | 297,9-9264,4 | 420 | 500 | 15 |
|  | 41 | 9-11 | - | 295,6-9264,1 | 435 | 200 | 145 |
| 182/3 | 1 | 19-7 | - | 280,7-9246, 1 |  | 300 | 70 |
|  | 2 | 19-7 | - | 282,3-9246,5 | 475 | 300 | 70 |
|  | 3 | 19-7 | - | 283,8-9247,1 | 470 | 300 | 70 |
|  | 4 | 24-7 | Ilonga SE | 284,2-9252,7 | 500 | 300 | 45 |
|  | 5 | 25-7 | - | 281,3-9247,4 | 490 | 600 | 30 |
|  | 6 | 25-7 | - | 282,3-9249,2 | 490 | 500 | 25 |
|  | 7 | 25-7 | Ilonga | 283,0-9250,6 | 490 | 175 | 25 |
|  | 8 | 25-7 | - | 285,4-9247,5 | 465 | 250 | 75 |
|  | 9 | 25-7 | - | 287,1-9248,4 | 455 | 700 | 35 |
|  | 10 | 26-7 | Mpirani | 288,5-9249,3 | 450 | 800 | 65 |
|  | 11 | 26-7 | Mpirani | 289,8-9250,7 | 450 | 800 | 40 |
|  | 12 | 27-7 | Kimamba | 293,8-9251,0 | 440 | 600 | 160 |
|  | 13 | 27-7 | Madoto SE | 293,4-9252,8 | 440 | 600 | 160 |
|  | 14 | 28-7 | - | 289,8-9252,5 | 440 | 700 | 80 |
|  | 15 | 28-7 | Kimamba | 293,8-9251,0 | 440 | 250 | 160 |
|  | 16 | 31-7 | - | 295,5-9249,2 | 435 | 300 | 160 |
|  | 17 | 31-7 | Kimamba | 294,8-9248,9 | 435 | 400 | 160 |
|  | 18 | 1-8 | Kimamba | 293,4-9249,9 | 440 | 400 | 100 |
|  | 19 | 1-8 | Kimamba | 293,5-9249,6 | 440 | 500 | 75 |
|  | 20 | 1-8 | Kimamba | 296,0-9250,1 | 435 | 700 | 75 |
|  | 21 | 2-8 | Tanzania SE | 280,1-9239,9 | 480 | 200 | 100 |
|  | 22 | 2-8 | Tanzania SE | 280,0-9240,0 | 480 | 350 | 100 |
|  | 23 | 4-8 | Kimamba | 294,7-9250,9 | 435 | 700 | 80 |
|  | 24 | 4-8 | Madoto SE | 296,3-9252,0 | 435 | 800 | 80 |
|  | 25 | 4-9 | Mbwade | 297,7-9252,9 | 430 | 600 | 170 |
|  | 26 | 4-9 | Mbwade | 298,5-9252,4 | 430 | 400 | 80 |
|  | 27 | 7-9 | - | 299,8-9252,6 | 430 | 200 | 80 |
|  | 28 | 7-9 | - | 300,9-9252,8 | 425 | 600 | 80 |
|  | 29 | 7-9 | - | 302,1-9252,9 | 420 | 500 | 80 |
|  | 30 | 8-9 | Mbwade | 298,8-9251,8 | 430 | 500 | 170 |
|  | 31 | 8-9 | - | 299,3-9253, 1 | 430 | 600 | 180 |
|  | 32 | 8-9 | - | 303,7-9252,6 | 420 | 700 | 100 |
|  | 33 | 11-9 | - | 305,8-9251,7 | 420 | 700 | 80 |
|  | 34 | 11-10 | Mkata Ujamaa | 302,1-9231,3 | 430 | 175 | 140 |
|  | 35 | 11-10 | Mkata Ujamaa | 303,3-9230,5 | 430 | 150 | 145 |


| map- <br> sheet | number <br> a) | $\begin{gathered} \text { date } \\ 1978 \end{gathered}$ | location <br> b) | co-ordinates <br> c) | eleva- <br> tion <br> d) | $\begin{aligned} & \mathrm{L} / 2 \\ & \max \end{aligned}$ | azimuth <br> e) |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| 182/3 | 36 | 11-10 | Mkata Ujamaa | 302,6-9230,8 | 430 | 350 | 45 |
|  | 37 | 12-10 | Mkata Ujamaa | 301,9-9231,0 | 430 | 75 | 45 |
|  | 38 | 12-10 | Mkata Ujamaa | 302,7-9231,0 | 430 | 400 | 45 |
|  | 39 | 12-10 | - | 304,5-9229,5 | 440 | 100 | 135 |
|  | 40 | 12-10 | - | 305,7-9228,4 | 450 | 200 | 165 |
|  | 41 | 24-10 | Kilangali | 287,2-9231,9 | 450 | 150 | 135 |
|  | 42 | 24-10 | Kwungu SE | 286,2-9232,3 | 450 | 250 | 130 |
|  | 43 | 24-10 | - | 285,0-9233,2 | 460 | 200 | 130 |
|  | 44 | 24-10 | Kondoa | 284,4-9246,2 | 465 | 350 | 70 |
|  | 45 | 25-10 | Malangali | 290,3-9244,1 | 450 | 200 | 90 |
|  | 46 | 25-10 |  | 288,8-9244,2 | 450 | 100 | 65 |
|  | 47 | 25-10 | Mabwere-bwere | 287,3-9243,8 | 455 | 200 | 55 |
|  | 48 | 25-10 | Manyema | 285,8-9243,6 | 460 | 400 | 95 |
|  | 49 | 26-10 | - | 282,7-9245,5 | 470 | 400 | 70 |
|  | 50 | 10-11 | - | 289,5-9244,4 | 450 | 500 | 90 |
|  | 51 | 10-11 | Mabwere-bwere | 287,1-9243,3 | 455 | 200 | 90 |
|  | 52 | 10-11 | - | 284,1-9244,9 | 465 | 100 | 90 |
|  | 53 | 13-10 | - | 293,0-9247,2 | 445 | 600 | 75 |
|  | 54 | 13-11 | - | 282,5-9244,9 | 470 | 800 | 85 |
| 182/4 | 1 | 11-9 | - | 307,6-9252,0 | 415 | 700 | 85 |
|  | 2 | 11-9 | - | 309,3-9252,2 | 415 | 700 | 85 |
|  | 3 | 12-9 | - | 310,7-9252,4 | 410 | 700 | 85 |
|  | 4 | 12-9 | - | 312,3-9252,6 | 410 | 700 | 86 |
|  | 5 | 12-9 | - | 313,9-9252,8 | 405 | 700 | 85 |
|  | 6 | 13-9 | $\cdots$ | 315,5-9252,9 | 400 | 700 | 85 |
|  | 7 | 13-9 | - | 317,4-9252,9 | 400 | 400 | 90 |
|  | 8 | 13-9 | - | 320,4-9252,9 | 400 | 250 | 105 |
|  | 9 | 13-9 | - | 319,4-9252,8 | 400 | 500 | 95 |
|  | 10 | 14-9 | - | 321,6-9252,1 | 400 | 200 | 145 |
|  | 11 | 14-9 | - | 322,3-9250,7 | 410 | 150 | 160 |
|  | 12 | 14-9 | - | 322,8-9249,1 | 420 | 300 | 160 |
|  | 13 | 14-9 | - | 323,1-9257,7 | 420 | 175 | 180 |
|  | 14 | 14-9 | - | 324,1-9246,6 | 430 | 150 | 125 |
|  | 15 | 14-9 | - | 325,5-9245,9 | 440 | 350 | 115 |
|  | 16 | 14-9 | - | 326,9-9245,2 | 450 | 200 | 115 |
|  | 17 | 15-9 | - | 327,8-9244,0 | 460 | 200 | 150 |
|  | 18 | 15-9 | * | 328,5-9242,6 | 465 | 175 | 0 |
|  | 19 | 15-9 | - | 328,6-9240,9 | 480 | 150 | 160 |
|  | 20 | 15-9 | - | 328,6-9240,8 | 480 | 150 | 160 |
|  | 21 | 15-9 | - | 329,1-9239,2 | 500 | 150 | 0 |
|  | 22 | 26-9 | - | 329,8-9236,6 | 520 | 175 | 75 |


| mapsheet | number <br> a). | $\begin{aligned} & \text { date } \\ & 1978 \end{aligned}$ | location <br> b) | co-ordinates <br> c) | eleva- <br> tion <br> d) | $\begin{aligned} & \mathrm{L} / 2 \\ & \max \end{aligned}$ | azimuth <br> e) |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| 182/4 | 23 | 26-9 | - | 331,8-9236,9 | 520 | 150 | 120 |
|  | 24 | 26-9 | - | 333,8-9235,8 | 570 | 150 | 135 |
|  | 25 | 13-10 | - | 308,3-9226,4 | 470 | 150 | 120 |
|  | 26 | 13-10 | - | 309,1-9226,1 | 470 | 175 | 110 |
| 183/1 | 1 | 15-6 | - | 341,2-9281,1 | 375 | 150 | 150 |
|  | 2 | 15-6 | - | 341,6-9279,5 | 385 | 125 | 0 |
|  | 3 | 15-6 | - | 341,9-9278,0 | 390 | 300 | 180 |
|  | 4 | 16-6 | - | 342,1-9276,5 | 395 | 125 | 155 |
|  | 5 | 16-6 | - | 342,6-9275,1 | 400 | 100 | 145 |
|  | 6 | 16-6 | - | 343,4-9273,9 | 415 | 300 | 150 |
|  | 7 | 16-6 | - | 344,1-9272,3 | 425 | 200 | 165 |
|  | 8 | 16-6 | - | 344,2-9270,9 | 430 | 100 | 155 |
|  | 9 | 16-6 | - | 344,9-9269,7 | 440 | 200 | 140 |
|  | 10 | 19-6 | - | 345,5-9268,1 | 460 | 75 | 150 |
|  | 11 | 19-6 | - | 346,2-9266,7 | 480 | 100 | 135 |
|  | 12 | 19-6 | - | 347, 0-9263,5 | 490 | 150 | 155 |
|  | 13 | 19-6 | - | 347,5-9263,9 | 510 | 200 | 160 |
|  | 14 | 20-6 | - | 344,2-9270,6 | 430 | 75 | 155 |
|  | 15 | 20-6 | - | 341,7-9279,9 | 380 | 100 | 60 |
|  | 16. | 20-6 | - | 342,0-9277,5 | 390 | 150 | 80 |
|  | 17 | 21-9 | - | 351,0-0254,5 | 510 | 100 | 150 |
|  | 18 | 21-9 | - | 350, 7-9256,2 | 520 | 150 | 165 |
|  | 19 | 21-9 | Mkundi | 350,4-9257,7 | 570 | 200 | 165 |
|  | 20 | 21-9 | Mkundi | 349,7-9259,2 | 580 | 200 | 145 |
|  | 21 | 21-9 | Mkundi | 348,7-9260,5 | 580 | 125 | 155 |
|  | 22 | 25-9 | - | 348,3-9262,0 | 560 | 250 | 160 |
|  | 23 | 28-9 | Tungi SE | 357,6-9254,8 | 480 | 100 | 85 |
|  | 24 | 28-9 | Tungi SE | 357,4-9256,0 | 490 | 100 | 85 |
|  | 25 | 28-9 | Tungi SE | 357,3-9256,2 | 500 | 200 | 85 |
| 183/2 | 1 | 4-10 | Mikese | 379,8-9255,9 | 380 | 30 | 10 |
| 183/3 | 1 | 30-5 | Kihonda | 351,9-9250,6 | 480 | 150 | 175 |
|  | 2 | 30-5 | Kihonda | 351,8-9251,6 | 480 | 100 | 160 |
|  | 3 | 21-9 | Kihonda | 351,2-9252,8 | 490 | 300 | 160 |
|  | 4 | 25-9 | - | 352,0-9249,7 | 500 | 300 | 175 |
|  | 5 | 25-9 |  | 352,3-9248,2 | 510 | 100 | 175 |
|  | 6 | 25-9 | Morogoro | 352,4-9246,0 | 500 | 100 | 175 |
|  | 7 | 25-9 | Morogoro | 352,9-9245,1 | 500 | 100 | 175 |
|  | 8 | 26-9 |  | 334,8-9234,6 | 600 | 100 | 120 |
|  | 9 | 26-9 | Kinyenze | 336,3-9233,4 | 570 | 175 | 90 |


| map- <br> sheet | number <br> a) | $\begin{aligned} & \text { date } \\ & 1978 \end{aligned}$ | location <br> b) | co-ordinates <br> c) | elevation <br> d) | L/2 <br> $\max$ | azimuth <br> e) |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| 183/3 | 10 | 28-9 | Tungi SE | 358,3-9252,4 | 480 | 100 | 90 |
|  | 11 | 28-9 | Tungi SE | 357,6-9250,9 | 480 | 100 | 90 |
|  | 12 | 28-9 | Tungi SE | 357,5-9249,4 | 500 | 175 | 90 |
|  | 13 | 29-9 | Morogoro | 353,3-9243,7 | 540 | 75 | 30 |
|  | 14 | 29.9 | Morogoro | 353, 7-9244, 2 | 540 | 100 | 145 |
|  | 15 | 29-9 | Tanke SE | 350,3-9242,9 | 520 | 75 | 40 |
|  | 16 | 2-10 | - | 338,3-9232,7 | 560 | 350 | 145 |
|  | 17 | 2-10 | Mlali | 338,3-9231,4 | 580 | 175 | 110 |
|  | 18 | 2-10 | Mlali | 338,2-9230,5 | 580 | 100 | 20 |
|  | 19 | 13-10 | Tanke SE | 349,5-9244,4 | 505 | 125 | 135 |
| 183/4 | 1 | 4-10 | - | 379,2-9250,7 | 420 | 100 | 10 |
|  | 2 | 4-10 | - | 379,3-9250,1 | 400 | 50 | 130 |
|  | 3 | 4-10 | - | 379,1-9251,3 | 400 | 75 | 170 |
|  | 4 | 4-10 | - | 379,2-9253,8 | 400 | 100 | 0 |
|  | 5 | 5-10 | - | 366,1-9248,4 | 490 | 100 | 50 |
|  | 6 | 5-10 | Mgugu | 72,6-9247,9 | 440 | 100 | 120 |
|  | 7 | 5-10 | - | 366,9-9248,8 | 485 | 300 | 60 |
|  | 8 | 5-10 | Pangawe | 365,2-9247,7 | 540 | 100 | 90 |
|  | 9 | 5-10 | Pangawe | 365,7-9247,6 | 520 | 150 | 130 |
| 184/1 | 1 | 6-10 | - | 393,4-9264,9 | 280 | 75 | 60 |
|  | 2 | 6-10 | - | 393,5-9264,8 | 280 | 125 | 140 |
|  | 3 | 6-10 | - | 392,6-9265,0 | 280 | 75 | 140 |
|  | 4 | 6-10 | - | 393,5-9265,5 | 280 | 30 | 35 |
|  | 5 | 6-10 | - | 394,1-9264,2 | 280 | 75 | 0 |
| 200/2 | 1 | 13-10 | - | 309,8-9225,6 | 480 | 175 | 40 |

DATA DD 7

THE GEO-ELECTRICAL SOUNDINGS AND THEIR INTERPRETATIONS





















































































































































































































































## DATA DD 8

TIME-DRAWDOWN AND TIME-RECOVERY GRAPHS OF PUUPING TESTS FROM MDWSP-BOREHOLES




meetpapier-wormerveer




neetpapier - wormerveer
17 Tr


PUMPING TESTS © AND © OF BOREHOLE 135c/78, DAKAWA


RECOVERY TESTS (1) AND (3) OF BOREHOLE 135e/78, DAKAWA


PUMPING TEST OF BOREHOLE 151/78, DIBAMBA


PUMPING TEST (1) OF BOREHOLE 162/78, MANDELA



RECOVERY TEST (3) OF BOREHOLE $162 / 78$, MANDELA


PUMPING TEST OF BOREHOLE 173/78, MAKUYU


## DATA DD 9

SUMMARY OF hand drillings in the ngerengere area

DD 9 - Sumary of MwCP hand drillings in the Ngerengere area

| village | hand drilling no. | total depth <br> ( m -GL) | aquifer depth <br> from - to $(m-G L)$ | aquifer thickness <br> (m) | aquifer type | water level <br> Juring drilling <br> (m-GL) | tested yield (1/HR) | maximum drawdown | ${ }_{\left(\mathrm{m}^{2} / \text { day }\right)}$ | $\begin{aligned} & \mathrm{K} \\ & \text { (요/day) } \end{aligned}$ | EC during remarks drilling (ms/m) |  |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| Kihonda | A-4 | 2.1 | 1.3-2.1 | 0.8 | loamy sand | 1.3 | - | - | - | - | 750 |  |
|  | A-5 | 4.0 | 2.0-3.0 | 1.0 | fine sand | 1.0 | - | - | - | - | 750 |  |
|  | A-6 | 5.0 | 1.9-4.5 | 2.6 | median sand | 0.6 | - | - | - | - | 350 |  |
|  | A-7 | 6.0 | 1.5-3.0 | 1.5 | median sand | 0.6 | - | - | - | - | 1375 |  |
|  | A-B | 11.5 | 1.3-3.5 | 2.2 | median sand | 0.6 | - | - | - | - | 7350 |  |
|  |  |  | 5.9-11.5 | 5.6 | fine sand |  |  |  |  |  |  |  |
|  | A-9 | 10.0 | 2.0-2.7 | 0.7 | clay sand | 0.3 | - | - | - | - | 7400 |  |
|  |  |  | 8.2-8.8 | 0.6 | fine sand |  |  |  |  |  |  |  |
|  | A-10 | 3.0 | 0.0-0.5 | 0.5 | fine sand | 0.2 | - | - | - | - | 150 |  |
|  |  |  | 2.8-3.0 | 0.2 | coarse sand |  |  |  |  |  |  |  |
|  | A-11 | 1.7 | 0.8-1.7 | 0.9 | clay sand | 1.26 | - | - | - | - | 100 |  |
|  | A-13 | 10.0 | 1.6-3.0 | 1.4 | median fine | 0.3 | - | - | - | - | 290 |  |
|  | A-14 | 11.0 | 3.0-4.0 | 1.0 | median <br> coarse | 0.8 | - | - | - | - | 1000 |  |
|  |  |  | 9.0-11.0 | 2.0 | loamy sand |  |  |  |  |  |  | m |
| Tungi | B-6 | 4.0 | 3.6-4.0 | 0.4 | coarse sand | 3.6 | - | - | - | - | 800 |  |
|  | B-9 | 2.8 | 1.7-2.8 | 1.1 | loamy sand | 1.75 | - | - " | - | - | 2050 |  |
|  | B-10 | 6.0 | 2.0-6.0 | 4.0 | median coarse | 1.3 | - | - | - | - | 1300 |  |
|  | B-11 | B. 0 | 2.0-3.0 | 1.0 | median sand | 0.8 | - | - | - - | - | 1050 |  |
|  | B-12 | 7.6 | 3.0-4.0 | 1.0 | median sand | 0.7 | - | - | - | - | 1200 |  |
|  | B-13 | 10.0 | 1.5-2.0 | 0.5 | median sand | 0.5 | - | - | - | - | 1900 |  |
|  | B-14 | 10.2 | 3.0-3.5 | 0.5 | median sand | 0.5 | - | - | - | - | 2500 |  |
|  | B-15 | 10.0 | 0.0-4.1 | 4.1 | median sand | 0.5 | - | - | - | - | 2350 |  |
|  |  |  | 7.0-10.0 | 3.0 | median saod |  |  |  |  |  |  |  |
|  | B-16 | 8.0 | 2.0-5.2 | 3.2 | median sand | 0.3 | - | - | - | - | 1100 |  |
|  | B-17 | 2.0 | 1.5-2.0 | 0.5 | clay sand | 0.4 | - | - | - | - | 800 |  |
|  | B-19 |  | 3.0-7.0 | 4.0 | median fine | 0.3 | - | - | - | - | 700 |  |
|  | B-20 | 7.0 | 1.2-1.8 | 0.6 | fine sand | 0.3 | - | - | - | - | 600 |  |
|  |  |  | 3.4-7.0 | 4.0 | median <br> coarse |  |  |  |  |  |  |  |
| Mkambarani | $\mathrm{c}^{1}-1$ | 6.0 | 2.8-5.6 | 2.8 | median fine | 0.5 | - | - | - | - | 650 |  |
|  | $c^{1}-2$ | 4.0 | 3.6-4.0 | 0.4 | fine sand | 0.3 | - | - | - | - | 580 |  |
|  | $c^{1}-3$ | 4.5 | 3.0-4.5 | 1.5 | clay sand | 0.2 | - | - | - | - | 380 |  |
|  | $\mathrm{C}^{1}-4$ | 5.0 | 2.3-3.5 | 1.2 | fine sand | 0.4 | - | - | - | - | 160 |  |
|  | $\mathrm{C}^{1}-5$ | 5.2 | 4.5-5.2 | 0.7 | median <br> coarse | 0.6 | - | - | - | - | 350 |  |
|  | $c^{1}-6$ | 2.0 | 1.4-2.0 | 0.6 | fine sand | 0.8 | - | - | - | - | 120 |  |
|  | $\mathrm{C}^{1}-7$ | 5.0 | 3.0-5.0 | 2.0 | median coarse | 0.5 | - | - | - | - | 135 |  |
|  | $\mathrm{c}^{2}-1$ | 5.0 | 4.0-4.5 | 1.5 | coarse median fine | 0.7 | - | - | - | - | 72 |  |
|  | $\mathrm{c}^{2}-2$ | 6.5 | 1.3-2.3 | 1.0 | median fine | 0.6 | - | - | - | - | 20 |  |




DATA DD 10

CHEMICAL ANALYSES OF GROUND WATER FROM HAND DRILLINGS IN THE NGERENGERE AREA

DD 10 - Chemical analyses of ground water from hand drillings in the Ngerengere area

| location borehole nr. date of sampling date of analysis sample method |  | $\begin{aligned} & \text { Kihonda } \\ & \text { A-4 } \\ & 12-12-78 \\ & 14-12-78 \\ & \text { bailer } \end{aligned}$ | $\begin{aligned} & \text { Kihonda } \\ & \text { A-6 } \\ & \text { 13-12-78 } \\ & 14-12-78 \\ & \text { bailer } \end{aligned}$ | Kihonda A-7 <br> 13-12-78 <br> 14-12-78 <br> bailer | $\begin{aligned} & \text { Kihonda } \\ & \text { A-8 } \\ & 19-12-78 \\ & 20-12-78 \\ & \text { bsiler } \end{aligned}$ | $\begin{aligned} & \text { Kihonda } \\ & \text { A-9 } \\ & 19-12-78 \\ & 20-12-78 \\ & \text { bailer } \end{aligned}$ | $\begin{aligned} & \text { Kinhonda } \\ & \text { A-11 } \\ & 18-12-78 \\ & 20-12-78 \\ & \text { bailer } \end{aligned}$ | $\begin{aligned} & \text { Kihonda } \\ & \text { A-13 } \\ & 18-12-78 \\ & 11-01-78 \\ & \text { bailer } \end{aligned}$ | $\begin{aligned} & \text { Kihonda } \\ & \text { A-14 } \\ & \text { 18-12-78 } \\ & \text { 11-01-78 } \\ & \text { bailer } \end{aligned}$ | bailer | bailer | bailer | bailer |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| Water analysis | unit |  |  |  |  |  |  |  |  |  |  |  |  |
| electrical conductivity | $\mathrm{ms} / \mathrm{m}\left(25^{\circ} \mathrm{C}\right)$ | 925 | 300 | 1475 | 400 | 400 | 127 | 295 | 1000 |  |  |  |  |
| pH |  | 8.0 | 7.1 | 7.3 | 7.8 | 8.2 | 7.5 | 7.1 | 7.5 |  |  |  |  |
| calcium | mg/I $\mathrm{Ca}^{2+}$ | 52 | 128 | 300 | 40 | 112 | 40 | 112 | 120 |  |  |  |  |
| magnesium | $\mathrm{mg} / \mathrm{Mg}^{2+}$ | 182 | 115 | 349 | 55 | 350 | 24.3 | 112 | 364 |  |  |  |  |
| sodium | $m \mathrm{~m} / \mathrm{I} \mathrm{Na}{ }^{+}$ | 1564 | 161 | 2737 | 805 | - | 299 | 322 | 1449 |  |  |  |  |
| total harduess | $\mathrm{mg} / \mathrm{ICaCO}_{3}$. | 880 | 795 | 2180 | 324 | 1640 | 200 | 740 | 1800 |  |  |  |  |
| iron | $\mathrm{mg} / \mathrm{T} \mathrm{Fe}$ | 0.03 | 0.2 | 0.15 | 1.0 | 0.23 | 0.58 | 0.05 | 0.09 |  |  |  |  |
| manganese | mg/ 1 Mm |  |  |  |  |  |  | 0.9 | 1.6 |  |  |  |  |
| fluoride | mg/ $\mathrm{FF}^{-}$ | 1.0 | 0.5 | 0.7 | 3.0 | 1.1 | 0.5 | 0.8 | 1.5 |  |  |  |  |
| chloride | $\mathrm{mg} / \mathrm{ICl}^{-}$ | 2525 | 630 | 5150 | 1050 | 513 | 200 | 700 | 3050 |  |  |  |  |
| bicarbonate | $\mathrm{mg} / \mathrm{L} \mathrm{HCO}_{3}{ }^{-}$ | 915 | 305 | 1098 | 610 | 671 | 440 | 560 | 790 |  |  |  |  |
| nitrate | $\mathrm{mg} / \mathrm{IHO} \mathrm{H}^{-}$ | 2,7 | 2.2 | 5.3 | 3.1 | 3.1 | 12.4 | 4 | 8.5 |  |  |  |  |
| sulphate | $\mathrm{mg} / \mathrm{ISO}_{4}{ }^{2-}$ | 28 | 17 | 54 | 50 | 52 | 180 | 14 | 17 |  |  |  |  |
| phosphate (ortho) |  | 1.4 | 1.1 | 0.25 | 1.5 | 0.3 | 1.8 | 0.5 | 0.2 |  |  |  |  |

Conlinued 1

| location borehole nr. date of sampling date of analysis sample method |  | $\begin{aligned} & \text { Tungi } \\ & \text { B-6 } \\ & 29-12-78 \\ & 8-01-79 \\ & \text { bailer } \end{aligned}$ | $\begin{aligned} & \text { Tungi } \\ & \mathrm{B}-9 \\ & 1-01-79 \\ & 8-01-79 \\ & \text { bailer } \end{aligned}$ | $\begin{aligned} & \text { Tungi } \\ & \text { B-10 } \\ & 2-01-79 \\ & 8-01-79 \\ & \text { bailer } \end{aligned}$ | $\begin{aligned} & \text { Tungi } \\ & \text { B-11 } \\ & 3-01-79 \\ & 8-01-79 \\ & \text { bailer } \end{aligned}$ | Tungi $\mathrm{B}-12$ $4-01-79$ $8-01-79$ <br> bailer | $\begin{aligned} & \text { Tungi } \\ & \mathrm{B}-13 \\ & 5-01-79 \\ & 8-01-79 \\ & \text { bailer } \end{aligned}$ | $\begin{aligned} & \text { Tungi } \\ & \text { B-14 } \\ & 8-01-79 \\ & 11-01-79 \\ & \text { bailer } \end{aligned}$ | $\begin{aligned} & \text { Tungi } \\ & \text { B-15 } \\ & 10-01-79 \\ & 11-01-79 \\ & \text { bailer } \end{aligned}$ | $\begin{aligned} & \text { Tuagi } \\ & \text { B-i6 } \\ & \text { Feb. } 79 \\ & \text { Feb. } 79 \\ & \text { bailer } \end{aligned}$ | $\begin{aligned} & \text { Tungi } \\ & \text { B-17 } \\ & \text { Feb. } 70 \\ & \text { Feb. } 79 \\ & \text { bailer } \end{aligned}$ | Tungi <br> B-19 <br> Feb. 79 <br> Feb. 79 <br> bailer | $\begin{aligned} & \text { Tungi } \\ & \text { B-20 } \\ & \text { Yeb. } 79 \\ & \text { Feb. } 79 \\ & \text { bailer } \end{aligned}$ |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| Water analysis | unit |  |  |  |  |  |  |  |  |  |  |  |  |
| electrical conductivity | $\mathrm{mS} / \mathrm{m}\left(25^{\circ} \mathrm{C}\right)$ | 800 | 1980 | 1100 | 1320 | 1325 | 3800 | 2550 | 2345 | 725 | 525 | 800 | 500 |
| pH |  | 7.7 | 7.1 | 8.0 | 7.2 | 7.8 | 7.0 | 7.5 | 7.3 | 8.2 | 8.3 | 8.1 | 7.9 |
| calcium | $\mathrm{mg} / \mathrm{I} \mathrm{Ca}{ }^{2+}$ | 720 | 1000 | 160 | 930 | 240 | 1440 | 200 | 350 | 170 | 60 | 150 | 100 |
| magnes ium | $\mathrm{mg} / \mathrm{Mg}^{2+}$ | 425 | 364 | 61 | 164 | 18 | 486 | 152 | 304 | 151.88 | 400.95 | 115.43 | 255.2 |
| sodium | mg/I $\mathrm{Na}^{+}$ | 1840 | 3588 |  | 2116 | 2714 | 6233 |  | 4531 | 1104 | 621 | 851 | 230 |
| total hardness | $\mathrm{mg} / \mathrm{ICaCO}_{3}$ | 3550 | 4000 | 650 | 3000 | 675 | 5600 | 1125 | 2125 | 1050 | 1800 | 850 | 1300 |
| iront | $\mathrm{mg} / \mathrm{I} \mathrm{Fe}$ | 1.5 | 3.2 | 0.65 | 1.2 | 0.2 | 1.0 | 0.5 | 1.3 | 0.04 | 0.06 | 0.08 | 1.3 |
| manganese | mg/I Mn | 0.4 | 0.5 | 2.4 | 10.0 | 0.6 | 1.6 | 0.04 | 5 | 0.4 | 0.7 | 0.8 | 0.2 |
| fhuoride | $\mathrm{mg} / \mathrm{T} \mathrm{F}^{-}$ | 0.4 | 0.2 | 0.5 | 0.3 | 0.6 | 0.2 | 0.3 | 0.3 | 0.9 | 1.0 | 0.7 | 1.0 |
| chloride | $\mathrm{mg} / \mathrm{I} \mathrm{Cl}^{-}$ | 4800 | 7900 | 3200 | 4875 | 3700 | 1275 | 6625 | 6375 | 1750 | 450 | 1550 | 200 |
| bicarbouate | $\mathrm{rag} / \mathrm{I} \mathrm{HCO}_{3}{ }^{-}$ | 976 | 732 | 854 | 824 | 1460 | 1220 |  | 1677 | 976 | 3050 | 488 | 1830 |
| nitrate | $\mathrm{mg} / 1 \mathrm{HO}_{3}{ }^{-}$ | 8.4 | 24 | 4.4 | 7 | 8 | 4.4 | 8.4 | 20 | 8.86 | 27.46 | 8.42 | 22.15 |
| sulplate | $\text { mg/ } \mathrm{I} \mathrm{SO}_{4}{ }^{2-}$ | 4 | 55 | 30 | 17 | 125 | 125 | 180 | 520 | 200 | 3 | 145 | 2 |
| phosphate (ortho) | $\mathrm{mg} / \mathrm{IPO}_{4}^{3-}$ | 0.7 | 0.3 | 0.5 | 0.5 | 0.35 | 0.4 | 0.15 | 0.1 | 0.8 | 1.1 | 0.3 | 0.3 |

Continued 2

| Jocation borehole irr. date of sampling date of analysis sample method |  | Mkambarani <br> $C^{1}$ - 1 <br> Feb. 79 <br> Feb. 79 <br> bailer | Mkamharani $C^{1}-2$ <br> Feb. 79 <br> Feb. 79 <br> bailer | Mkambarani $C^{1-3}$ <br> Feb. 79 <br> Feb. 79 <br> bailer | Mkamharani $C^{1-4}$ <br> Feb. 79 <br> Feb. 79 bailer | Mkambarani $C^{2}-1$ <br> Feb. 79 <br> Feb. 79 bailer | Mkambarani $c^{2}-2$ <br> Feb. 79 <br> Feb. 79 <br> bailer | Mkambarani $c^{2}-3$ <br> Feb. 79 <br> Feb. 79 bailer | Mkambarani $c^{2}-4$ <br> Feb. 79 <br> Feb. 79 bailer | bailer | bailer | bailer | bailer |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| Water analysis | unit |  |  |  |  |  |  |  |  |  |  |  |  |
| electrical conductivity | $\mathrm{mS} / \mathrm{mL}\left(25^{\circ} \mathrm{C}\right)$ | 625 | 575 | 401 | 130 | 80 | 28 | 81 | 200 |  |  |  |  |
| pH |  | 7.3 | 7.3 | 7.5 | 7.5 | 7.6 | 5.8 | 6.8 | 6.1 |  |  |  |  |
| calcium | mg/I Ca ${ }^{2+}$ | 410 | 400 | 270 | 112 |  |  |  | 144 |  |  |  |  |
| magnes ium | $\mathrm{mg} / \mathrm{Hg}^{2+}$ | 176.2 | 109.35 | 115.43 | 38.88 |  |  |  | 32.48 |  |  |  |  |
| sodium | $\mathrm{mg} / \mathrm{I} \mathrm{Na}{ }^{+}$ | 483 | 483 | 391 | 46 |  |  |  |  |  |  |  |  |
| total hardness | $\mathrm{mg}^{\text {/ }}$ [ $\mathrm{CaCO}_{3}$ | 1750 | 1450 | 1150 | 440 |  |  |  | 720 |  |  |  |  |
| iron | $\mathrm{mg} / \mathrm{Y} \mathrm{Fe}$ | 0.04 | 0.05 | 0.2 | 0.35 | 0.95 |  | 1.9 | 0.8 |  |  |  |  |
| manganese | mg/I Mn | 3.0 | 5.0 | 1.8 | 1.4 | 0.4 |  | 0.8 | 2.5 |  |  |  |  |
| fluoride | mg/I $\mathrm{F}^{-}$ | 0.5 | 0.2 | 0.4 | 0.2 |  | 0.1 |  | 0.2 |  |  |  |  |
| chloride | $\mathrm{mg} / \mathrm{I} \mathrm{Cl}{ }^{-}$ | 1600 | 1650 | 1200 | 310 | 0.3 |  | 0.1 | 150 |  |  |  |  |
| bicarbonate |  | 732 | 244 | 396.5 | 97.6 |  |  |  | 134.2 |  |  |  |  |
| nitrate | $\mathrm{mg} / \mathrm{I} \mathrm{HO}_{3}{ }^{-}$ | 0.89 | 1.33 | 1.77 | 0.89 | 0.89 |  | 0.89 | 0.89 |  |  |  |  |
| sulphate | $\mathrm{mg} / \mathrm{I} \mathrm{SO} 4^{2-}$ | 30 | 40 | 45 | 5 | 10 |  | 5 | 10 |  |  |  |  |
| phosphate (ortho) | $\mathrm{mg} / \mathrm{PO}_{4}{ }^{3-}$ | 0.4 | 0.1 | 0.2 | 0.35 |  |  |  | 0.25 |  |  |  |  |

Continued 3

| location boretiole mr. date of sampling riate of analysis sample methor |  | Lubungo $\mathrm{F}-2$ <br> March 79 <br> March 79 <br> bailer | L.ubungo F-4 <br> Harch 79 <br> March 79 bailer | Lubungo $\mathbf{F - 5}$ <br> March 79 <br> March 79 <br> bailer | bailer | bailer | bailer | bailer | bailer | bailer | bailer | bailer | bailer |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| Hater analysis | unit |  |  |  |  |  |  |  |  |  |  |  |  |
| electrical conductivity | - $\mathrm{S} / \mathrm{m}\left(25^{\circ} \mathrm{C}\right)$ | 450 | 577 | 147 |  |  |  |  |  |  |  |  |  |
| pli |  | 6.5 | 6.6 | 6.7 |  |  |  |  |  |  |  |  |  |
| calcium | $\mathrm{mg} / \mathrm{ICa}{ }^{2+}$ | 210 | 200 | 152 |  |  |  |  |  |  |  |  |  |
| magnesium | $\mathrm{mg} / \mathrm{Mg}^{2+}$ | 151.88 | 145.8 | 48.6 |  |  |  |  |  |  |  |  |  |
| sodium | mg/I $\mathrm{Na}^{+}$ | 391 | 690 | 46 |  |  |  |  |  |  |  |  |  |
| total hardness | $\mathrm{mg} / \mathrm{ICaCO}_{3}$ | 1150 | 1100 | 580 |  |  |  |  |  |  |  |  |  |
| iron | $\mathrm{mg} / \mathrm{IFe}$ | 0.1 | 3.0 | 3.2 |  |  |  |  |  |  |  |  |  |
| mangatese | tag/I Mn | 1.0 | 4.5 | 4.3 |  |  |  |  |  |  | . |  | - |
| fluoride | mg/ $/ 1 \mathrm{~F}^{-}$ | 0.2 | 0.4 | 0.3 |  |  |  |  |  |  |  |  |  |
| chloride | mg/I $\mathrm{Cl}^{-}$ | 1200 | 1500 | 180 |  |  |  |  |  |  |  |  |  |
| bicarbonate | $\mathrm{mg} / \mathrm{I} \mathrm{HCO}_{3}{ }^{-}$ | 183 | 448 | 561.2 |  |  |  |  |  |  |  |  |  |
| mitrate | mg/ $\mathrm{HO}_{3}$ | 10.19 | 10.63 | 17.72 |  |  |  |  |  |  |  |  |  |
| satphate | $\mathrm{mg} / \mathrm{I} \mathrm{SO}_{4}{ }^{2-}$ | 200 | 130 | 0 |  |  |  |  |  |  |  |  |  |
| phosphate (ortho) | $\mathrm{mg} / \mathrm{I} \mathrm{PO}_{4}{ }^{3-}$ | 0.25 | 1.0 | 1.1 |  |  |  |  |  |  |  |  |  |

Continued 4

| location borehole nr. date of sampling date of analysis sample method |  | $\begin{aligned} & \text { Visalaka } \\ & \text { J-5 } \\ & \text { March } 79 \\ & \text { March } 79 \\ & \text { bailer } \end{aligned}$ | $\begin{aligned} & \text { Visalaka } \\ & \text { J-8 } \\ & \text { March } 79 \\ & \text { Harch } 79 \\ & \text { bailer } \end{aligned}$ | bailer | bailer | bailer | bailer | bailer | bailer | bailer | bailer | bailer | bailer |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| Hater analysis | unit |  |  |  |  |  |  |  |  |  |  |  |  |
| electrical conductivity | $\mathrm{mS} / \mathrm{m}\left(25^{\circ} \mathrm{C}\right.$ ) | 350 | 320 |  |  |  |  |  |  |  |  |  |  |
| pH |  | 7.5 | 7.8 |  |  |  |  |  |  |  |  |  |  |
| calcium | $\mathrm{mg} / \mathrm{ICa}{ }^{2+}$ | 152 | 32 |  |  |  |  |  |  |  |  |  |  |
| magnesimm | $\mathrm{mg} / \mathrm{Mg}^{2+}$ | 82.62 | 14.5 |  |  |  |  |  |  |  |  |  |  |
| sodium | mg/I $\mathrm{Na}^{+}$ | 575 | 805 |  |  |  |  |  |  |  |  |  |  |
| total hardness | $\mathrm{mg} / \mathrm{CaCO}_{3}$ | 720 | 140 |  |  |  |  |  |  |  |  |  |  |
| iron | $\mathrm{mg} / \mathrm{IFe}$ | 0.1 | 0.15 |  |  |  |  |  |  |  |  |  |  |
| manganese | mg/ I Hs | 0.08 | 0.02 |  |  |  |  |  |  |  |  |  |  |
| fluoride | mg/ $/ \mathrm{F}^{-}$ | 0.5 | 3.0 |  |  |  |  |  |  |  |  |  |  |
| chloride | mag/I $\mathrm{Cl}^{-}$ | 820 | 760 |  |  |  |  |  |  |  | . |  |  |
| bicarbonate | $\mathrm{mg}_{\mathrm{I}} \mathrm{HCO}_{3}^{-}$ | 805.2 | 732 |  |  |  |  |  |  |  |  |  |  |
| nitrate | $\mathrm{mg} / \mathrm{I} \mathrm{HO}_{3}{ }^{-}$ | 39.43 | 31 |  |  |  |  |  |  |  |  |  |  |
| silphate | $\mathrm{mg} / \mathrm{I} \mathrm{SO} 4{ }^{2-}$ | 200 | 220 |  |  |  |  |  |  |  |  |  |  |
| phosphate (ortho) | $\mathrm{mg} / \mathrm{I} \mathrm{PO} 4{ }^{3-}$ | 1.3 | 0.4 |  |  |  |  |  |  |  |  |  |  |

Continued 5

| location borchole nr. date of sampling date of anslysis sample method |  | $\begin{aligned} & \text { Masimbu } \\ & \text { K-3 } \\ & \text { March } 79 \\ & \text { March } 79 \\ & \text { bailer } \end{aligned}$ | $\begin{aligned} & \text { Masimbu } \\ & \text { K-4 } \\ & \text { PFarch } 79 \\ & \text { March } 79 \\ & \text { bailer } \end{aligned}$ | $\begin{aligned} & \text { Mas imbu } \\ & \text { K-10 } \\ & 14-3-79 \\ & 20-3-79 \\ & \text { bailer } \end{aligned}$ | $\begin{aligned} & \text { Mas ímbu } \\ & \text { K-11 } \\ & 14-3-79 \\ & 20-3-70 \\ & \text { bailer } \end{aligned}$ | $\begin{aligned} & \text { Masimbu } \\ & \mathrm{K}-12 \\ & 14-3-79 \\ & 20-3-79 \\ & \text { bailer } \end{aligned}$ | bailer | bailer | bailer | bailer | bailer | bailer | bailer |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| Water amalysis | unit |  |  |  |  |  |  |  |  |  |  |  |  |
| clectrical conductivity | $\mathrm{ms} / \mathrm{m}\left(25^{\circ} \mathrm{C}\right)$ | 450 | 93 | 175 | 80 | 275 |  |  |  |  |  |  |  |
| pH |  | 6.8 | 7.3 | 6.8 | 7.1 | 7.0 |  |  |  |  |  |  |  |
| calciur | mg/L $\mathrm{Ca}^{2+}$ | 220 | 188 | 100 | 64 | 140 |  |  |  |  |  |  |  |
| magnesiam | $\mathrm{mg} / \mathrm{Hg}^{2+}$ | 182.25 | 36.45 | 97.2 | 48.6 | 103.28 |  |  |  |  |  |  |  |
| sodium | mg/I $\mathrm{Ha}^{+}$ | 414 |  | 138 |  | 207 |  |  |  |  |  |  |  |
| total hardness | $\mathrm{mg} / \mathrm{ICaCO}_{3}$ | 1300 | 400 | 650 | 360 | 775 |  |  |  |  |  |  |  |
| i ron | $\mathrm{mg} / \mathrm{T} \mathrm{Fe}$ | 0.5 | 0.32 | 1.0 | 0.14 | 0.6 |  |  |  |  |  |  |  |
| manganese | mg/I Mn | 9.2 | 2.2 | 2.5 | 0.6 | 0 |  |  |  |  |  |  |  |
| fluoride | mg/L $\mathrm{F}^{-}$ | 0.2 | 0.4 | 0.6 | 0.6 | 0.5 |  |  |  |  |  |  |  |
| chloride | $\mathrm{mg} / \mathrm{CCl}{ }^{-}$ | 1200 | 75 | 325 | 50 | 625 |  |  |  |  | . |  |  |
| bicarhonate | $\mathrm{mg} / \mathrm{HCO}_{3}{ }^{-}$ | 610 | 366 | 488 | 268 | 427 |  |  |  |  |  |  |  |
| nitrate | $\mathrm{mg} / \mathrm{I} \mathrm{HO}_{3}{ }^{-}$ | 8.86 | 4.43 | 3.1 | 22.2 | 4.43 |  |  |  |  |  |  |  |
| sulphate | $\mathrm{mg} / \mathrm{I} \mathrm{SO}_{4}{ }^{2-}$ | 4 | 24 | 100 | 10 | 35 |  |  |  |  |  |  |  |
| phosphate (ortho) | $\mathrm{mg} / \mathrm{IPO}_{4}{ }^{3-}$ | 7.6 | 1.8 | 0.6 | 0.4 | 0.9 |  |  |  |  |  |  |  |


| location borehole nr. date of sampling date of analysis sample method |  | Mkambarani <br> Lokobe River <br> Feb. 79 <br> Feb. 79 <br> bailer | Horogoro <br> Tennis Court <br> Feb. 79 <br> Feb. 79 <br> bailer | Morogar <br> Golf Court <br> Feb. 79 <br> Feb. 79 <br> bailer | bailer | bailer | bailer | bailer | bailer | bailer |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| Water analysis | unit |  |  |  |  |  |  |  |  |  |
| electrical conductivity | $\mathrm{mS} / \mathrm{mf}\left(25^{\circ} \mathrm{C}\right)$ | 520 | 190 | 182 |  |  |  |  |  |  |
| plt |  | 7.9 | 8.0 | 7.6 |  |  |  |  |  |  |
| calcium | $\mathrm{mg} / \mathrm{ICa}{ }^{2+}$ | 300 | 66 | 86 |  |  |  |  |  |  |
| magnesium | $\mathrm{mg} / \mathrm{Hg}^{2+}$ | 121.5 | 96 | 127.58 |  |  |  |  |  |  |
| sodium | mg/I ${ }^{\text {Na }}{ }^{+}$ | 1725 | 184 | 92 |  |  |  |  |  |  |
| total hardness | $\mathrm{mg} / \mathrm{ICaCO}_{3}$ | 1250 | 560 | 740 |  |  |  |  |  |  |
| iron | $\mathrm{mg} / \mathrm{IFe}$ | 0.2 | 0.03 | 0.01 |  |  |  |  |  |  |
| manganese | mg/ m H | 0.5 | 0.4 | 0.3 |  |  |  |  |  |  |
| fluoride | $\underline{m g / I ~} \mathrm{~F}^{-}$ | 0.2 | 1.3 | 1.2 |  |  |  |  | - |  |
| chloride | $\mathrm{mg} / \mathrm{I} \mathrm{Cl}^{-}$ | 1600 | 260 | 50 |  |  |  |  |  |  |
| bicarbonate | $\mathrm{mg} / \mathrm{IHCO} 3-$ | 335,5 | 732 | 1122.4 |  |  |  |  |  |  |
| nitrate | $\mathrm{mg} / \mathrm{L} \mathrm{HO}_{3}{ }^{-}$ | 2.66 | 8.86 | 0.89 |  |  |  |  |  |  |
| sulphate | $\mathrm{mg} / \mathrm{I} \mathrm{SO}{ }_{4}{ }^{2-}$ | 20 | 15 | 20 |  |  |  |  |  |  |
| phosphate (ortho) | $\mathrm{mg} / \mathrm{IPO}_{4}{ }^{3-}$ | 0.9 | 0.6 | 0.46 |  |  |  |  |  |  |

## DATA DD 11

THE GEO-ELECTRICAL SOUNDINGS IN THE BEREGA AREA AND THEIR INTERPRETATIONS











WELL LOGGINGS WITH CORRESPONDING LITHOLOGICAL PROFILES












DATA DD 13

FLUCTUATIONS OF GROUND WATER LEVELS AND EC'S, MEASURED IN PIEZOMETERS; BEREGA AREA

FLUCTUATIONS OF GROUNDWATERLEVELS AND EC'S
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## FLUCTUATIONS OF GROUNDWATERLEVELS AND EC'S

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